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# High-resolution nitrogen stable isotope sclerochronology of bivalve shell carbonate-bound organics

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#### 24 Abstract.

Nitrogen stable isotope ratios ( $\delta^{15}$ N) of organic material have successfully been used to track 25 food web dynamics, nitrogen baselines, pollution, and nitrogen cycling. Extending the  $\delta^{15}N$ 26 27 record back in time has not been straightforward due to a lack of suitable substrates in which 28  $\delta^{15}$ N records are faithfully preserved, thus sparking interest in utilizing skeletal carbonate-bound 29 organic matter (CBOM) in mollusks, corals, and foraminifera. Here we test if calcite Pecten 30 maximus shells from the Bay of Brest and the French continental shelf can be used as an archive of  $\delta^{15}$ N values over a large environmental gradient and at a high temporal resolution 31 (approximately weekly). Bulk CBOM  $\delta^{15}$ N values from the growing tip of shells collected over a 32 large nitrogen isotope gradient were strongly correlated with adductor muscle tissue  $\delta^{15}N$  values 33  $(R^2=0.99, n=6, p<0.0001)$ . We were able to achieve weekly resolution (on average) over the 34 35 growing season from sclerochronological profiles of three shells, which showed large seasonal 36 variations up to 3.4‰. However, there were also large inter-specimen differences (up to 2.5‰) between shells growing at the same time and location. Generally, high-resolution shell  $\delta^{15}N$  data 37 follow soft-tissue  $\delta^{15}$ N values, but soft-tissues integrate more time, hence soft-tissue data are 38 39 more time-averaged and smoothed. Museum-archived shells from the 1950s, 1965, and 1970s do not show a large difference in  $\delta^{15}$ N values through time despite expected increasing N loading to 40 41 the Bay over this time, which could be due to anthropogenic N sources with contrasting values. Compiling shell CBOM  $\delta^{15}$ N data from several studies suggests that the offset between soft-42 tissue and shell  $\delta^{15}$ N values ( $\Delta_{tissue-shell}$ ) differs between calcite and aragonite shells. We 43 44 hypothesize that this difference is caused by differences in amino acids used in constructing the 45 different minerals, which should be specific to the CaCO<sub>3</sub> polymorph being constructed. Future 46 work should use compound specific isotope analyses (CSIA) to test this hypothesis, and to

- 47 determine whether certain amino acids could specifically track N sources or possibly identify
- 48 amino acids that are more resistant to diagenesis in fossil shells. In conclusion, bivalve shell
- 49 CBOM  $\delta^{15}$ N values can be used in a similar manner to soft-tissue  $\delta^{15}$ N values, and can track
- 50 various biogeochemical events at a very high-resolution.
- 51

## 52 1.0 Introduction

53 Nitrogen stable isotope signatures ( $\delta^{15}$ N) of organic matter are a powerful tool for studying food 54 webs and tracking nitrogen dynamics in terrestrial and aquatic systems (Fry 1988; Cabana and Rasmussen 1996: Cole et al., 2011). Nitrogen in consumers is usually enriched in <sup>15</sup>N relative to 55 56 their diet, typically by +2% to +4%, or on average 3.4\%, which allows estimations of trophic 57 positions of consumers relative to the base of the food web (De Niro and Epstein, 1981; 58 Minigawa and Wada, 1984; Vander Zanden and Rasmussen 2001; Post 2002; Caut et al. 2009). 59 It can however be difficult to estimate the isotopic composition of N sources at the base of the 60 food web (isotopic baseline) since ecosystem biogeochemistry is dynamic and multiple nutrient 61 sources available to phytoplankton can have distinct stable isotope signatures that vary 62 temporally and spatially (McMahon et al. 2013). Bivalve soft-tissues have been proposed as 63 good proxies of this isotopic baseline as they are sessile and integrate this variability (Jennings and Warr 2003, Vokhshoori and McCarthy 2014). Natural variations in  $\delta^{15}$ N values of 64 65 particulate N caused, for example, by upwelling (e.g., Mollier-Vogel et al. 2012) could thus be preserved in tissues, potentially serving as an upwelling / El Niño Southern Oscillation proxy in 66 certain locations. In addition,  $\delta^{15}$ N values of organics can be used as a wastewater pollution 67 indicator (e.g., Costanzo et al 2005). Processed wastewaters are  $^{15}$ N enriched, with  $\delta^{15}$ N values 68 of particulate N typically around +15‰ (Heaton, 1986), but much higher values have been 69 70 recorded (e.g., Schlacher et al., 2007; Marwick et al., 2014; see Bouillon et al., 2012, for a 71 review).

