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RESEARCH LETTER

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Key Points:

- We monitor the seasonal growth and retreat of both polar caps over MY 36
- We monitor the presence of CO₂ ice on the surface of Mars over MY 36, through the season and through all times of day
- CO₂ ice appears at the surface at equatorial latitudes during the second half of the night around the equinoxes

Supporting Information:

Supporting Information may be found in the online version of this article.

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Diurnal and Seasonal Mapping of Martian Ices With EMIRS

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Abstract Condensation and sublimation of ices at the surface of the planet is a key part of both the Martian H₂O and CO₂ cycles, either from a seasonal or diurnal aspect. While most of the ice is located within the polar caps, surface frost is known to be formed during nighttime down to equatorial latitudes. Here, we use data from the Emirates Mars Infrared Spectrometer onboard the Emirates Mars Mission to monitor the diurnal and seasonal evolution of the ices at the surface of Mars over almost one Martian year. The unique local time coverage provided by the instrument allows us to observe the apparition of equatorial CO₂ frost in the second half of the Martian night around the equinoxes, to its sublimation at sunrise.

Plain Language Summary The H₂O and CO₂ ices that form at the surface on Mars play an important role in the exchange between the atmosphere and the surface of the planet. While most of the ice is located within the two polar caps that grew and shrink seasonally, ice is also known to condensate as surface frost during the night and sublimate during the day. This nighttime surface frost deposition can be observed even at equatorial latitudes. In this paper we use data from the Emirates Mars Infrared Spectrometer onboard the Emirates Mars Mission to detect the H₂O and CO₂ ices at the surface of the planet at all local times over almost one Martian Year, which allows us to monitor both the seasonal and diurnal evolution of the distribution of ices at the surface of Mars. We observe that nighttime CO₂ frost forms at equatorial latitudes in the second half of the night to disappear at sunrise around the Martian equinoxes.

1. Introduction

The Martian polar caps are the main reservoirs of both H₂O and CO₂ ices on the surface of the red planet, at the interface between the surface and the atmosphere with an active exchange of volatiles by sublimation/condensation. Their seasonal growth and retreat (e.g., Appéré et al., 2011; Calvin et al., 2015, 2017; Kieffer & Titus, 2001; Kieffer et al., 2000; Langevin et al., 2005, 2007; Oliva et al., 2022) are an important process in both the H₂O and CO₂ cycles on the present-day Mars, which are two major features of the global Martian atmospheric circulation (Forget et al., 1999; Montmessin et al., 2017; Titus et al., 2017).

Every year, a significant fraction of the atmospheric CO₂ condensates into the seasonal polar caps which results in annual variations of about one-third of the global atmospheric mass (Forget, 1998; Hourdin et al., 1993, 1995; James et al., 1992; Leighton & Murray, 1966). In addition, the presence of ice at the surface of the planet changes the albedo and the surface thermal inertia, thus affecting the energy budget at a local but also planetary scale.

While the ices are essentially located in the polar regions, both CO₂ & H₂O ices can also be observed seasonally under low latitudes, in the shadows of pole-facing slopes or at the bottom of some craters (Brown et al., 2008; Carrozzo et al., 2009; Conway et al., 2012; Lange et al., 2022; Schorghofer & Edgett, 2006; Vincendon et al., 2010). From a more transient perspective, frost has also been observed to be deposited on the surface during nighttime, and remaining until early morning (K. L. Jones et al., 1979; Landis & Mer Athena Science Team, 2007; Piqueux et al., 2016). The presence of this daily cycle of CO₂ frost has been shown to have an impact on surface processes such as gullies or slope streaks formation (Diniaga et al., 2021; Khuller et al., 2021; Lange et al., 2022; Pilorget & Forget, 2016). Thus, better constraining this diurnal frost cycle and the properties of the CO₂ ice in these regions is of importance to better understand current active processes at the surface of Mars. In addition, as the observations of the seasonal growth and retreat of the polar caps over the past years have shown interannual variations

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(Piqueux et al., 2015), it is of interest to monitor the evolution of both the Seasonal North Polar Cap (SNPC) and Seasonal South Polar Cap (SSPC).

In this paper, we use for the first time data from the Emirates Mars Mission to map and monitor the evolution of the Martian polar caps along with the nighttime surface CO₂ frost at low latitudes. First, we describe in Section 2 the data set and methods used in this study for the detection and mapping of the surface ices. Then, Section 3 presents the seasonal and diurnal monitoring of the polar caps and the midlatitude nighttime CO₂ frost. Finally, Section 4 summarizes the main points of the study.

