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# 1 Construction of a Tephra-Based Multi-Archive Coherent Chronological Framework for the Last 2 Deglaciation in the Mediterranean Region

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4

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## 9 Abstract

10 Proxy records from different climate archives such as ice cores, speleothems or sediment cores **are**  
11 **essential to define** the sequence of events over to the last deglaciation. However, multi-archive  
12 comparison and compilation of data, necessary to assess the robustness of climate models, are  
13 rapidly limited by inconsistencies between **archives' chronology**. Here we present the development  
14 and validation of the Datice **chronological integration tool** for the construction of multi-archive  
15 coherent chronologies. This **chronology building tool**, first developed to date ice cores only, can now  
16 integrate deposition-like archives such as sediment cores and speleothems, **independently or**  
17 **coherently**. The robustness of this dating method resides in its capacity to build coherent  
18 chronologies for multiple archives with a proper calculation of chronological uncertainties. Using this  
19 tool, we were able to construct a coherent chronology for the last deglaciation in the Mediterranean  
20 region based on volcanic tephra layers correlation **in terrestrial and marine sediment cores**. We  
21 confirm the synchronicity, **within chronological errors**, of the sequence of events characterizing the  
22 last deglaciation between Greenland and the Mediterranean region, independently of any climatic  
23 alignment assumptions. Using this chronological framework, we however highlight some regional  
24 expression of this transition period in term of vegetation cover over the Mediterranean region.

25 **Key words:** last deglaciation, Mediterranean region, multi-archive comparison, coherent chronology,  
26 tephra, vegetation changes.

## 27 1. Introduction

28 Understanding the mechanisms driving climate changes at the Earth's surface has largely  
29 benefited from measurements of geochemical tracers from various climate archives. Marine cores  
30 have provided a first relatively precise timing of reference for the succession of glacial and

31 interglacial periods over the last millions of years (Imbrie and Imbrie, 1980). Later on, polar ice cores  
32 were the first to clearly evidence the abrupt climate variability characterizing the last glacial period  
33 (Dansgaard-Oeschger events, Dansgaard et al., 1993). Concomitantly, marine cores from North  
34 Atlantic revealed the occurrence of strong iceberg-rafted episodes during the last glacial period,  
35 named Heinrich events, that are related to some Dansgaard-Oeschger events (Heinrich, 1988). Since  
36 this pioneer period in paleoclimatology, many different archives from the continent and the ocean  
37 have recorded the succession of glacial and interglacial periods as well as the millennial-scale climate  
38 variability of the last glacial periods (Martrat et al., 2004; Jouzel et al., 2007; Cheng et al., 2009).  
39 Obtaining records from different latitudes is essential to accurately depict the regional as well as  
40 global variability associated with past climate changes. With the improvement of climate models  
41 (Eyring et al., 2016), it is becoming clear and necessary to combine records from different archives in  
42 order to further assess the robustness of model outputs and their capacity to reproduce the  
43 duration and pace of past climate changes at different spatial scales (Kageyama et al., 2018).

44 Efforts have been made to combine records of marine and continental origin covering the last  
45 deglaciation (Clark et al., 2012; Shakun et al., 2012; Moreno et al., 2014). This period framed  
46 between 19 and 11 ka ago corresponds to the transition between the last glacial period and the  
47 current interglacial (i.e. Holocene). It was induced by changes in insolation, associated with changes  
48 in greenhouse gases (GHG) concentration, atmospheric and ocean circulation reorganisations as well  
49 as the melting of polar ice sheets. The last deglaciation is widely recorded within various  
50 paleoclimatic archives. Even if numerous radiocarbon ages constrain the sequence of events of this  
51 key period, the absolute timing of changes is not always consistent from one site to another  
52 essentially because of the spatial and temporal variability of reservoir ages, the calibration of the  $^{14}\text{C}$   
53 dates (Siani et al., 2001; Reimer et al., 2013), but also the sample selection (e.g. shells, wood, bulk  
54 sediment; Müller et al., 2011). The combination of climate data compilations and model simulations  
55 generally agree with the major role of GHGs and Atlantic Meridional Overturning Circulation (AMOC)  
56 variations during the last deglaciation. However, the relative leads and lags between different  
57 regions and/or type of paleoclimate archives remain difficult to assess and therefore largely  
58 unknown. Indeed, even if efforts were made in order to calibrate  $^{14}\text{C}$  dates on the same reference  
59 curves (Clark et al., 2012; Shakun et al., 2012), the harmonization of chronologies between the  
60 ocean and the continent, when it exists, remains largely based on alignment assumptions. For  
61 example, Moreno et al., (2014) have combined terrestrial records covering the period 60-8 ka from  
62 lake sediments, speleothems, an ice core and pollen records from marine cores. They choose not to  
63 revise the published age models. However, most of the marine cores included in their compilation  
64 possess chronologies derived from the alignment of their surface records (planktonic foraminifera

65  $\delta^{18}\text{O}$  or Sea Surface Temperature - SST) with a Greenland reference record ( $\delta^{18}\text{O}_{\text{ice}}$  of NGRIP usually).  
66 While such an assumption can be easily verified close to Greenland, thanks to the numerous volcanic  
67 ash layers (Davies et al., 2010; Austin and Hibbert, 2012), it can be questionable down to the Iberian  
68 margin and the Mediterranean region, especially when there is no assessment of the  
69 synchronization uncertainty.

70

71 Here we propose an assessment of relative climate changes associated with the last deglaciation  
72 using a coherent chronological framework based on tephra layers that are used as anchors to  
73 synchronize records. These volcanic derived deposits are by nature independent from climatic  
74 alignment assumptions (e.g. Lane et al., 2013; Lowe et al., 2015). In order to verify the consistency of  
75 the timing of changes associated with this transition, we primarily focus our chronological efforts on  
76 the Mediterranean region (Figure 1), which presents two main advantages. First, it offers a large  
77 variety of climatic records obtained from different archives on the continent and from the sediments  
78 of the Mediterranean Sea. Second, this region is characterized by widely dispersed tephra layers  
79 found within both continental and marine sedimentary archives. Their geochemical characteristics  
80 allow us to identify common tephra horizons in different archives (Albert et al., 2017). In some cases  
81 these tephra layers are even precisely and accurately dated using the  $^{40}\text{Ar}/^{39}\text{Ar}$  method (Galli et al.,  
82 2017; Giaccio et al., 2017; Albert et al., 2019), or by  $^{14}\text{C}$  dating of the surrounding material (Lee,  
83 2013; Albert et al., 2015; Bronk Ramsey et al., 2015).

84 The consistency between chronologies of the different Mediterranean paleoclimate archives  
85 included in our study is obtained using the Datice **chronological integration tool** (Lemieux-Dudon et  
86 al., 2010). Datice was initially developed for coherent ice core dating, and used to produce the  
87 reference chronology of ice cores (Antarctic Ice Core Chronology - AICC - 2012, Bazin et al., 2013;  
88 Veres et al., 2013). For the purpose of our study, we improved the Datice chronological integration  
89 tool and developed a multi-archive version allowing us to consistently build chronologies for  
90 sediment cores and speleothems in addition to ice cores.

91 The article is organized as follows. First we present the Datice Multi-archive tool and its new  
92 development to build chronologies for deposition-like archives. Second, we list the different sites  
93 included in this study and explain the rationale behind the construction of a tephra-based  
94 chronological framework for the last deglaciation in the Mediterranean region. The comparison of  
95 our coherent chronology with previously published age scales highlights the strength and usefulness  
96 of the Datice tool for multi-archive compilation applications. Finally, we discuss the coherency of

97 climatic changes as recorded by different archives within the Mediterranean region and compare  
98 them with NGRIP  $\delta^{18}\text{O}_{\text{ice}}$  reference record from Greenland.