72

Fixed the extending organic  $\delta^{15}$ N records back through time to develop isotopic baselines, or gather data on past pollution events, is challenging. Sediment  $\delta^{15}$ N records can be used to track relative

changes, but sediment trap and surface sedimentary  $\delta^{15}$ N values suggest alteration during early 75 76 burial, cautioning against the use of sediment cores for reconstructing isotope baselines 77 (reviewed in Robinson et al., 2012). Although there are excellent archives of preserved organic 78 material in museums (e.g., animal soft-tissues), the effect of long-term preservation in formalin 79 and/ or ethanol are not well characterized. Delong and Thorp (2009) recorded a -0.2 ‰ shift in freshwater mussel tissue  $\delta^{15}$ N values after 12 months in ethanol. On the other hand, Carabel et al. 80 (2009) reported a positive shift in  $\delta^{15}$ N values of bivalve tissues of about +1 ‰ after storage in 81 ethanol for two years. While these are not large effects considering the strong <sup>15</sup>N enrichment 82 83 associated with anthropogenic N loading, it is not clear if longer time periods would result in 84 larger isotopic shifts. For example, Versteegh et al. (2011) found more than a 5‰ difference in shell organic  $\delta^{15}$ N values between shells stored dry and shells stored in ethanol for 73 years. 85 Clearly, dry stored specimens would be the safer option for extending  $\delta^{15}N$  data back in time. 86 87 Typically, shells are stored dry in museum collections making them ideal archives for reconstructing past  $\delta^{15}$ N values. Moreover, pristine unaltered fossils may also provide insights 88 89 into the nitrogen cycle into the geological past (e.g., O'Donnell et al., 2003, 2007).

90

91 The study of nitrogen isotopes in carbonate bound organics has recently received renewed

92 attention. Studies investigating N in carbonate bound organics in foraminifera (Ren et al.,

93 2012a,b), corals (Marion et al. 2005; Williams and Grottoli, 2010; Yamazaki et al., 2011a, b,

2013; Wang et al., 2015), fish otoliths (Vandermyde and Whitledge, 2008; Rowell et al, 2010;

95 Grønkjær et al., 2013), and bivalves (O'Donnell et al., 2007; Carmichael et al., 2008; Watanabe

96 et al., 2009; Kovacs et al., 2010; Versteegh et al., 2011; Dreier et al., 2012; Graniero et al., 2016)

97 have been increasing. Continuous long-term (>150 years) N isotope records have been developed

98 from coral skeletons (Erler et al., 2015), and foraminiferal N isotope records have been extended 99 back 30 Ka suggesting lower N fixation in the Atlantic during the last ice age (Ren et al. 2009), 100 both illustrating the power of this technique. Bivalves have a global distribution in many 101 environments, generally withstand pollution, are common, and are not mobile over large 102 distances allowing for spatial reconstructions. In addition, the dense 'closed' shells of bivalves 103 (Marin et al., 2007) exclude foreign organic material and make them relative diagenically 104 resistant (c.f., Engel et al., 1994). Nevertheless, both corals and mollusk shells have been shown 105 to maintain their carbonate bound organics for hundreds to thousands of years (Engel et al 1994; 106 Ingalls et al., 2003). Therefore, similar to corals and foraminifera, bivalve shell carbonate is a 107 suitable structure to preserve high-resolution carbonate bound N. 108 109 As previously noted, bivalves are also a good substrate to target for N isotope studies because they are low-level consumers and therefore record baseline  $\delta^{15}N$  values (Jennings and Warr 110 111 2003). Finally, bivalves can be long-lived and have high growth rates allowing both high-112 resolution environmental reconstruction (down to daily) and provide records extending back in 113 time (e.g., Arctica islandica shell chronologies have been extended back more than 1300 years; 114 Butler et al., 2013).

115

Although bivalve shells have higher %N than corals and foraminifera, they typically have only a few percent organic matter bound within the carbonate (typically 1 to 5% organic matter; Marin and Luquet 2004). This presents analytical challenges when determining such small amounts of nitrogen in a large carbonate matrix, especially in a sclerochronological context. To circumvent this problem, several studies have removed the carbonate via acidification (e.g., Carmichael et

121	al., 2008; Watanabe et al., 2009; Kovacs et al., 2010). However, several studies have shown that
122	acidifying organic matter with high CaCO <sub>3</sub> content alters the $\delta^{15}$ N value (Jacob et al., 2005; Ng
123	et al., 2007; Mateo et al., 2008; Serrano et al., 2008), likely because acidification removes all or
124	part of the acid soluble N, leading to analysis of an unknown part of the bulk organic N. This has
125	prompted other groups to directly combust the bulk shell to release the entire organic N pool,
126	with good results (e.g., Vandermyde and Whitledge, 2008; Rowell et al, 2010; Versteegh et al.,
127	2011; Graniero et al., 2016).
128	
129	The aims of this study are 1) illustrate that low %N samples can be accurately measured on a
130	standard elemental analyzer - isotope ratio mass spectrometer (EA-IRMS) configuration, 2)
131	determine if bivalve shells can be used as an archive of $\delta^{15}N$ values, providing the same or
132	parallel information as soft-tissues, and 3) determine if shells can provide an ultra-high-
133	resolution $\delta^{15}$ N record (i.e., at a daily or weekly resolution).
134	