2. Data Set and Methods

2.1. Data Set

The Emirates Mars InfraRed Spectrometer (EMIRS) instrument onboard the Emirates Mars Mission (EMM) “Hope” probe is a Fourier Transform Infrared spectrometer that is observing the Martian surface and atmosphere between 6 and 100 μm with a selectable spectral resolution of 5 cm^{-1} or 10 cm^{-1} from February 2021 (Amiri et al., 2022; Edwards et al., 2021). The unique orbit of EMM allows EMIRS to observe the whole Martian surface across all local times in ~ 4 orbits, which corresponds to $\sim 5^\circ$ of L_s , or 10 Earth days. For every daytime EMIRS observation, high-resolution UV-visible images of the full Martian disk are acquired immediately before by the Emirates eXploration Imager (EXI) instrument (A. R. Jones et al., 2021). The coordination of these two instruments allows a direct comparison between the two datasets.

With a pixel size typically between 100 and 300 km, it is important to consider the spatial extent of each pixel to compute accurate maps, especially for the study of the polar regions where the emission angles are high. However, although polygons of the pixels footprints as computed by the SPICE kernels (C. H. Acton, 1996; C. Acton et al., 2018) are provided to users, their use to generate maps is not straightforward and significantly time-consuming. Thus, we have developed a new Python module called “SPiP” (*Spacecraft Pixel footprint Projection*) that generates an approximation of the pixel footprints projected on a regular longitude/latitude map using 3D trigonometry and assuming (for now) a spherical planet (Stcherbinine, 2023).

In this study, we use the Martian surface temperature previously retrieved using a multiple-step algorithm applied on a large portion of the EMIRS spectra between 7.6 and 40 μm excluding the strong CO₂ absorption band at 15 μm (Smith et al., 2022) between $L_s = 6^\circ$ (MY 36) and $L_s = 11^\circ$ (MY 37), that is, EMM orbits 8 to 323. The uncertainties on the retrieved surface temperature values have been estimated to be about 1 K (Smith et al., 2022).

2.1.1. Data filtering

Then, in order to prevent our processing for instrumental bias and data artifacts, we only consider in our maps the data from pixels that meet the conditions below:

- The emission angle is lower than 80° .
- The retrieved surface temperature is between 140 and 300 K.
- The entire field of view of the pixel is within the Martian disk.

2.2. Diurnally Stable Ice Maps Processing

As ice has higher thermal inertia than the regular Martian soil (typically $>2,000 \text{ J K}^{-1} \text{ m}^{-2} \text{ s}^{-1/2}$ for the water ice (Schorghofer, 2010) and $>1,000 \text{ J K}^{-1} \text{ m}^{-2} \text{ s}^{-1/2}$ for the CO₂ ice (Ciazela et al., 2019) versus $\sim 200 \text{ J K}^{-1} \text{ m}^{-2} \text{ s}^{-1/2}$ (Putzig & Mellon, 2007)), we can detect the presence of surface ice that is stable across the day from the low amplitude of the surface temperature diurnal variations of these regions. Thus, we produce surface temperature maps from EMIRS retrievals at all Martian local times and compute the daily variations to convert them into maps of the presence of diurnally stable surface ice. We converged to this temperature variations method instead of using the absolute surface temperatures in order to include H₂O ice in our retrievals. Although the presence of CO₂ ice can confidently be derived from the absolute surface temperature (see Section 2.3), identifying H₂O ice is more challenging as we have to take into account the presence of available water, and some places can exhibit temperature compatible with water ice while being actually ice-free. One can note that this method does not allow us to retrieve any information on the composition of the ice. Thus, for these maps, the term “ice” can refer to either H₂O or CO₂ ice.

A noteworthy point: in the following, the data are duplicated on both sides of the maps (i.e., in terms of longitude) when running the interpolation and smoothing to prevent edge effects and any dependence on the choice of the center longitude.

First, we gather all the surface temperature values retrieved from EMIRS that meet the data filtering conditions described in Section 2.1 over 4 consecutive EMM orbits to have a full spatial coverage for all local times. Then, we compute surface temperature maps for bins of 1 hr of local time with a spatial resolution of $0.5^\circ \times 0.5^\circ$ by (a) projection of the pixel footprints, (b) linear interpolation of the data over the entire longitude/latitude grid with `scipy.interpolate.griddata` (Virtanen et al., 2020), and (c) Gaussian smoothing with a standard deviation $\sigma = 5$. We also compute for each map the Gaussian kernel-density estimation (KDE) of our data for each map via `scipy.stats.gaussian_kde` (Virtanen et al., 2020), and we flag as “low data density” all the pixels of the final maps associated with a KDE lower than 5×10^{-6} , this value has been tuned experimentally by comparing with the pixel footprints maps.