## 99 2. Tools

100 We use the Datice chronological integration tool to build a multi-archive coherent chronology  
101 (<https://datice-multi-archives.ipsl.fr/>). Information on where and how to install Datice are given in a  
102 dedicated paragraph at the end of the article. Datice was initially developed to produce coherent  
103 chronologies for ice cores by integrating absolute and relative age constraints between several ice  
104 cores for both the ice and gas phases (Lemieux-Dudon et al., 2010; Buiron et al., 2011; Bazin et al.,  
105 2013; Veres et al., 2013). Following this work, additional developments were necessary to constrain  
106 annually resolved archives such as the Greenland ice cores back to 60 ka (GICC05 chronology,  
107 Andersen et al., 2006; Rasmussen et al., 2006; Vinther et al., 2006; Svensson et al., 2008), and led to  
108 the integration of duration constraints (Lemieux-Dudon et al., 2015). On a mathematical point of  
109 view, Datice is an inverse modelling method based on Bayesian statistics. It makes the best  
110 compromise between an initial age-depth model, and its associated uncertainty estimation, with  
111 absolute and relative age constraints obtained at various depths. The prior chronology is deduced  
112 from the integration with depth of background parameters and corresponding uncertainties, all  
113 defined by the user (Table 1). Taking into account the chronological constraints (i.e. all background  
114 parameters and age markers) of all sites, as well as the stratigraphic links defined across sites, Datice  
115 produces a single chronology that is coherent to all sites and benefits from a proper assessment and  
116 propagation of age uncertainties. Moreover, after obtaining the final chronology, uncertainties are  
117 also propagated to the parameters used to calculate the final age for the different archives (e.g.  
118 accumulation rate, deposition rate, thinning function, Lock-in depth).

119 For ice core dating, it is necessary to date both the ice and gas phases. In Datice, the prior ice  
120 chronology is calculated from given scenarios of snow accumulation and thinning function, while the  
121 gas chronology is deduced from the evolution with depth of the reconstructed lock-in depth (i.e. the  
122 depth at which the air is trapped within the ice lattice, typically 50-100 m under the surface, Table 1,  
123 Lemieux-Dudon et al., 2010).

124 For the purpose of multi-archive coherent dating, we further developed a new version of Datice in  
125 order to extent the use of this tool to date paleoclimate archives such as speleothems and  
126 (terrestrial and marine) sediment cores. For these mono-phase archives, the age calculation is  
127 simpler than for ice cores. Contrary to ice, which is formed from the compaction of snow and  
128 continues to thin during burying, speleothems are not affected by post-depositional compaction

129 effects. For sediment cores, compaction effects are generally not taken into account due to the  
130 difficulty to properly quantify or model them. Consequently, the age-depth relationship for such  
131 climate archives (i.e. speleothems and sediment cores) is directly obtained from the integration with  
132 depth, **from the top to the current depth**, of the deposition rate  $L(z)$  (e.g. the growth rate for  
133 speleothems or the sedimentation rate for marine or lake sediment cores in **centimetres per year**;  
134 equation 1).

$$135 \quad \text{Age}(z) = \int_0^z \frac{1}{L(z')} dz' \quad (1)$$

136 As the age is parameterised as increasing with depth, the depth reference is taken at the surface of  
137 the sediment or the top of the speleothem. Similarly to the background parameters for ice cores, the  
138 error of the prior estimation of the deposition rate can be parameterised as constant or varying with  
139 depth. So far, Datice Multi-archive **can only** integrate continuous archives. When dealing with  
140 hiatuses (regardless of the type of archives), each section has to be integrated separately (before  
141 and after the hiatus). **Users have the possibility to associate error correlation with the prior**  
142 **estimation of the deposition rate. Adding such an error correlation allows Datice Multi-archive to**  
143 **propose a final deposition rate with smoother variations than in the case of no error correlation**  
144 **(Table 1).**

145 In order to obtain precise and coherent chronologies, different sets of markers can be included in  
146 Datice Multi-archive (Table 1). Some are specific to ice cores, such as gas age markers or gas  
147 stratigraphic links. However, the markers applied to the ice phase (e.g. age markers, duration  
148 constraints, stratigraphic links) can also be applied to the deposition-like archives. The uncertainty  
149 associated with all markers must be given in years and should take into account all sources of error  
150 (e.g. the identification of markers, resolution of data, measurement uncertainties, calibration). **Ages**  
151 **deduced from  $^{14}\text{C}$  measurements should be calibrated into calendar years before their integration**  
152 **into Datice Multi-archive.** Datice Multi-archive **can also** add error correlation between different sets  
153 of markers if needed. The coherency of the chronology between the sites is ensured by the  
154 integration of stratigraphic links. These markers must be defined *a priori* between two cores and  
155 have to be associated with an uncertainty in years. The final chronology uncertainty calculated by  
156 Datice Multi-archive is then a compromise between the uncertainties of the background chronology  
157 and all age markers. Large uncertainties associated with the background parameters (e.g.  
158 sedimentation rate, accumulation rate, thinning function, Table 1) necessarily result in a large  
159 uncertainty for the prior chronology. When such large uncertainties occur, the optimisation module  
160 of Datice Multi-archive has sufficient freedom to modify accordingly the background parameters and

161 the chronology to fit the more robust dating constraints (i.e. the age markers and the stratigraphic  
 162 links). Inversely, when we are confident in our background parameters, the final chronology will  
 163 remain close to its initial age-depth model.

164 **Datice Multi-archive** is a complementary tool to the already existing applications used to date  
 165 paleoclimate archives. The other Bayesian tools such as OxCal, BChron or Bacon (Haslett and Parnell,  
 166 2008; Bronk Ramsey, 2009; Blaauw and Christeny, 2011; Parnell et al., 2011), are not suited for the  
 167 specific problematic of ice core dating, but they offer the possibility to calibrate <sup>14</sup>C ages as well as  
 168 build consistent chronologies for sediment cores and/or speleothems (Bronk Ramsey et al., 2014,  
 169 2015). On the other hand, IceChrono (Parrenin et al., 2015) builds coherent chronologies for ice  
 170 cores only, similarly to the previous Datice version. Thanks to our recent developments, **Datice**  
 171 **Multi-archive** is the only chronological tool allowing to build coherent chronologies for ice cores as  
 172 well as other paleoclimatic archives such as speleothems or (terrestrial and marine) sediment cores.  
 173 **Datice Multi-archive** offers an easy way to parameterize the relative constraints between the  
 174 different archives through the implementation of the stratigraphic links with the possibility of adding  
 175 error correlation between markers.

176 Consequently, the strength of **Datice Multi-archives** when building chronologies for paleoclimatic  
 177 archives is its capacity to reconstruct chronologies for ice cores and deposition-like archives, with  
 178 the proper transfer of uncertainties. **Datice Multi-archive** is therefore especially useful when  
 179 addressing questions about the relative timing of changes associated with major climatic events that  
 180 are recorded by proxies from multiple sites and archives, such as in the case of the last deglaciation.

181 Table 1: input parameters required for **Datice Multi-archive**

Background parameters			Age markers
Ice cores	Sediment cores	Speleothems	Absolute ages
Accumulation rate	Sedimentation rate	Deposition rate	Duration constraints
Thinning function			Stratigraphic links
Lock-in depth			(+ gas absolute ages, gas stratigraphic links, delta-depth markers, all specific for ice cores)
+ associated uncertainties for each parameter and each dating constraint (possibility of adding error correlations for paired markers)			

182

183 **3. Material and Methods**

184 To test this new Datice version, we choose to integrate paleoclimate archives from the  
185 Mediterranean region that cover continuously the last deglaciation. The selected archives were  
186 taken on the continent (speleothems and lake/peat sediments), and from the Mediterranean Sea,  
187 and exhibit multiple high-resolution proxies. The selected marine and lake sediment cores present  
188 numerous common and independently dated tephra layers (Figure 2). This unique set of  
189 chronological tie points is of key importance to safely discuss the relative changes between the  
190 sediment records without any assumption of climatic synchronicity, as usually implied in case of  
191 chronological alignment. Consequently, we selected one marine core (MD90-917), three continental  
192 sediment cores (Monticchio, Ohrid, Tenaghi Philippon) and two speleothems (Sofular, La Mine)  
193 spread over the Mediterranean perimeter (Figure 1, Table 2). The two selected speleothems, while  
194 being independently dated and not integrated in our tephra-based Mediterranean coherent  
195 chronological framework, bring complementary information about the expression of the last  
196 deglaciation within the Mediterranean region. They, moreover, serve as test cases to validate the  
197 developments and use of Datice Multi-archive to date deposition-like archives (see section 4.1). We  
198 summarized in Table 2 the independent radiometric constraints and relative tie-points available for  
199 each site.