#### 135 **2.0 Materials and Methods**

136 2.1 Shell collection

This study presents data from three sets of *Pecten maximus* shell samples collected from the
French coast: shells collected along a depth gradient in 2008, shells from the Bay of Brest
collected in 2000, and three archived shells collected from the Bay over the past century. Shells
were collected along a transect from the Bay of Brest, France (40 m depth, 4° 40' W, 48° 18' N),
to the edge of the continental shelf (220 m depth, 8° 15' W, 48° 12' N) in 2008 (Fig 1; shell
collection details and other data are provided in Nerot et al., 2012). Six shells from different
depths (one per depth at 40, 78, 120, 140, 153, and 190 m) were analyzed along their ventral

margin covering the most recent one to five years of growth. Soft-tissue  $\delta^{15}$ N values from the 144 145 adductor muscle (including those from the shells analyzed in this study) are from Nerot et al. 146 (2012) and range from 2.0 to 10.4‰ (n = 95). Three additional shells collected in the year 2000 147 from 30 m depth in the Bay were analyzed over their last year of growth with a resolution of 2 to 29 days per sample (average =  $7.4 \pm 6.4$  days, median = 5 days). Monthly soft-tissue  $\delta^{15}$ N values 148 149 from the adductor mussel and digestive gland collected over the 2000 calendar year are taken 150 from Lorrain et al. (2002) and represent average and standard deviations of five individuals all 151 collected at the same time and site as the shells. Shells were selected from those collected on 152 December 14, 2000; the last collection date of the Lorrain et al. (2002) study. Finally, archived 153 shells from the 1950s, 1965 and 1970s (one shell per time period, collected in the Bay; stored dry 154 without tissues, each covering about four years of growth) were also analyzed.

155

## 156 2.2 Shell cleaning and sampling

157 Pecten maximus produce dense, non-porous shells of foliated calcite microstructure without 158 pockets or chalky layers. Shell exteriors were scrubbed with a hard brush and cleaned with a 159 weak acid (acetic) for a few seconds to remove any extraneous particles and the periostracum 160 and to expose the underlying carbonate. Carbonate powder was milled from the exterior of the 161 cleaned shell surface following the daily growth striae this species produces using a hand held 162 drill. These striae, along with the known date of collection, allowed an absolute chronology to be 163 developed by counting back from the collection date (see Chauvaud et al., 1998; Lorrain et al., 164 2004; Gillikin et al., 2008 as examples). Milling followed the striae along the curve of the shell, 165 only sampling the bottom of the rays (Fig 2). Care was taken to only sample the outer shell layer. 166 This sampling produced about 5 mg (5.47  $\pm$  0.56 mg) of calcite powder, corresponding to ~5  $\mu$ g

167 N (see Results) and resulting in a peak area of  $8.5 \pm 2.8$  Volt-seconds (Vs) on the IRMS 168 (equivalent to an amplitude of  $\sim 200 \text{ mV}$  on mass 28). While the elemental analyzer can easily 169 accommodate 40 mg, the small sample size was necessary to reduce time averaging required in a 170 sclerochronological study. To double-check the reproducibility of this method, we resampled one 171 of the shells collected in the year 2000 (shell B). Samples were milled between the samples taken 172 during the first sampling, and were analyzed on a different date. These can be considered 173 replicates, but also will contain real differences as they are effectively different samples and 174 represent different time. Furthermore, an additional shell was used to test reproducibility. A 175 larger section of this shell was milled and homogenized, from this, 7 samples were analyzed and were highly reproducible ( $\delta^{15}N = 8.98 \pm 0.27$  ‰; %N = 0.07 ± 0.002%). 176

177

178 As noted earlier, bivalves produce dense non-porous skeletons and therefore likely do not require 179 the aggressive oxidative cleaning required for corals and foraminifera. In contrast, porous coral 180 skeletons are known to contain foreign organic N from endolithic algae and fungi (Bentis et al., 2000) as well as humic acids (Susic et al., 1991), which can alter carbonate bound organic  $\delta^{15}N$ 181 182 values therefore necessitating an oxidative cleaning step (Erler et al., 2016). In addition to thin 183 porous skeletons, foraminifera also have empty internal chambers, which can fill with sediments, 184 also necessitating aggressive cleaning procedures (Ren et al., 2009). Bivalve shell organics on 185 the other hand can be considered a "closed system" (Marin et al., 2007). In addition, foliated 186 calcite bivalve shell microstructure is typically rather compact (in the absence of pockets or 187 chalky layers such as in oysters) and is probably rather diagenetically stable, especially when 188 compared with porous for a corals. We tested this hypothesis on another shell by 189 analyzing growth lines from half the shell cleaned as mentioned above, and half the shell soaked

190 in a strong sodium hypochlorite solution (reagent grade, 10–15%) for 24 hours. Four samples 191 were milled from each half of the shell taking care to match timelines across shell halves (each 192 sequential sample moved forward in time). Differences between bleached and unbleached halves 193 ranged from 0.04 to 0.88 ‰ (Fig. 3). A paired T-test showed there is no statistical difference 194 between the bleached and unbleached halves (p = 0.31), and that the mean difference (0.34 ‰) is 195 less than expected analytical precision (0.5 ‰, see further). Moreover, both the bleached and unbleached halves both show a decrease in  $\delta^{15}$ N values along the shell (though time). The small 196 197 differences between the bleached and unbleached halves likely represent variable loss of acid 198 soluble N along the shell.