Next, in order to increase the data coverage and prevent possible spurious temperature retrievals from some pixels, we bin the data into 3-hr maps. To do so, we crop each 1-hr map to keep only the pixels with a high data density and compute the median between the 3 consecutive maps. Then, we apply again a linear interpolation and a Gaussian smoothing filter (with $\sigma = 10$) to reconstruct the data over the whole longitude/latitude grid, and we flag as “low data density” the pixels that have been reconstructed by the interpolation (i.e., without any data from one of the 3 initial cropped 1-hr maps). This gives us a set of 8 maps of the surface temperature for local times ranging from 00:00 to 24:00 with the following binning: 00/03/06/09/12/15/18/21/24.

Then, we compute a map of the amplitude of diurnal temperature variations using only the pixels for which we have at least 7 data points (over 8) across the day. We exclude the maximum and minimum values for each pixel to prevent our results from data artifacts that may affect one map and compute the difference between the maximum and minimum temperature values from the remaining data ΔT . This provides us a map of the amplitude of the temperature variation between daytime and nighttime.

This ΔT map is then converted into a map of the ice diurnally stable at the surface from a calibration made from a comparison with EXI images. To do so, we compared for 9 different values of L_s spanning from $L_s = 58^\circ$ to $L_s = 290^\circ$ the latitudinal extent of the polar caps derived from several EXI images with the corresponding ΔT maps computed with EMIRS, and identified the ΔT values in the EMIRS observations that are associated with the presence of surface ice on the EXI images acquired over the same time. This ice map distinguishes 3 categories:

- “Ice” for pixels associated with $\Delta T \leq 25$ K
- “Maybe ice” for pixels associated with $25 \text{ K} < \Delta T \leq 35$ K
- “Not ice” for pixels associated with $\Delta T > 35$ K

These ΔT thresholds may seem relatively high compared to what will be expected for surfaces continuously covered by ice, but due to the large size of the EMIRS pixel footprints, especially at high latitudes, some of the pixels may include a mix of icy and non-icy regions. This will tend to increase the retrieved temperature when some portion of the pixel footprints.

Finally, we flag as “low data density” the pixels of the ice map associated with less than 7 “high data density” pixels from the 3-hr binned surface temperature maps; and we flag the entire map as “low quality” if more than 50% of its data points are flagged as “low data density.”

These maps are also included in the L3 products released by the EXI instrument as ice masks associated with the clouds' optical depth, as the retrievals cannot distinguish between atmospheric and surface ice (Wolff et al., 2022).

2.3. Local Time CO₂ Ice Maps

Similarly to what is done for the diurnally stable ice maps (see Section 2.2), we select all the EMIRS surface temperature values previously retrieved (Smith et al., 2022) over 4 consecutive orbits, and compare them to the freezing temperature of CO₂ (T_{CO_2}) computed for each pixel according to Clapeyron's law (Piqueux et al., 2016):

$$\ln P = \alpha - \frac{\beta}{T_{\text{CO}_2}} \quad (1)$$

with $\alpha = 23.3494$, $\beta = 3182.48$, and P the CO_2 partial pressure taken as 96% of the total surface pressure (in mbar in Equation 1) (Piqueux et al., 2016). As P cannot be reliably retrieved from EMIRS data alone, it is taken from the Mars Climate Database (MCD) (Forget et al., 1999; Millour et al., 2018) and computed to match the observational parameters of each EMIRS observation (Smith et al., 2022). In the following, as CO_2 is the principal component of the Martian atmosphere, the formation of CO_2 frost on the surface is not limited by the presence of gaseous CO_2 (unlike H_2O frost), so we will consider that CO_2 frost is present at the surface anywhere the retrieved surface temperature T is lower than the predicted T_{CO_2} . However, if an EMIRS pixel is only partially covered by CO_2 ice, the retrieved temperature (averaged over the entire pixel footprint) will be higher than the expected CO_2 freezing temperature as part of the observed surface will be unfrozen and warmer than the icy part. Thus, we will only be able to detect surface CO_2 ice that entirely covers an entire EMIRS pixel.