200 Both speleothems chronologies are based on numerous  $^{230}\text{Th}$  ages (Genty et al., 2003; Fleitmann et  
201 al., 2009; Göktürk et al., 2011). For Monticchio, three tephra layers corresponding to ashes from the  
202 Somma-Vesuvius volcanic field have been historically recognised and are associated with a 1-sigma  
203 uncertainty of 1 year ("\*" in Figure 2). The absolute ages of the Mercato and Y2 tephra layers were  
204 obtained after combining together numerous  $^{14}\text{C}$  calibrated ages from different archives (Zanchetta  
205 et al., 2011; Lee, 2013; Bronk Ramsey et al., 2015). This method statistically reduces the uncertainty  
206 associated with one volcanic event that is well-identified and  $^{14}\text{C}$ -dated in numerous paleoclimate  
207 archives, especially when no  $^{40}\text{Ar}/^{39}\text{Ar}$  date is available. The remaining absolute ages correspond to  
208  $^{40}\text{Ar}/^{39}\text{Ar}$  dates obtained on inland deposits of well-known eruptions ("o" in Figure 2, Pappalardo et  
209 al., 1999; Galli et al., 2017; Giaccio et al., 2017; Albert et al., 2019). As their published ages were  
210 obtained from different calibrations, we decided to homogenize all the  $^{40}\text{Ar}/^{39}\text{Ar}$  ages. Ages are then  
211 recalculated using the more recent ACs-2 standard of Niespolo et al., (2017) (1.1891 Ma), and based  
212 on the  $^{40}\text{K}$  total decay constant of Renne et al., (2011) that is independent from astronomical tuning  
213 (Figure 2). Consequently, the corresponding ages used in this work, and presented in Figure 2, may  
214 slightly differ from the published ages due to this homogenization and recalibration process (see  
215 supplementary material). Including the recent work of Albert et al., 2019, all three major eruptions  
216 of the Campi Flegrei (i.e. NYT, Y3, Y5) that are recognised at the different sites are now precisely and

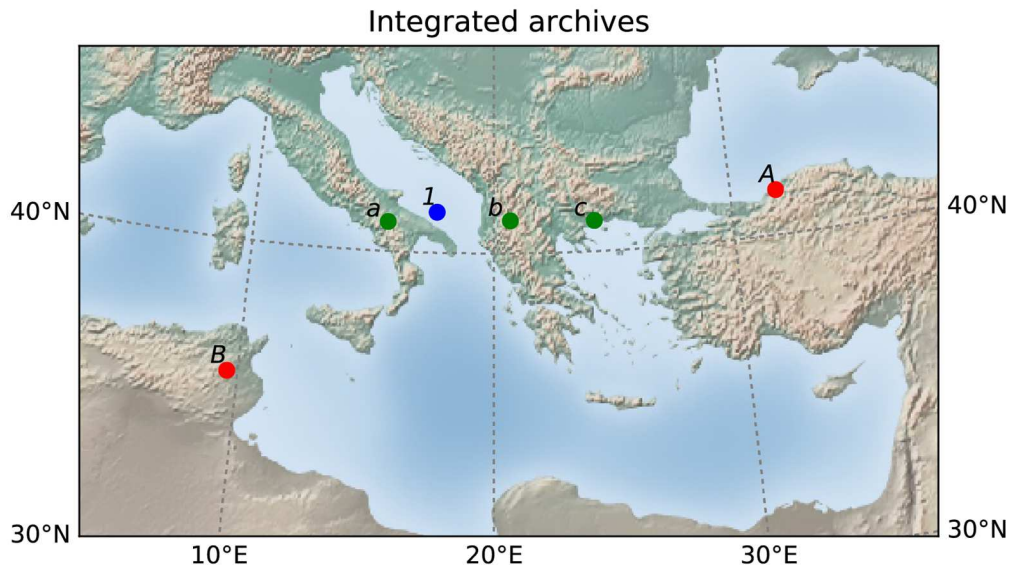


217 accurately dated by  $^{40}\text{Ar}/^{39}\text{Ar}$ . They all remain in good agreement with the  $^{14}\text{C}$  ages previously  
218 published.

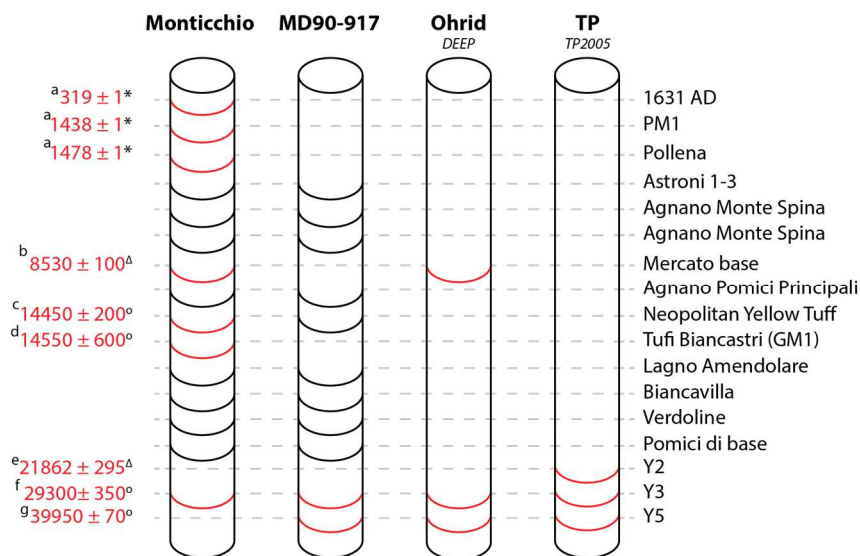
219 Finally, the chronological coherency of all sediment cores is assured by the identification of common  
220 tephra layers, which transfer the absolute age of an independently and better dated record to a  
221 poorly dated one (Figure 2). The identifications and correlation of the tephra layers are based on the  
222 published geochemical composition of tephra and their already proposed correlations (Siani et al.,  
223 2001, 2004, Wulf et al., 2004, 2008, 2018, Tomlinson et al., 2012, 2014; Leicher et al., 2016; Albert et  
224 al., 2019). We further confirmed these assignments using the RESET database (Lowe et al., 2015).  
225 We assigned an uncertainty of 1-year to these volcanic stratigraphic links. Tephra deposition after an  
226 eruption can be considered instantaneous (i.e. <1 year). While we may encounter some bioturbation  
227 perturbation in the top-most centimetres of the sediment locally, the chronological error associated  
228 to this parameter remains difficult to model and quantify (Bard et al., 1987; Manighetti et al., 1995;  
229 Carey, 1997; Charbit, 2002; Barsanti et al., 2011). However, enlarging the uncertainty of our  
230 volcanic-based stratigraphic links up to 10 years to account for these potential perturbations do not  
231 affect significantly the final chronology (not shown). Such a test gives us confidence in the tephra-  
232 based chronology using a 1-year uncertainty associated to the tephra correlation.

233 To summarize, our coherent chronological context of the last deglaciation in the Mediterranean  
234 region is constrained by 53 independently dated age markers and 20 relative stratigraphic links  
235 (Table 2 and Figure 2) for the sediment cores. Additional 59 and 13 absolute age markers constrain  
236 the Sofular and La Mine speleothem chronologies respectively, which remain here independent from  
237 the tephra-based common chronology.

238 To complete the set of *a priori* parameters (Table 1), we need to propose a prior estimation of the  
239 sedimentation/deposition rates and associated uncertainties for each deposition-like archive  
240 integrated into Datice Multi-archive. Ideally these estimates have to be independent from the age  
241 constraints, which is not always achievable. For Monticchio, we deduced the prior sedimentation  
242 rate from its published varved chronology covering the last deglaciation (Allen et al., 1999) and  
243 assigned a 10% uncertainty, as proposed in the original paper (Table 3). For the other Mediterranean  
244 archives, no prior estimation of the sedimentation rate independent from the absolute age markers  
245 were available. Consequently we decided to estimate a constant background sedimentation rate  
246 from the linear regression of all the age markers for each site and associated it with a large constant  
247 uncertainty ( $\geq 50\%$ , Table 3). Such a parameterization allows Datice Multi-archive to substantially  
248 modify the background chronology and, as a result, respect the collection of age constraints. The  
249 chronology produced by Datice then better agrees with the age markers than with the prior guess.