199

## 200 2.3 Analytical methods and validation

201 To test if our IRMS was capable of analyzing low N samples, various masses of the IAEA-N1 standard (ammonium sulfate;  $\delta^{15}N = 0.4 \pm 0.2\%$ ) were analyzed. To facilitate preparing 202 203 standards with very low amounts of N, finely powdered IAEA-N1 material was mixed with 204 IAEA-CH6 sucrose standard (which does not contain N). Standards were weighed into standard 205 tin cups and combusted on a standard EA-IRMS setup (Thermo Flash HT with a Costech zero 206 blank autosampler, coupled to a Thermo DeltaV Advantage via a Thermo ConFlo IV) at KU 207 Leuven, Belgium. No modifications were made to the standard operating procedure (e.g., 208 combustion column and reduction column at 1070 °C and 640 °C respectfully; an inline CO<sub>2</sub> trap 209 (soda lime/ Carbosorb) was installed prior to the H<sub>2</sub>O trap). The various masses of IAEA-N1 210 used contained between 2.4 and 41.1 µg N. Versteegh et al. (2011) have shown that large masses 211 of CaCO<sub>3</sub> do not interfere with low N samples combusted in this manner on a similar

instrumental setup. Percent N was determined on a subset of samples using acetanilide as astandard (10.36% N).

215	Thirty-three IAEA-N1 standards were analyzed and data were organized into size class bins (Fig.
216	4). The lower N samples had poorer accuracy ( $1\sigma = -0.5\%$ compared to $-0.2\%$ ), but this was
217	not considered problematic considering the large variations expected in the shell $\delta^{15}N$ values.
218	Most of the shell samples produced peaks with an area around 8.5 Vs, which considering the
219	IAEA-N1 data should have an error of less than $\pm 0.5$ ‰ (Fig. 4). Empty tin cups produced very
220	small peaks with an amplitude of ~2 mV or less, i.e. 30 times smaller than our smallest IAEA-N1
221	peak. We therefore consider $\pm 0.5$ ‰ as a representative uncertainty for all shell $\delta^{15}$ N data, which
222	is conservative considering the reproducibility of 7 homogenized samples noted above ( $\pm 0.27$
223	‰). Shell N concentrations ranged from 0.06 to 0.12 % with an average of 0.09 $\pm$ 0.02 % (N=18,
224	collected from two shells). The samples milled adjacent to the original samples in the shell
225	collected in 2000 used to test reproducibility (shell B) had an average difference of $0.2 \pm 0.5\%$
226	(n=9) and ranged from 0.0‰ to 1.2‰ different. Although one sample was 1.2‰ different, it was
227	only 0.1‰ off from the preceding sample (see further), thus the samples follow the same pattern
228	as the first set and the new data do not increase the variability in the data. Therefore, the
229	reproducibility is similar to that reported based on the IAEA-N1 data shown in Fig. 4.

# **3.0 Results**

*3.1 Depth-transect shells* 

233 Carbonate bound organics in shells collected along the depth transect exhibited a wide range of 234  $\delta^{15}$ N values, 1.0‰ to 10.3‰, with a general trend of lower  $\delta^{15}$ N values with increasing depth

235 (Fig 5). The range of values in each shell (intra-shell variability) was highest in the shells from 236 the deepest sites (both = 4.2%) and lower in shells from shallower sites (1.3% to 2.1%) (Fig. 5). Shell  $\delta^{15}$ N values were similar to soft-tissue (adductor muscle)  $\delta^{15}$ N values, with the data from 237 238 the shell edge (representing the most recent time) offering the best correlation with tissue data 239 for shallow water specimens, and the average shell data better correlating for deeper water 240 specimens (Fig 6). Simple linear regression resulted in strong correlations between soft-tissue  $\delta^{15}N$  values and both growing shell tip  $\delta^{15}N$  values (Fig. 7; R<sup>2</sup>=0.992; p<0.0001; n = 6) and 241 average shell  $\delta^{15}$ N values (R<sup>2</sup>=0.792, p=0.017; n = 6) (Fig. 7). Including data from the three 242 243 shells collected in 2000 in the regression statistics does not statistically change the goodness of fit (growing shell tip: shell  $\delta^{15}N = 0.62 \pm 0.03 * \text{tissue } \delta^{15}N + 3.45 \pm 0.27$ , R<sup>2</sup> = 0.98, p < 0.0001; 244 average shell: shell  $\delta^{15}N = 0.80 \pm 0.13$  \* tissue  $\delta^{15}N + 2.115 \pm 0.98$  (intercept is not significantly 245 246 different from 0, p = 0.07),  $R^2 = 0.847$ , p = 0.0004; N = 9 shells for both; standard errors 247 included in regression equations; shell tip slopes (p = 0.51) and intercepts (p = 0.71) were not 248 different, and average shell slopes (p = 0.63) and intercepts were not different (p = 0.74) when 249 the three additional shells were included in the regression).