We project the results over the footprints of every pixel into slices of 3 hr of local time, centered on every hour of the day, with a spatial resolution of $0.5^\circ \times 0.5^\circ$. This results in 24 maps containing 3 values:

- “ CO_2 ice” if $T \leq T_{\text{CO}_2}$
- “No CO_2 ice” if $T > T_{\text{CO}_2}$
- “Maybe CO_2 ice” if there is an overlap between pixels footprints associated with “ CO_2 ice” and “No CO_2 ice” temperatures

Then, we use a nearest-pixel interpolation to reconstruct a whole map of the presence of CO_2 ice using `scipy.interpolate.griddata` (Virtanen et al., 2020). Finally, we compute the Gaussian KDE of our data for each map via `scipy.stats.gaussian_kde` (Virtanen et al., 2020), and we flag as “low data density” all the pixels of the final maps associated with a KDE lower than 5×10^{-6} , this value has been tuned experimentally by comparing with the pixel footprints maps. One may note that consecutive maps have a temporal overlap of 1 or 2 hr, but this allows us to have a smoother view to better see when the CO_2 ice starts to form at the surface, without being biased by the choice of the local time binning.

3. Results and Discussion

3.1. Seasonal Variations

Figure 1 shows the seasonal evolution of the median latitudes of the North and South seasonal polar caps between $L_s = 57^\circ$ (MY 36) and $L_s = 11^\circ$ (MY 37) derived from EMIRS observations. This figure has been obtained by computing the median value of the diurnally stable ice maps (cf. Section 2.2) over all longitudes for observations with a high-quality flag. The temporal range encompasses the Northern Summer and the SNPC recession/SSPC progression along with part of the Northern Winter. In addition, we also include in Figure 1 the edges of the polar caps as derived from other orbital instruments (OMEGA & MOC) for previous Martian years (Appéré et al., 2011; Benson & James, 2005; Schmidt et al., 2010), and predicted by the Mars PCM version 6 numeric model assessing a “nominal dust scenario” (Forget et al., 1999, 2022; Naar et al., 2021; Navarro et al., 2014) as a comparison with our MY 36 EMIRS retrievals. For the Mars PCM, the caps areas are defined as the regions where the surface H_2O ice layer at LST = 12 averaged over all longitudes is greater than $10^{-3} \text{ kg m}^{-2}$.

A similar figure representing the seasonal variations of the CO_2 SSPC and SNPC only computed from the median values of the CO_2 local time maps over the day (described in Section 2.3) is provided as Figure S1 in Supporting Information S1, but will not be discussed in details here.

3.1.1. North Polar Cap

We observe that the edge of the SNPC remains stable between 70°N and 75°N between $L_s = 58^\circ$ and $L_s = 143^\circ$ (i.e., during the Northern Summer, when only the perennial North Polar Residual Cap (NPRC) remains), then moves progressively equatorward to reach $\sim 40^\circ\text{N}$ at $L_s = 250^\circ$ and remains there until $L_s = 290^\circ$. Unlike most of the previous studies of the evolution of the SNPC, we do not observe here the recession of the cap but its growth, which will provide a noticeable contribution to our overall understanding of the annual cycle of the SNPC.

Considering uncertainties of $\sim 3^\circ$ in latitude due to the spatial extent of the EMIRS pixels at these latitudes our SNPC retrievals match previous measurements by MOC and OMEGA at $L_s \sim 57^\circ$ (MY 36) and $L_s \sim 11^\circ$ (MY 37) (Appéré et al., 2011; Benson & James, 2005). However, we can see that after the Northern Summer solstice ($L_s = 90^\circ$) a larger discrepancy occurs between our EMIRS retrievals that identify the edge of the cap around