252 Figure 1: Map of all the sites discussed in this study. Speleothems are in red (A - Sofular, B - La Mine),  
 253 lake/peat sediments are in green (a - Monticchio, b - Ohrid, c - Tenaghi Philippon) and the MD90-917  
 254 marine core is in blue (1). All identifications are consistent with Table 2.



256 Figure 2: Common and well-dated tephra layers in the Mediterranean region identified in the  
 257 selected terrestrial and marine sediment cores (Monticchio, MD90-917, Ohrid and Tenaghi Philippon  
 258 – TP, (Siani et al., 2001, 2004, Wulf et al., 2004, 2008, 2018, Tomlinson et al., 2012, 2014; Lowe et al.,  
 259 2015; Leicher et al., 2016; Albert et al., 2019). Tephra in red are absolutely dated with the reference

260 year of 1950 AD (\* Somma-Vesuvius historical eruption,  $^{40}\text{Ar}/^{39}\text{Ar}$  dates or  $\Delta$  average  $^{14}\text{C}$  ages with 1-  
 261 sigma errors). For more consistency, all  $^{40}\text{Ar}/^{39}\text{Ar}$  ages are recalculated using the more recent ACs-2  
 262 standard and the  $^{40}\text{K}$  total decay constant independent from astronomical calibration (Renne et al.,  
 263 2011; Niespolo et al., 2017). Ages can therefore slightly differ from the original publications cited  
 264 here. References : a - Wulf et al., 2008, b - Zanchetta et al., 2011, c - Galli et al., 2017, d - Pappalardo  
 265 et al., 1999, e - Bronk Ramsey et al., 2015, f - Albert et al., 2019, g - Giaccio et al., 2017.

266

267 Table 2: List of Mediterranean sites and age constraints considered here for the last deglaciation.

ID	Name	Lat. (°N)	Long. (°E)	Absolute markers	Stratigraphic links (relative)	References
1	MD90-917	41.00	17.52	25 $^{14}\text{C}$ cal. ages 2 tephras	10 to Monticchio 3 to TP 2 to Ohrid	Siani et al., 2001, 2004, 2010; this study
a	Monticchio	40.93	15.58	3 historical eruptions 4 tephras	10 to MD90-917 2 to Ohrid 1 to TP	Allen et al., 1999; Wulf et al., 2008
b	Ohrid	41.03	20.7	3 tephras	2 to Monticchio 2 to TP 2 to MD90-917	Francke et al., 2016
c	Tenaghi Philippon (TP)	40.97	24.22	13 $^{14}\text{C}$ cal. ages 3 tephras	2 to Ohrid 3 to MD90-917 1 to Monticchio	Müller et al., 2011; Wulf et al., 2018
A	Sofular	41.42	31.93	59 U/Th ages		Fleitmann et al., 2009; Göktürk et al., 2011
B	La Mine	35.83	9.58	13 U/Th ages		Genty et al., 2006

268

269 Table 3: Values of constant sedimentation rates and corresponding errors used as background  
 270 parameters in Datice Multi-archive.

	MD90-917	Monticchio	Ohrid	Tenaghi Philippon	Sofular	La Mine
<b>Sed. rate</b> (cm/a)	0.0346	Varved chronology <sup>a</sup>	0.04	0.0338	0.008	0.001
<b>Sigma</b> (cm/a)	0.015	10%	0.02	0.02	0.01	0.01

271 <sup>a</sup>Allen et al., 1999

## 272 4. Results

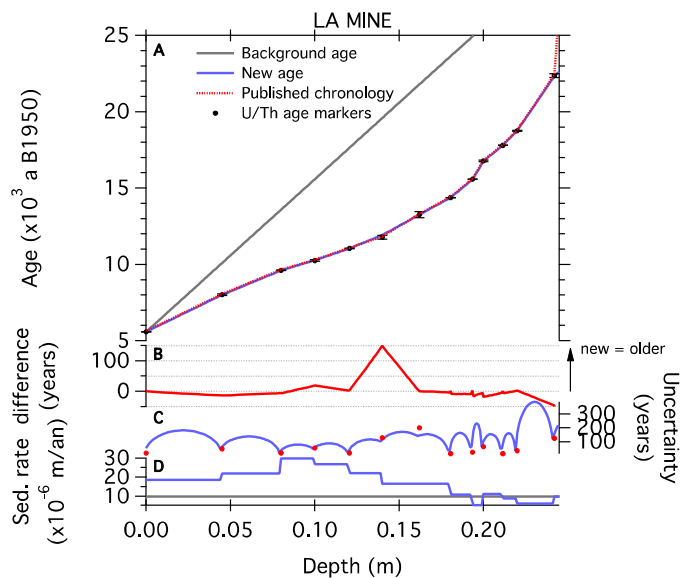
### 273 4.1. Datice chronologies of deposition-like archives

274 We validate Datice Multi-archive developments to **build chronologies for** deposition-like archives,  
275 such as speleothems and sediment cores, using the Tunisian speleothem of La Mine and the Turkish  
276 speleothem So1 of Sofular cave (Figure 1, Table 2). The published chronologies of La Mine and So1  
277 speleothems were obtained through linear interpolation between the 13 and 59 <sup>230</sup>Th dates for La  
278 Mine and So1, with 2 $\sigma$  uncertainties ranging between 0.17 % - 3.03 % and 0.26% - 7.53%, of ages,  
279 respectively (Genty et al., 2003, 2006; Fleitmann et al., 2009; Göktürk et al., 2011). The speleothem  
280 of La Mine grew continuously between 5.5 and 23.0 ka, and the So1 speleothem of Sofular grew  
281 nearly continuously over the last 21.2 ka (Genty et al., 2003, 2006; Fleitmann et al., 2009; Göktürk et  
282 al., 2011). Following our background parameterization, the chronologies calculated by Datice Multi-  
283 archive are then mostly constrained by the age markers with a minor influence of the prior growth  
284 rate estimations.

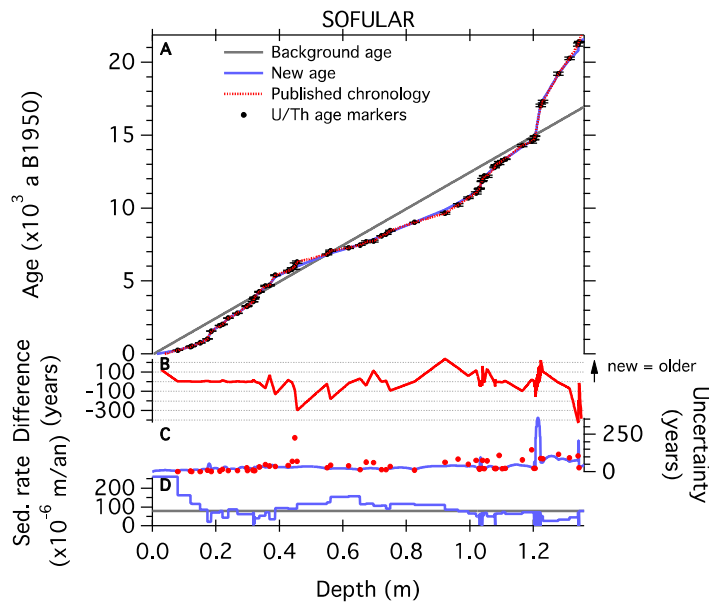
285 In figures 3 and 4, we can note that the chronologies obtained with Datice are overall in very good  
286 agreement with the published age scales of La Mine and Sofular. Compared to a simple linear  
287 regression method, which leads to constant sedimentation rates in-between markers and no  
288 uncertainty propagation (e.g. red markers on panels C of Figures 3 and 4), Datice Multi-archive  
289 produces smoother variations of the sedimentation rate with the calculation of the chronological  
290 uncertainty at all points (blue curves in figures 3 and 4). The resulting Datice Multi-archive  
291 chronology slightly differs from the published age scale of La Mine and Sofular over these intervals  
292 (panels B of Figures 3 and 4). For La Mine (Figure 3), the largest difference between the chronologies  
293 reaches about 150 years, and occurs when Datice Multi-archive does not strictly respect the  
294 absolute age of the nearby marker. The final chronology remains, however, in agreement with the  
295 absolute dates within the uncertainty range (i.e. 170 years at 12.6 ka). The difference between both  
296 chronologies is induced by the smoother evolution of age with depth with Datice Multi-archive than  
297 for the original chronology of La Mine. **This difference occurs for the first marker encountered with a**  
298 **significantly larger uncertainty than all three previous age constraints. In such a case, Datice allows**  
299 **the prior estimation of the deposition rate to have more impact on the final chronology relatively to**  
300 **the previous, more precise, age markers.** For Sofular (Figure 4), the difference between both  
301 chronologies ranges between +150 years/-300 years for Datice Multi-archive chronology. Similarly to  
302 La Mine, the largest differences between the published and the Datice Multi-archive chronologies  
303 are observed when our chronological integration tool does not strictly respect the age markers. The  
304 first large age difference of 300 years is observed at around 0.45 m from the top of the speleothem  
305 at a time when Datice Multi-archive does not respect the marker at ~6 ka, which is associated with a  
306 larger uncertainty than the surrounding ones (226 years at 1 $\sigma$ , against 69 years for the two closest  
307 markers). A strict fit to this age marker at ~6 ka would have induced a strong shift in the deposition