250

# 251 3.2 High-resolution shell $\delta^{15}N$ profiles

The three shells collected from the Bay of Brest in 2000 and sampled at high resolution show a similar  $\delta^{15}$ N pattern, with a high value in late March/ early April, a low around late April/ early May, rising to a high value in late Spring/ early Summer, then decreasing to the end of the growth year (Fig. 8). However, there are large differences between individual shells, up to ~2.5‰ at times, whereas at other times data from different shells converge. The shell  $\delta^{15}$ N pattern roughly tracks the digestive gland offset by +3.5‰ pattern (the digestive gland was emptied prior to analysis and should not contain food; Lorrain et al., 2002). The shell tip δ<sup>15</sup>N
values versus November adductor muscle tissue δ<sup>15</sup>N values plot close to the regression line of
the shells collected along the depth transect (Fig. 7). *3.3 Archived shells*The shells from the 1950s, 1965 and 1970s did not show markedly different δ<sup>15</sup>N values when
compared to shells collected more recently (2000, 2008) in the Bay of Brest (Fig. 9). The 1950s
shell has slightly higher δ<sup>15</sup>N values compared to the other shells (ANOVA post-hoc Tukey HSD

Test p<0.01), but overlaps with values from the more recent shells (see Fig. 9).

267

#### 268 **4.0 Discussion**

These data illustrate that shells can be used as archives for nitrogen isotope studies (Fig. 7). The strong correlation between shell and soft-tissue  $\delta^{15}$ N values show that shells can be utilized in a similar manner as soft-tissues, which have been extensively utilized to trace pollution, food webs, metabolism, and developing baseline isoscapes (see Fry 1988, 1999; Lorrain et al., 2002; Paulet et al., 2006; Carmichael et al., 2012; Nerot et al., 2012; Vokhshoori and McCarthy, 2014).

274

275 4.1 Inter- and intra-shell variability

276 *Pecten maximus* shells collected over a large environmental range of  $\delta^{15}$ N values strongly

277 correlate with soft-tissues. Individuals that grew in the same area and time can exhibit large

- 278 offsets in  $\delta^{15}$ N profiles, yet exhibit similar intra-annual trends (e.g., up to 2.5% between shell B
- and C in early September; Fig. 8). The differences between the high-resolution shell  $\delta^{15}N$
- 280 profiles may be caused by several factors. Nitrogen in consumers is typically enriched in <sup>15</sup>N by

281 about 3.4‰ (De Niro and Epstein, 1981; Minigawa and Wada, 1984; Vander Zanden and 282 Rasmussen 2001). However, it is well known that N isotope fractionation is a complex process 283 and other factors aside from just trophic enrichment also play a role. Therefore, there is often a 284 large variability in the trophic enrichment factor with standard deviations of more than 1‰ being 285 common (e.g., Post, 2002; Bouillon et al., 2012). Factors such as condition index (animal health), 286 age, food availability, and food quality affect N fractionation (Minagawa and Wada. 1984; 287 Adams and Sterner, 2000; Overman and Parrish, 2001; see Fry, 2006, Caut et al., 2009, and 288 Bouillon et al., 2012, for review). We can rule out food and age as these animals (shells A, B, 289 and C) were growing in the same area and were all the same age class (see Lorrain et al., 2002). The trophic <sup>15</sup>N enrichment between a consumer and its food is due to the light isotope, <sup>14</sup>N, 290 291 reacting faster during amino acid deamination thereby enriching <sup>15</sup>N in the tissues (Macko et al., 292 1986). Variations in excretion and nutrient assimilation caused by, for example, condition index 293 or metabolic rate, can therefore result in differences between co-occurring animals (see Fry, 294 2006 for a discussion on N excretion and fractionation), which may explain the differences we 295 see between individuals in Fig. 8.

296

Soft-tissues also show variability between individuals (Fig. 6), albeit typically less than seen in
these shells (Fig. 8), but soft-tissues integrate more time. Soft-tissues are constantly degrading
and are replaced with newly synthesized components in a state of dynamic equilibrium (Bender,
1975), thereby updating the isotopic signature of the tissues (e.g., Tieszen, 1978; Tieszen et al.,
1983). In general, tissues that are more metabolically active have faster turn-over rates
(Thompson and Ballou, 1956; Libby et al., 1964; Paulet et al., 2006). The shells do not have any
turn-over and once calcified, the isotope signature is locked in. Sampling of the shells time-

304 averaged about one week of growth (on average; due to shell sampling resolution), whereas the 305 adductor muscle, for example, can integrate several months to years of time due to its slow 306 metabolic turn-over (Lorrain et al. 2002; Paulet et al., 2006; Poulain et al., 2010). Soft-tissue 307  $\delta^{15}$ N values in adductor muscle tissues of this scallop population varied up to 0.8% between co-308 occurring individuals (at the end of March 2000) and up to 1.8% in the gonad (at the beginning 309 of May 2000) (Lorrain et al., 2002). Pecten maximus adductor muscle tissue has a slower 310 metabolic turn-over than gonad tissue (Paulet et al., 2006), so the adductor muscles are more 311 time-averaged thereby reducing any high-frequency variations. Shells time-average very little 312 time (<1 week), so larger differences between individuals (up to 2.5% in our study) are not 313 unexpected. Moreover, the average of high-resolution data (mean of shell A, B and C) shown in 314 Fig. 7 show much less variability between shells (range = 0.4%), further illustrating the effect of 315 time-averaging on data variability. This can explain why the soft-tissues exhibit much less 316 temporal variability than shells; the soft-tissues average much more time than the shell samples. 317