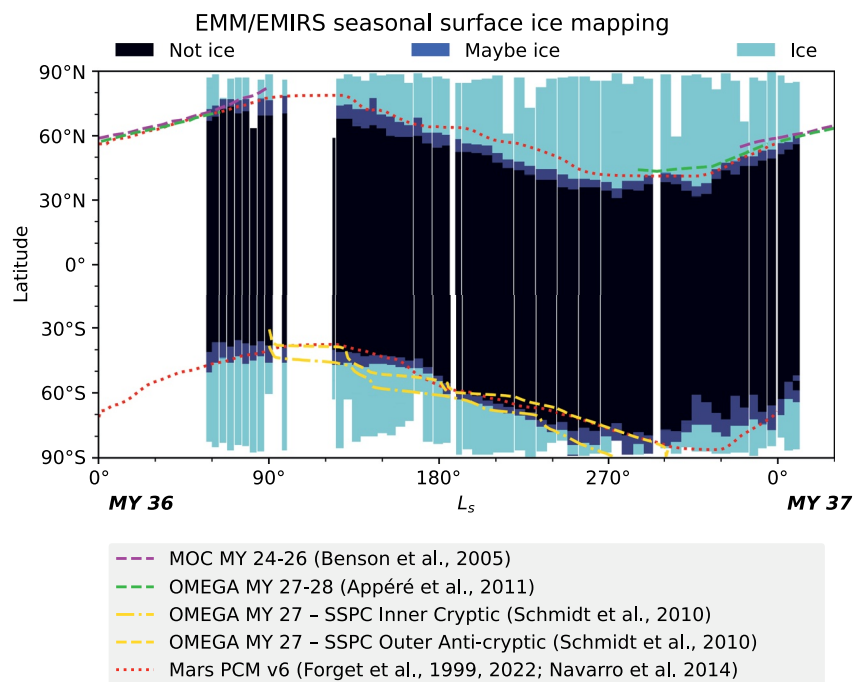


Figure 1. Seasonal variations of the latitudinal median extent of the Martian polar caps from EMM/EMIRS retrievals for diurnally stable surface ice. Only maps with a high-quality flag have been considered here. The red dotted line shows the limits of the H₂O polar caps derived from the Mars Planetary Climate Model (PCM) with a nominal dust scenario (Forget et al., 1999; Navarro et al., 2014), the green and violet dashed lines show the limits of the Seasonal North Polar Cap derived from MOC and OMEGA observations respectively acquired during MY 24–26 (Benson & James, 2005) and MY 27–28 (Appéré et al., 2011), and the yellow dotted lines show the limits of the Seasonal South Polar Cap during its MY 27 recession derived from OMEGA observations (Schmidt et al., 2010).

70°N–75°N and the previous MOC observations that report a limit around 80°N (Benson & James, 2005). By looking at previously obtained maps of the NPRC we can see that it barely reaches 80°N, but we also observe the presence of an additional region of perennial water ice between latitudes 74°N and 80°N for longitudes ranging from 95°E to 245°E (e.g., Langevin et al., 2005; Stcherbinine et al., 2021). Thus, as this area represents ~40% of the longitudes below the polar cap and the EMIRS pixels can span over a few tenths of degrees in longitude under these latitudes, it is likely that the cap boundary that is detected here includes these icy deposits, which explains the mismatch with the MOC data that only consider the NPRC without the frost outlier at lower latitudes that is present between longitudes 95°E and 245°E (Kieffer & Titus, 2001; Langevin et al., 2005).

The comparison with the Mars PCM predictions shows an asymmetry over the year: the PCM SNPC boundary matches previous OMEGA & MOC observations during the SNPC retreat phase ($L_s \sim 330^\circ$ – 80°) and EMIRS retrievals for $L_s = 330^\circ$ – 11° and $L_s = 58^\circ$ – 74° , then it is $\sim 5^\circ$ – 10° above the EMIRS boundary during the expanding phase from $L_s = 185^\circ$ to $L_s = 270^\circ$. The PCM predicts a linear growth from $L_s = 134^\circ$ to $L_s = 170^\circ$ followed by a plateau until $L_s = 195^\circ$ and a second expansion phase until $L_s = 290^\circ$, while our results show a linear expansion from $L_s = 140^\circ$ to $L_s = 250^\circ$ follow by a plateau. The presence of this plateau between $L_s = 170^\circ$ and $L_s = 190^\circ$ in the model which is not observed in the EMIRS retrievals leads to the smaller extent of the SNPC predicted by the PCM compared to our retrievals during the second half of the expansion phase. One may also consider that the PCM has been run here for a “nominal dust scenario” (Forget et al., 1999, 2022; Millour et al., 2022; Montabone et al., 2015), that is, not including the specificities of MY 36. Previous studies (e.g., Calvin et al., 2015; Piqueux et al., 2015) have shown the presence of interannual variations of a few degrees in latitude (typically up to 3–4°) in the extent of the seasonal polar cap deposits for the same values of L_s . Thus, considering the uncertainties of $\sim 3^\circ$ in latitude in our retrievals due to the spatial extent of EMIRS pixels footprints, the discrepancies observed between our retrievals and the PCM may be reflecting interannual variations of the evolution of the SNPC, with a faster expanding phase in MY 36 compared to the nominal scenario. Otherwise, the model matches the EMIRS and OMEGA observations during the Northern Winter ($L_s \sim 290^\circ$) and the EMIRS retrievals during Northern

Summer ($L_s \sim 90^\circ$), thus also including the Northern icy deposits outside the primary perennial polar cap with the longitudinal average.