308 rate. Instead, Datice Multi-archive prefers to optimize and bring in agreement the chronology with  
 309 the age markers presenting smaller uncertainties, preventing an abrupt change in deposition rate  
 310 not primarily documented in the background estimate. Finally, the periods where the chronology  
 311 uncertainties produced by Datice Multi-archive are the largest correspond to tipping periods close to  
 312 a significant change in the deposition rate of calcite. In this case, Datice Multi-archive tries to  
 313 reconcile the influence of two very close markers with significantly different absolute ages.

314 Using Datice Multi-archive to date mono-archives provides a proper assessment of the chronology  
 315 uncertainty and sedimentation rate in-between the age markers. This method is more integrative  
 316 than a simple linear interpolation between the age markers, **as it was implemented in the Sofular  
 317 and La Mine original publications.**



318  
 319 Figure 3: Comparison of chronologies for La Mine. A: depth-age relationship and age makers used to  
 320 constrain the chronologies. The published chronology is in red and the new chronology obtained  
 321 with Datice is in blue. B: Difference between the new chronology and the published one of Genty et  
 322 al., 2006 (positive values mean that the new chronology produced by Datice is older than the  
 323 published ages). C: Uncertainties of chronologies (blue for Datice, red markers for the original  
 324 chronology). D: sedimentation rates. The grey curves represent the prior estimation given to Datice,  
 325 the blue curves are the resulting parameters as calculated by Datice after optimization of the  
 326 chronology.



327

328 Figure 4: Same as Figure 3 for the Sofular So1 speleothem (published chronology from Fleitmann et  
 329 al., 2009 and Göktürk et al., 2011).

330

331 **4.2. Datiche Multi-archive coherent chronology for the Mediterranean region**

332 In this section we present the results of the Mediterranean coherent chronological framework  
 333 integrating all the selected sediment/peat archives (Table 2). The chronological coherency between  
 334 the different sedimentary sites is independent of any assumption of climatic alignment and only  
 335 based on the identification of common and independently-dated tephra layers (Figure 2). We choose  
 336 to keep the speleothem chronologies independent from the Mediterranean sediment cores in order  
 337 to remain consistent and avoid unverified assumptions of climatic alignments.

338 The new chronology obtained with Datiche for MD90-917 is in generally good agreement with the  
 339 previously published chronology covering the last 24 ka (Siani et al., 2004, 2010, Figure 5). The  
 340 published chronology of MD90-917 is based on 21 calibrated radiocarbon ages combined with 1  
 341 tephra layer back to 24 ka and using a simple linear regression between the markers (Siani et al.,  
 342 2010, 2013). Before this period, ages were extrapolated based on the last  $^{14}\text{C}$  calibrated age used in  
 343 this study. Compared to the published chronology, we have revised the calibration of the  $^{14}\text{C}$  ages  
 344 into calendar ages using the IntCal13 calibration curve (Reimer et al., 2013). The largest difference  
 345 between the two chronologies, up to 1115 years, is observed at the proximity of the Y3 tephra. A  
 346 second significant difference occur at the proximity of the Biancavilla tephra (352 years difference at  
 347  $\sim 16.85 \pm 0.27$  ka), common with Monticchio, and is induced by the combination of the tephra  
 348 correlation and the revision of the  $^{14}\text{C}$  age marker measured just below the tephra layer. The  
 349 remaining differences between the two chronologies are within the uncertainty range and can be

350 explained by the revision of the calibration of the  $^{14}\text{C}$  ages and possibly the different age calculation  
351 methods (i.e. best compromise for Datice Multi-archive vs. linear interpolation for the published  
352 chronology). The close resemblance of the two chronologies confirms the reservoir ages estimates of  
353 (Siani et al., 2001) for the Adriatic Sea from the late glacial to the Holocene. The final uncertainty of  
354 MD90-917 chronology is strongly modulated by the occurrence of age markers. Indeed, the largest  
355 uncertainty values are observed at periods without age makers (e.g. **~480 years between 21-30 ka**  
356 **and 30-40 ka,  $1\sigma$** ). The new Datice Multi-archive chronology for MD90-917 is updated and more  
357 robust than the published chronology of Siani et al., 2010 over the last 24 ka, especially through  
358 providing a comprehensive estimation of the chronological uncertainty and the transfer of  
359 Monticchio's absolute constraints through the numerous stratigraphic links between the two cores.

360 For Monticchio, the tephra-based chronology appears a bit different from the varve-counted  
361 chronology covering the last deglaciation proposed by Allen et al. (1999), which presents an  
362 uncertainty of 5-10%. The timing of the vegetation changes corresponding roughly to the onset of  
363 the Bølling Allerød can be considered consistent within the uncertainty range of both chronologies  
364 (i.e.  $15.30\pm 0.77$  ka for the chronology of Allen et al., (1999) and  $15.54\pm 0.14$  ka for Datice Multi-  
365 archive chronology as recorded by the temperate taxa, Figure 5). However, further back in time the  
366 varve-counted chronology and the new Datice Multi-archive age scale start to significantly differ at  
367 around 18 ka (Figure 5), at the proximity of the Verdoline tephra layer and down to the bottom of  
368 the core section. Due to the correlated tephra layers, the absolute age of Monticchio is indirectly  
369 constrained by the  $^{14}\text{C}$  calibrated ages of MD90-917 from the Tufi Biancastri ( $GM1 - 14.55\pm 0.60$  ka)  
370 back to the Pomici di base tephra ( $21.27\pm 0.30$  ka on the new chronology), leading to older ages for  
371 this core section. The older Datice Multi-archive chronology is further confirmed and constrained by  
372 the absolutely dated Y3 tephra layer, common for all the sediment cores included here. Similarly to  
373 MD90-917, the uncertainty of the Datice Multi-archive chronology is larger when no marker  
374 constrain the chronology, e.g. up to **~500 years between 20 ka and 30 ka** and up to **~680 years back**  
375 **to 38 ka**, but remain smaller than the 5-10% estimated uncertainty of the varved chronology.