318 In addition to variability between shells, there are also large seasonal differences in the high-319 resolution profiles, with up to 3.4‰ variations between April and June for shells B and C (Fig. 320 8). Lorrain et al. (2002) proposed that differences in metabolism and energy allocation have a strong impact on  $\delta^{15}$ N values in soft-tissues. For example, large seasonal differences were 321 observed in the digestive gland  $\delta^{15}$ N values that were attributed to a shift between utilization of 322 323 reserves from muscles during the winter to directly from food during phytoplankton events in the spring. Similarly, seasonal variations in adductor muscle  $\delta^{15}$ N values were attributed to changes 324 325 in energy allocation to reproduction. Seasonal changes in energy allocation and metabolism

326 likely also results in seasonal changes in shell  $\delta^{15}$ N values. The analysis of carbonate bound 327 organics therefore opens the possibility to study rapid changes in the metabolism of bivalves. 328

329 Despite the observed variations, all three shells analyzed at high-resolution do exhibit similar patterns in  $\delta^{15}$ N values (Fig. 8). Moreover, shell  $\delta^{15}$ N values are similar to adductor muscle and 330 digestive gland (with a +3.5 offset)  $\delta^{15}$ N values, which are smoothed due to slower turn-over 331 rates (Fig. 8). In addition, when the  $\delta^{15}$ N values from the growing tips of the shells sampled at 332 high resolution are plotted against the  $\delta^{15}$ N values of the adductor muscles at the time of 333 334 collection, the data lie along the same regression line from the depth transect in Fig. 7. The correlation between soft-tissue and shell  $\delta^{15}$ N values suggests that if a sudden large shift in food 335  $\delta^{15}$ N values (i.e., a baseline shift) occurred, for example due to a wastewater pollution event, it 336 may be recorded in high-resolution shell  $\delta^{15}$ N values offering the possibility to precisely date 337 338 such an event. This could prove useful to determine the exact timing and duration of coastal 339 pollution events.

340

## 341 4.2 The shell and soft-tissue $\delta^{15}N$ record of a large nitrogen isotope gradient

342 The clear decrease seen in both the soft-tissue and shell  $\delta^{15}$ N data from the Bay to the edge of the

343 shelf (Fig. 4) is likely caused by less anthropogenic N loading farther from land (see Nerot et al.,

- 344 2012). However, as noted by Nerot et al. (2012), the  $\delta^{15}$ N values of soft-tissues were
- unexpectedly low in the deeper waters. Considering the 3.5‰ trophic shift, data from deeper
- scallops suggest a food source with a  $\delta^{15}$ N value around 0‰, which is highly unlikely in this
- region. Nerot et al. (2012) hypothesized that these animals have low metabolic rates due to
- 348 constant cold-water temperatures coupled with both low food supply and poor food quality (see

also Nerot et al., 2015). Our shell data support this hypothesis, but illustrate that the  $\delta^{15}$ N values 349 350 vary through time, perhaps seasonally. Data from the deepest shell ranged from 1.0 to 5.2% 351 (Fig. 5). This variability suggests that there are variations in food supply and quality over time, as well as variations in metabolic activity. As explained above, seasonal variations in shell  $\delta^{15}N$ 352 353 values can be attributed to variations in metabolic activity or energy allocation. The large 354 seasonal differences observed at the deepest sites (as compared to coastal stations) also suggests 355 higher variability in metabolic rates in animals deeper than 150 m water depth, probably due to 356 the sporadic nature of access to food resources.

357

358 Metabolic rates and time-averaging can also be used to explain why shells do not plot on the 1:1 359 line with tissues at the time of collection (Fig. 7). It is possible that during the time of shell 360 collection, metabolic rates in the animals from deeper waters were higher than other times of the year, and therefore there was more N isotope fractionation leading to higher  $\delta^{15}$ N values in the 361 362 shell tip (Fig. 6). The shell tips time-average the recent conditions (~about 1 week) whereas the 363 adductor muscle time-averages several months or longer. This is evident in the deepest shell, 364 where the average shell matches tissue data better than the shell tip (Fig. 6). This illustrates a 365 clear advantage of high-resolution shell sampling; it allows time resolved N isotope data, which 366 provide insight into environmental and biological mechanisms. Nevertheless, it should be noted 367 that there are no a priori expectations that the two tissues (muscle and CBOM) should plot on a 368 1:1 line, as other tissues also do not plot on the 1:1 line, such as muscle and digestive gland 369 tissues (see Fig. 8).