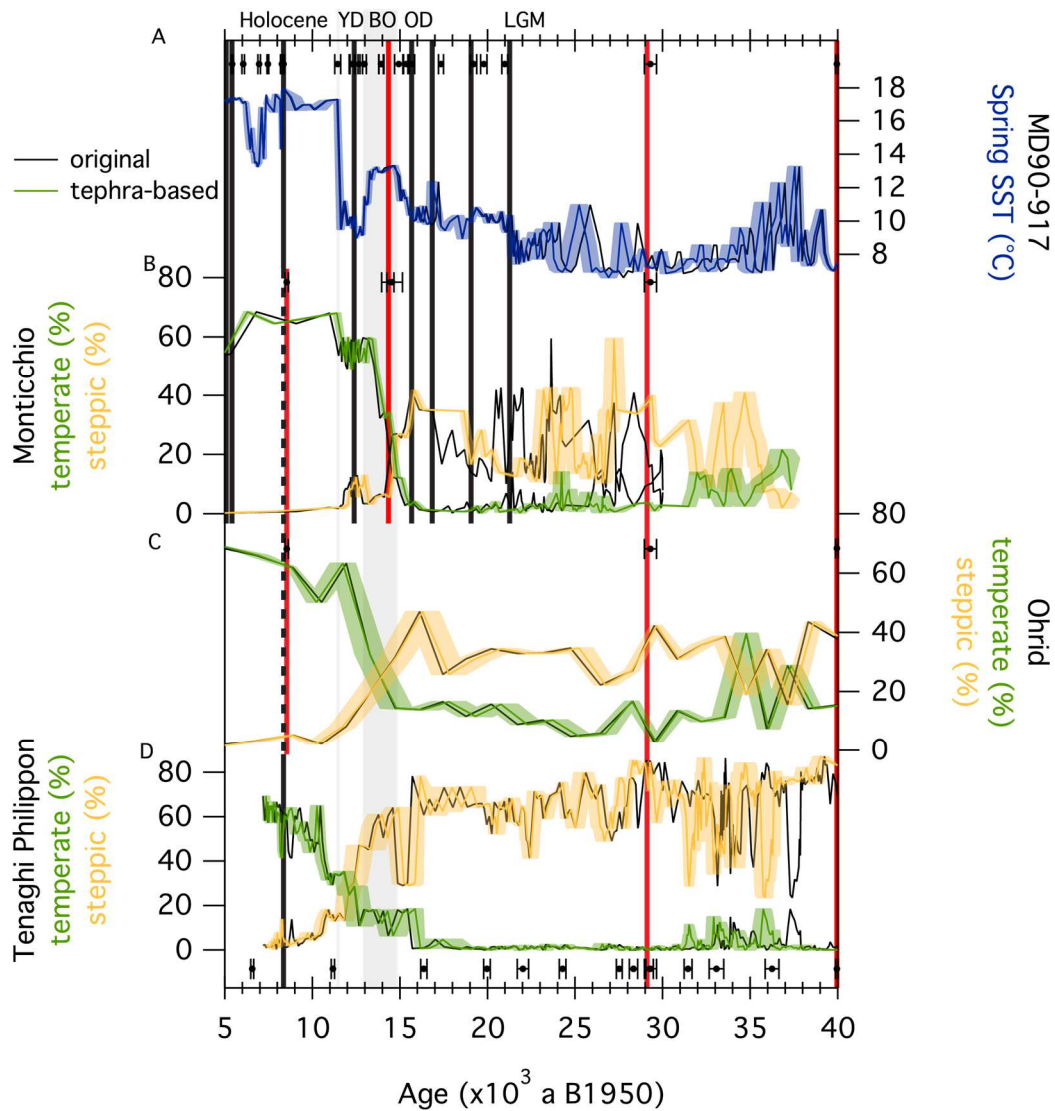
376 At Ohrid, the chronology of Leicher et al., (2016) was obtained after linear interpolation between the  
377 absolute ages of numerous tephra layers back to 637 ka. We used the same tephra layers to  
378 constrain our chronology over our period of interest (i.e. the Mercato, Y3 and Y5 tephtras, Figure 2)  
379 with revised ages using harmonized  $^{40}\text{K}$  decay constant and the recent ACs-2 standard for all  
380  $^{40}\text{Ar}/^{39}\text{Ar}$  dated tephra (see section 3). **Both chronologies are very consistent, always in agreement**  
381 **within uncertainties**. Moreover, similarly to the other sediment cores, the final uncertainty of the  
382 chronology tends to increase away from the age markers (Figure 5).

383 The original chronology of Tenaghi Philippon was obtained after the calibration of 20 <sup>14</sup>C ages from  
384 wood or peat bulk samples following the method of Weninger and Jöris (2008). The Datiche Multi-  
385 archive chronology of Tenaghi Philippon is constrained with the revised <sup>14</sup>C calibrated ages proposed  
386 by Wulf et al., 2018, combined with the revised age of the Y2 tephra layer (black markers on panel D  
387 of Figure 5, Bronk Ramsey et al., 2015). The coherency of its chronology with the other  
388 Mediterranean sites is only assured at the proximity of the E1, Y3 and Y5 tephra layers (Figures 2 and  
389 5). When comparing both chronologies, we notice two periods with major differences: ~465 years at  
390 around 16 ka and up to 1648 years between the Y3 and Y5 tephra layers. The differences for the  
391 older period most probably originate from the influence of the tephra layers and their revised  
392 absolute ages. For the younger period, it occurs between two age markers separated by ~5 ka,  
393 where the background parameter can have an influence on the final chronology. However, when  
394 considering both chronologies uncertainties, changes remain consistent.

395 By comparing our tephra-based coherent chronological framework with the published chronologies  
396 of each sediment site, we present the potential of Datiche Multi-archive as a tool to build coherent  
397 chronologies that are common to various paleoclimate archives. The common tephra-based  
398 chronology allows us now to discuss the relative timing of changes during the last deglaciation as  
399 recorded by different proxies from multiple climate archives in the Mediterranean region.

400





401

402 Figure 5: Comparison of chronologies between the published and the tephra-based coherent  
 403 chronological framework (black lines = published, colored lines = tephra-based chronology with 1-  
 404 sigma uncertainty envelopes). The positions of volcanic-based stratigraphic links are indicated by the  
 405 vertical lines between the different sites. A: SST reconstruction based on planktonic foraminifera  
 406 assemblages for core MD90-917 (Siani et al., 2010, 2013, this study). The top black markers indicate  
 407 the position of age markers constraining its chronology. B: pollen records of Monticchio (green =  
 408 temperate, orange = steppic; Allen et al., 1999). The absolute constraints for Monticchio (black  
 409 markers) are present on top of the record. C: pollen records of Ohrid (green = temperate, orange =  
 410 steppic ; Sadori et al., 2016). Absolute constraints are represented by the black markers on top of  
 411 the records. D: pollen records of Tenaghi Philippon (green = temperate, orange = steppic; Müller et  
 412 al., 2011). The position of <sup>14</sup>C calibrated age and absolutely dated markers of Tenaghi Philippon are  
 413 indicated at the bottom of the records. Selected periods are indicated on top of the figure: YD for  
 414 Younger Dryas (~13-11.5 ka), BA for Bølling Allerød (~15-13 ka), OD for Older Dryas (~16-15 ka) and  
 415 LGM for Last Glacial Maximum.

416

417 **5. Discussion: The last deglaciation in the Mediterranean region**

418 Thanks to the different chronological constraints integrated for our different archives, we now have  
419 a coherent chronology for the sediment cores of MD90-917, Monticchio, Tenaghi Philippon and  
420 Ohrid, independent from any climatic alignment assumption. The climatic representation in the  
421 Mediterranean region is complemented here with the speleothem records of La Mine and Sofular,  
422 used to validate the development of Datice Multi-archive. It is now interesting to look at how our  
423 new coherent and improved Mediterranean chronology for the last deglaciation fits within the more  
424 global context of this transition.

425 The  $\delta^{18}\text{O}_{\text{ice}}$  of NGRIP in Greenland records regional air temperature variability over the last glacial-  
426 interglacial cycle with an annual/seasonal resolution (NGRIP community Members, 2004). This  
427 record is dated back to 60 ka using annual layer counting (Andersen et al., 2006; Rasmussen et al.,  
428 2006; Vinther et al., 2006; Svensson et al., 2008) and is often used as a reference for comparison  
429 with new data. Compared to previous studies, we can now compare the absolute timing of changes  
430 recorded in Greenland and within the Mediterranean region independently of any alignment  
431 assumptions, thanks to our precise and coherent volcanic-based chronological framework. In figure  
432 6, we compare the absolute timing of changes as recorded by Greenland  $\delta^{18}\text{O}_{\text{ice}}$ , to the  
433 Mediterranean SST reconstruction of core MD90-917 presented on the tephra-based coherent  
434 chronology, as well as La Mine and Sofular speleothems records (panels A, B, G and H of Figure 6  
435 respectively). Consistent ages are observed for the sharp transitions recorded by the  $\delta^{18}\text{O}_{\text{ice}}$  record of  
436 NGRIP and the MD90-917 SST reconstruction characterizing the Older Dryas (OD) - Bølling Allerød  
437 (BA), BA – Younger Dryas (YD) and the YD – Holocene transitions (Figure 6). The timing of changes for  
438 MD90-917 is not significantly changed when compared to its published chronology over these  
439 periods. However, the construction of the tephra-based Datice Multi-archive chronology for MD90-  
440 917 confirms the reservoir age values proposed by Siani et al., 2001 over the LGM - Holocene  
441 transition (i.e.  $\sim 820 \pm 120$  years over the OD,  $\sim 390 \pm 80$  years for the LGM, BA, YD and Holocene). The  
442 consistent timing of changes associated with these sharp transitions has already largely been  
443 recorded by European speleothems (Genty et al., 2006) and western Mediterranean Sea records,  
444 however relying on  $^{14}\text{C}$  calibrations with constant reservoir ages and/or tuning to NGRIP records  
445 (Cacho et al., 2001; Martrat et al., 2004; Jiménez-Amat and Zahn, 2015). Consequently, using an  
446 independent volcanic-based age model for the Mediterranean sediment cores, we validate the  
447 general assumption of synchronous temperature changes associated with the last deglaciation over  
448 Greenland and within the Mediterranean region, independently of any climatic alignment.

449 Thanks to our volcanic-based coherent chronological framework, we can discuss the relative changes  
450 of vegetation between the marine core MD90-917 (Combourieu-Nebout et al., 1998, 2013),

451 Monticchio (Allen et al., 1999), Ohrid (Sadori et al., 2016) and Tenaghi Philippon (Müller et al., 2011)  
452 (panels C, D, E and F of Figure 6 respectively). **With our new coherent chronological framework, we**  
453 **now have a proper assessment of the chronological uncertainties for each site, rendering possible**  
454 **the inter-comparison of relative vegetation changes between Monticchio, MD90-917, Tenaghi**  
455 **Philippon and Ohrid.** First, all records agree in indicating a predominance of temperate taxa during  
456 the Holocene, and more steppic taxa during the last glacial period. However, the timing of changes  
457 in the predominance of temperate versus steppic pollen taxa at the different sites differs  
458 significantly and is now precisely established thanks to our multi-archive coherent chronology. At  
459 Monticchio, temperate tree taxa start to increase simultaneously with the SST increase recorded in  
460 the Adriatic Sea by the MD90-917 core, and become predominant at the onset of the Bølling-Allerød  
461 period (i.e. at  $14.63 \pm 0.14$  ka on our coherent chronological framework). Ohrid records indicate a  
462 similar behaviour slightly later (i.e.  $14.04 \pm 0.87$  ka), that can be considered synchronous with  
463 Monticchio within uncertainties. On the other hand, records from MD90-917 and Tenaghi Philippon  
464 show a shift in the predominance of temperate vs. steppic taxa at  $\sim 12$  ka (i.e. at  $11.98 \pm 0.16$  ka for  
465 MD90-917 and  $12.10 \pm 0.64$  ka for Tenaghi Philippon). In MD90-917 records, the BA period is  
466 characterized by a dominance of steppic taxa and also higher, but still low, amounts of temperate  
467 taxa. In contrast, Monticchio and Ohrid pollen records indicate a continuous increase in temperate  
468 pollens from the initial gradual warming recorded in the Adriatic Sea until the BA-YD sharp  
469 transition. Such regional differences could originate from the difference in altitude between the  
470 sites. While MD90-917 and Tenaghi Philippon are low-altitude sites, where pollen are mostly  
471 transported by wind and basin-wide hydrological transport, Monticchio and Ohrid sites are higher in  
472 altitudes (i.e. 656 m a.s.l. for Monticchio and 693 m a.s.l. for Ohrid). They may be potentially biased  
473 by high-altitude vegetation changes and less recording vegetation changes from lower altitudes. One  
474 should note that while the temperate and steppic taxa curves present different timing of relative  
475 predominance between the four sites, the records of individual pollen species of MD90-917 (e.g.  
476 *Quercus*, *Alnus*, *Betula*) indicate an increase of trees in altitude on the mountainous slopes around  
477 the Adriatic Sea, while herb species remain dominant (e.g. *Artemisia* and *Ephedra*, Combourieu-  
478 Nebout et al., 1998). Such individual pollen behaviour is consistent with the progressive extension of  
479 forests under wetter and possibly warmer conditions during the BA than the OD, and the continuous  
480 persistence of herbs at low altitudes. Finally, the YD – Holocene transition also indicate local  
481 differences between the pollen records (Figure 6). At Monticchio, the YD-Holocene transition  
482 corresponds to a final sharp increase in the temperate pollen record, which is now concomitant with  
483 the SST and Greenland  $\delta^{18}\text{O}_{\text{ice}}$  sharp changes in our tephra-based chronology. For MD90-917 and  
484 Tenaghi Philippon, the timing of this transition remains unchanged compared to the published

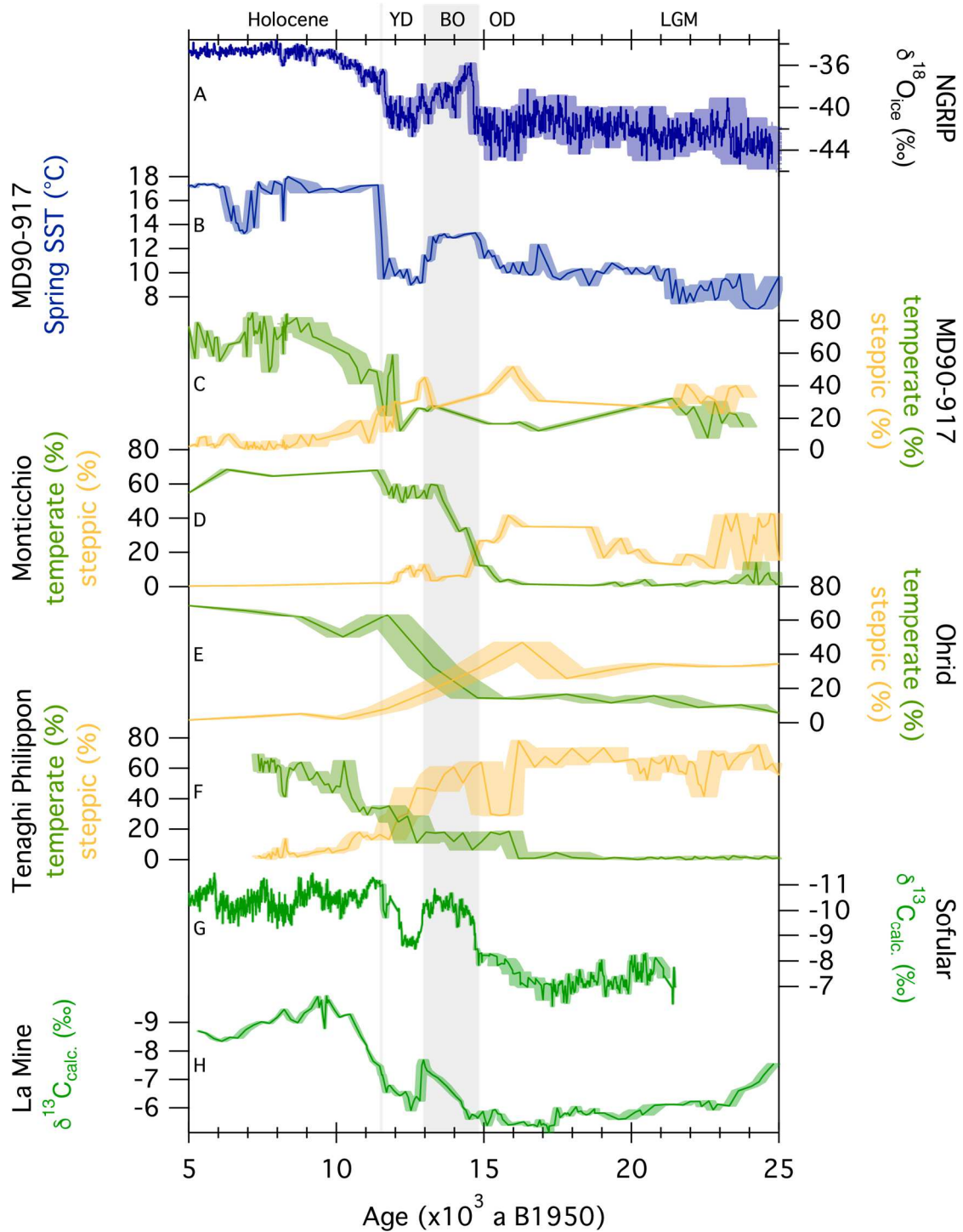
485 chronologies. Nonetheless, their pollen records present significantly different shapes in the  
486 temperate taxa records compared to Monticchio. Indeed, the YD-Holocene transition seems to be  
487 characterized by a primary increase, up to Holocene percentages in MD90-917 and halfway to  
488 Holocene amounts at Tenaghi Philippon, preceding the final shift that is synchronous with the other  
489 sites' temperature changes (Figure 6). The pollen records of Ohrid are too poorly resolved to  
490 differentiate in details the different periods characterizing the last glacial to Holocene transition. We  
491 moreover refrain from more discussion about the absolute timing of changes recorded by Tenaghi  
492 Philippon records because of the lack of stratigraphic links with the other sedimentary cores over  
493 the transition.