370

371 *4.3 Archived shells and pollution history* 

372 Despite the potential effects of diagenesis on very old shells, dry-stored museum archived shells 373 should remain pristine (discussed in the Introduction; also see Versteegh et al., 2011). In this study, shells collected from the Bay of Brest between the 1950s and 1970s show  $\delta^{15}N$  values 374 375 within the same range of the shells collected in 2000 and 2008, with slightly higher values in the 376 shell collected in the 1950s (Fig. 9). This suggests the Bay would not have undergone significant 377 change over this time period regarding anthropogenic N loading. This corroborates an earlier 378 study that showed there is no indication of any long-term trends in nutrient concentrations 379 between the mid-1970s and early 1990's in the Bay of Brest (see Chauvaud et al., 2000). 380 However, it is known that the sources of N loading to the Bay have changed, with increases in 381 both N from inorganic fertilizer and animal manure (Sebillotte, 1989; Peyraud, et al., 2012), but 382 this is likely rapidly removed due to strong tidal pumping and heavy biological N draw down 383 (Treguer and Queguiner, 1989; Chauvaud et al., 2000). Also, the N isotopic signature of 384 synthetic fertilizer is between +2 and -2‰ because it is synthesized from air via the Haber-Bosch method, whereas pig waste is expected to have higher  $\delta^{15}$ N values up to +10 or +20‰ (Heaton, 385 386 1986). Therefore, it is possible the N-isotope baseline did not significantly shift because the 387 increase in N loading includes N sources with both low (inorganic fertilizer) and high (pig waste)  $\delta^{15}$ N values. 388

389

## 390 4.4 Potential effects of biomineralization on bulk shell N fractionation

391 Overall, shell  $\delta^{15}$ N values were similar to, or higher than, soft-tissue  $\delta^{15}$ N values (Figs. 6 and 8). 392 A survey of literature presenting coupled shell and soft-tissue  $\delta^{15}$ N data illustrates that many 393 species with calcite shells also exhibit this pattern, whereas most studies of aragonite shells show 394 the opposite, i.e. lower  $\delta^{15}$ N values in the shells than in soft-tissues (Table 1). Considering the 395 caveat of averaging different amounts of time in soft-tissues (with relatively slow turn-over 396 rates) and shells (with time averaging dependent on sampling resolution and shell growth rate), it can be generalized that carbonate bound organic  $\delta^{15}$ N values from calcite shells match soft-tissue 397  $\delta^{15}$ N values, whereas carbonate bound organic  $\delta^{15}$ N values from aragonite shells are closer to 398  $\delta^{15}$ N values of particulate organic matter (i.e., the bivalve's food). It has been shown that calcite 399 400 prisms and aragonitic nacre are built using very different proteins (Marie et al., 2012); therefore 401 it can be assumed that calcite shells and aragonite shells also use a different set of proteins. It is 402 also well known that different proteins or amino acids (the building blocks of proteins) have largely different  $\delta^{15}$ N values (McClelland and Montoya, 2002; Vokhshoori and McCarthy, 403 404 2014). Therefore, we hypothesize that the differences observed between calcite and aragonite soft-tissue / shell  $\delta^{15}$ N values ( $\Delta_{tissue-shell}$ ) could be the result of different amino acid groups 405 406 present in the different mineral types, and that these different amino acids are somewhat 407 common in shells with the same CaCO<sub>3</sub> polymorph.

408

## 409 4.5 Suggestions for future research: CSIA

Future work on nitrogen isotopes in carbonate bound organics should investigate the  $\delta^{15}N$  values 410 411 of specific amino acids for several reasons. First, this could help develop the biomineralization 412 hypothesis we propose above in Section 4.4. Second, compound specific isotope analysis (CSIA) could also pinpoint amino acids that specifically record baseline  $\delta^{15}$ N values (cf., Vokhshoori 413 414 and McCarthy, 2014; McMahon et al., 2015). Finally, CSIA could potentially reveal amino acids 415 that are more resistant to diagenesis. Studies have shown that on time scales of 10,000 years to 416 millions of years, carbonate bound organic material is subject to diagenesis depending on burial 417 history, shell thickness, and mineralogy (Robbins and Ostrom, 1995; Risk et al., 1996).

418 However, other studies have shown some amino acids to be stable on century time scales in

419 porous coral skeletons (Goodfriend et al., 1992; Ingalls et al., 2003), and even more than 100,000

420 year time scales in well-preserved mollusk shells (Engel et al., 1994).

421

422 5.0 Conclusions

423 In conclusion, we show that simple combustion of shell material in an EA permits very low %N 424 samples to be easily analyzed with minimal processing, allowing increased temporal resolution 425 up to approximately one week. Our high-resolution results suggest that metabolism and/or energy allocation can affect shell  $\delta^{15}$ N values (as it does soft-tissue  $\delta^{15}$ N values), but that 426 seasonal and spatial variation shows good correlation with soft-tissue  $\delta^{15}$ N values. These 427 carbonate bound organic  $\delta^{15}$ N values would then be particularly useful to track both modern and 428 past isotope spatial gradients (isoscapes of  $\delta^{15}$ N values) across various systems, monitoring 429 430 nitrogen regime shifts due to pollution or natural N cycle processes. In addition, other natural 431 variations in particulate  $\delta^{15}$ N values caused, for example, by upwelling, could also be preserved 432 in shell organics, potentially serving as an upwelling / El Niño Southern Oscillation proxy. Overall, carbonate-bound organic matter is a good archive of soft-tissue  $\delta^{15}N$  values and offers 433 the possibility to examine  $\delta^{15}$ N shifts at high-resolution. 434

435

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448	
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683	<b>Table 1.</b> Differences between soft-tissue and carbonate bound organics ( $\Delta_{tissue-shell}$ ) and
684	shell mineralogy (C = calcite, A = aragonite) for 11 bivalve species.

685

Species	Mineralogy	$\Delta_{ ext{tissue-shell}}$ (‰)	Source
Pecten maximus	С	-1.3	This study
Mytilus edulis	С	-2.3	Versteegh et al. (2011)
Crassostrea virginica	С	-0.2	Kovacs et al. (2010)
Mytilus edulis	C+A?	-0.1*	LeBlanc (1989)
Pinctada imbricata	C+A?	~0.0*	Graniero et al. (2016)
Isognomon alatus	C+A?	~0.0*	Graniero et al. (2016)
Brachidontes exustus	C+A?	~0.0*	Graniero et al. (2016)
Arctica islandica	А	+2.7	LeBlanc (1989)
Ruditapes philippinarum	А	+1.1	Wananabe et al. (2009)
Mercenaria mercenaria	А	+2.4	Carmichael et al. (2008)
Mercenaria mercenaria	А	+1.0	O'Donnell et al. (2003)
Chambardia wissmanni	А	+3.0	Gillikin et al. (2012)
Venerupis aurea	А	+6.2	Dreier et al. (2012)

To be consistent, adductor muscle values were used when  $\delta^{15}$ N values were available from

687 multiple tissues. Only non-transplanted specimens were used from Kovacs et al. (2010)

because the muscle tissues of transplanted animals may have not yet equilibrated with the

689 new environment. All species are estuarine or marine except for *C. wissmanni*, which is an

690 African tropical freshwater mussel. \*These studies may have mixed aragonite and calcite

691 layers, but this is not certain.

### 692 **Figure legends:**

693 <u>Figure 1</u>. Map of study area in northern Bay of Biscay with shell collection sites (black circles)
694 and collection depth. Bathymetry contours are in meters.

695 <u>Figure 2</u>. Schematic of shell sampling. Fine grey lines represent daily striae and dark lines are

696 winter lines indicating growth cessation. Short thick lines represent areas milled for CBOM

samples. A) lower resolution sampling of older shells along depth transect. B) High-resolution

698 sampling of younger shells from the Bay of Brest (collected in 2000).

Figure 3. Comparison of  $\delta^{15}$ N data measured on bleached and unbleached halves of the same

growth striae from the same *Pecten maximus* valve. Growth and time is from sample 1 to 4 (left

701 to right). Error bars represent  $\pm 0.5\%$  (see text).

702 Figure 4. Average and standard deviations  $(1\sigma)$  for various sized IAEA-N1 standards. Values are

binned into ranges of nitrogen mass (x-axis). Standard deviations  $(1\sigma)$  are also listed below each

data point. The average and standard deviation for N1 reported by the IAEA is also shown for

comparison (shown as IAEA on x-axis). Number of standards analyzed from low to high N is 10,

706 10, 9, and 4. Peak amplitude for mass 28 (in mV) ranged from 60 to 100, 125 to 220, 300 to 425,

800 to 1200, whereas peak area (in Vs) ranged from 2 to 5, 5 to 10, 10 to 20, and > 20 from low
N to high N.

709 <u>Figure 5</u>. Shell  $\delta^{15}$ N values for samples milled from shells collected at different depths (one

shell per depth). The white symbols represent the sample taken from the tip of the shell, so the

most recent shell material, and black symbols represent time earlier in the growth year. Error bar

512 shows  $\pm 0.5\%$  (see text).

713 Figure 6. Average and shell tip  $\delta^{15}$ N values plotted against soft-tissue  $\delta^{15}$ N values from *P*.

714 *maximus* collected along the same transect. Soft-tissue  $\delta^{15}$ N values from Nerot et al. (2012).

715	<u>Figure 7</u> . Simple linear regression of average (grey) and shell tip (black) $\delta^{15}$ N values plotted
716	against soft-tissue $\delta^{15}$ N values (standard errors included). Also shown are the averages of the
717	shells sampled at high resolution (small black symbols). The grey dashed line is the 1:1 line.
718	<u>Figure 8</u> . High-resolution shell $\delta^{15}$ N values plotted against growth date (3 shells). Also shown
719	are $\delta^{15}$ N values of adductor muscle and digestive gland. For easier comparison, 3.5‰ was added
720	to the digestive gland data (shaded circles). Thin grey line represents daily shell growth rate
721	based on growth striae. Soft-tissue and shell growth data from Lorrain et al. (2002). Open
722	symbols of Shell B are additional (replicate) samples analyzed on a different day (see text).
723	<u>Figure 9</u> . Shell $\delta^{15}$ N values from museum archived shells. Black symbols represent data over the
724	annual growth of one specimen; white circle is the average of the shell. 21 <sup>st</sup> century shell data are
725	from the shells sampled at high-resolution in Fig. 7 and the shell from the Bay (40m) in Fig. 4
726	(totaling 4 shells).

