494 Depending of the local major influencing factors controlling the isotopic composition of calcite of  
495 speleothems, we can further discuss the relative changes of vegetation types based on the  $\delta^{13}\text{C}_{\text{calc.}}$ .  
496 Due to the geological setting of La Mine speleothem, its  $\delta^{13}\text{C}_{\text{calc.}}$  record is not significantly affected by  
497 vegetation changes over the cave because today's conditions consists in a thin soil with sparse trees  
498 and bushes (Genty et al., 2006). Instead, the  $\delta^{13}\text{C}_{\text{calc.}}$  at this site is more affected by changes in the  
499 soil  $\text{CO}_2$ , which mostly results from the soil biogenic production, then indirectly from temperature  
500 and humidity conditions, and is controlled by the atmospheric  $\text{CO}_2$ . However, it is not possible to  
501 differentiate the  $\delta^{13}\text{C}_{\text{calc.}}$  of La Mine in terms of quantified changes in temperature, precipitation or  
502  $\text{CO}_2$ . The  $\delta^{13}\text{C}_{\text{calc.}}$  of Sofular, on the other hand, is mostly influenced by the type and density of  
503 vegetation as well as the soil microbial activity over the cave (Fleitmann et al., 2009). The low values  
504 of  $\delta^{13}\text{C}_{\text{calc.}}$  ( $\sim -12\text{‰}$ ) recorded at Sofular are interpreted as a predominance of C3 plants ( $\sim$ trees) over  
505 the cave, while values around  $\sim -6\text{‰}$  are associated with a predominance of C4 plants ( $\sim$ grasses).

506 **In order to have a broader idea of vegetation changes in the Mediterranean region during the last**  
507 **deglaciation, we now further compare the pollen records of the integrated sediment cores against**  
508 **the  $\delta^{13}\text{C}_{\text{calc.}}$  of Sofular**, which is interpreted in terms of relative dominance of C3/trees and  
509 C4/grasses plants over the cave (located at 440 m a.s.l.). The Northwestern Turkey record of Sofular  
510 exhibits a predominance of trees during the Bølling Allerød, in agreement with the general trends of  
511 warmer and more humid conditions associated with this period. The  $\delta^{13}\text{C}_{\text{calc}}$  changes are moreover  
512 simultaneous with the sharp SST transitions recorded in the Adriatic Sea. Such a behaviour remains  
513 consistent with the interpretation of the differences between the other Mediterranean pollen  
514 records in terms of altitude sensitivity. Moreover, similarly to MD90-917 and Tenaghi Philippon, the  
515 YD – Holocene transition is associated with a two-step increase in  $\delta^{13}\text{C}_{\text{calc.}}$ , and the second part of the  
516 transition is synchronous with Greenland and Mediterranean SST changes. This preceding event has

517 already been recorded by other pollen and SST records within the western Mediterranean region  
518 (Combourieu Nebout et al., 2009; Jiménez-Amat and Zahn, 2015).

519 While the principal transitions characterizing the last deglaciation in the Mediterranean region seem  
520 to be well recorded and synchronous with Greenland changes, some differences appear in terms of  
521 integrated vegetation changes between the different sites. Even if changes recorded by individual  
522 vegetation species may be synchronous with temperature changes, the predominance of taxa curves  
523 from one site to another may differ. These differences most probably correspond to altitudinal  
524 sensitivities of vegetation species and potentially different pollen transport modes. This is especially  
525 the case for sites located at high altitudes, that can remain close to shelter areas for specific species  
526 over certain periods (Fleitmann et al., 2009). Inversely, pollen records from low altitude sites are  
527 potentially more basin-wide integrated, and may be over-imprinted, compared to the higher-altitude  
528 sites, by low-altitude pollen species. Moreover, as our tephra-based coherent chronological  
529 framework evidences asynchronous changes between temperature and vegetation changes, in  
530 terms of taxa dominance, within the Mediterranean region, we recommend caution when using  
531 synthetic pollen records for climatic alignments with reference curves, such as SST or Greenland  
532 records, to derive age models for sediment cores.



533

534 Figure 6: Comparison of the last deglaciation records, presented on the Mediterranean coherent  
 535 chronology between 5 ka and 25 ka B1950, with NGRIP reference record  $\delta^{18}\text{O}_{\text{ice}}$  (GICC05 chronology;  
 536 NGRIP community Members, 2004). All records are presented with their respective  $1\sigma$  chronological  
 537 uncertainty envelope. Shaded areas highlight the position of prominent events as recorded by the  
 538 SST of MD90-917 (same as in Figure 5). Selected periods are indicated on top of the figure (same as  
 539 Figure 5).

540

541 **Conclusion**

542 In this paper we have presented the development and validation of Datice Multi-archive in order to  
543 **build common chronologies for** deposition-like archives such as speleothems or sediment records.  
544 The advantages of using Datice Multi-archive for paleoclimate studies are threefold: 1- it **can**  
545 **simultaneously build chronologies for** several sites, regardless of the type of archive (ice cores,  
546 speleothems, sediment cores), 2- it **can** build one single coherent and precise chronology common  
547 to all sites, and 3- it gives a proper estimation and propagation of all the chronological uncertainties.  
548 For the first time, we have validated the use of Datice Multi-archive by focusing on the last  
549 deglaciation. We combined together records from one marine core, three lake sediment cores and  
550 two independently-dated speleothems from the Mediterranean region. We have built a coherent  
551 and precise chronology based on the identification of common tephra layers between the sediment  
552 cores. Using this chronology, independent from climatic alignment assumptions, we showed that the  
553 major climatic transitions characterizing the last deglaciation were synchronous, **within**  
554 **uncertainties**, in Greenland and within the Mediterranean region. The combination of records from  
555 sediment cores and speleothems highlights local differences in terms of the timing of predominance  
556 of vegetation taxa during the last deglaciation within the Mediterranean region. However, we point  
557 to caution when combining together records from different archives and recommend a proper  
558 assessment of the parameters affecting these records prior to multi-archive data compilation and  
559 alignment of records.

560 Through this first multi-archive coherent chronological context, we evidence the usefulness of Datice  
561 Multi-archive for assessing the relative timing of changes between records from different climatic  
562 archives. Datice Multi-archive can be used to build coherent chronologies from regional to global  
563 scales while focusing on specific periods for different paleoclimatic applications.

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## 569 **Availability:**

570 **The Datice Multi-archive software is available under a CeCILL license and can be downloaded from**  
571 **the INRIA Forge after registration (see <https://gforge.inria.fr/account/register.php>).** Documentation  
572 **and support for the install of Datice Multi-archive can be found on the INRIA Forge, or upon request**  
573 **to [lucie.bazin@lsce.ipsl.fr](mailto:lucie.bazin@lsce.ipsl.fr) or [lemieux.benedicte@gmail.com](mailto:lemieux.benedicte@gmail.com). The git version control tool is necessary**

574 to clone the INRIA Forge repository. Compiling the executable requires a Fortran compiler with  
575 libraries BLAS-Lapack and netcdf (incidentally Openmp library is necessary to test the multi-threaded  
576 and shared memory beta version). Datice Multi-archive sources include the INRIA modulopt  
577 optimization solvers library (Gilbert and Jonsson, 2009). Information on how to run Datice Multi-  
578 Archive, the different input, param and output files can be found at [https://datice-multi-  
580 archives.ipsl.fr/](https://datice-multi-<br/>579 archives.ipsl.fr/). A visualization tool is also available when installing Datice, requiring Python (version  
2.7) with modules numpy, scipy, matplotlib, wxPython and netcdf4.

#### 581 **Supplementary Material:**

- 582 - 1 file with all background parameters and age markers used to constrain the chronologies  
583 for all sites
- 584 - 1 file with the new Datice chronologies (depth, age, uncertainty) and data for all sites

585

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