3.1.2. South Polar Cap

The Mars SSPC is highly asymmetric and can be divided into two regions: the “cryptic” and the “anti-cryptic” (Kieffer et al., 2000; Schmidt et al., 2010). They do not have the same sublimation rate during the polar cap recession and do not extend to the same latitudes. In particular, during the Southern Summer, the perennial South Polar Residual Cap (SPRC) is only present in the “anti-cryptic” region (Langevin et al., 2007; Schmidt et al., 2010). Thus, considering the methodology used here to map the seasonal evolution of the polar caps with EMIRS, we expect our EMIRS latitudinal boundary for the SPRC to be located between the “cryptic” and “anti-cryptic” ones derived from OMEGA observations (Schmidt et al., 2010).

Indeed, we observe a good agreement between the EMIRS results and the OMEGA observations during the Southern Winter ($L_s \sim 95^\circ$ – 130°) and the second half of the SSPC recession ($L_s \sim 200^\circ$ – 295°). Between $L_s = 133^\circ$ and $L_s = 190^\circ$ the EMIRS SSPC boundary is detected up to 6° northern than the external OMEGA edge (“outer anti-cryptic”). However, the boundaries derived by OMEGA are “crocus” lines (Schmidt et al., 2009, 2010), that is, the limits of the CO_2 deposits, while our methods capture equally all the ices that may be present on the surface (CO_2 or H_2O). Figure 12 from Langevin et al. (2007) shows that between $L_s = 130^\circ$ and $L_s = 155^\circ$ the seasonal H_2O deposits extend outside the CO_2 ones, with detections up to $\sim 45^\circ\text{S}$ which is consistent with our observations.

Regarding the limits predicted by the Mars PCM, we observe that it matches our detections from $L_s = 58^\circ$ to $L_s = 92^\circ$ and for $L_s \sim 150^\circ$ – 225° . Then, the PCM line is at slightly more equatorial latitudes (up to 5°) compared to our detections but matches the “outer anti-cryptic” limit derived by OMEGA (Schmidt et al., 2009, 2010) from $L_s \sim 225^\circ$ to $L_s \sim 265^\circ$. Considering the uncertainties of $\sim 3^\circ$ in latitude in our retrievals, these discrepancies may be reflecting the inter-annual variations of the SSPC retreat (Piqueux et al., 2015), especially as the PCM run does not include the specific scenario of MY 36. Or it could be the consequence of a mix between the cryptic and anti-cryptic areas of the SSPC while computing the median over all longitudes, with a predominance of the cryptic area in the EMIRS retrievals and the anti-cryptic area in the PCM results, which can be related to the higher uncertainties in term of longitude in the EMIRS retrievals at high latitudes due to the spatial extension of the pixel footprints.

3.2. Diurnal Variations

From the surface CO_2 ice maps generated for bins of 3 hr of LST on a 4-orbits basis as described in Section 2.3, we compute for each bin of local time a map of the percentage of CO_2 frost detections over wider ranges of L_s , considering only the pixels flagged as “high data density.”

Figure 2 shows the frequency of CO_2 surface ice detections for each local time between $L_s = 152^\circ$ and $L_s = 203^\circ$. We can see that apart from the polar caps, CO_2 ice is detected at night under equatorial and mid-latitude regions mostly around 100°W (Tharsis region), and also around 40°E (Arabia Terra), which corresponds to low thermal inertia areas (Putzig & Mellon, 2007) where nighttime surface CO_2 frost have been detected through MCS or THEMIS measurements (Khuller et al., 2021; Lange et al., 2022; Piqueux et al., 2016). Regarding the daily evolution of these CO_2 frost deposits, we observe that it starts to appear at midnight (panel a) and last after 6 a.m. to disappear during daytime, with a maximum of intensity between LST of 03:00 and 06:00 (panel b). Condensation of the CO_2 occurs during the second half of the Martian night until sunrise.

Figure 3 shows the frequency of detections of surface CO_2 ice between 3 and 6 a.m. by EMIRS from $L_s = 6^\circ$ to $L_s = 290^\circ$ for 6 temporal bins of $\sim 50^\circ$ of L_s . One may note that panel a of Figure 3 exhibits some large and awkwardly shaped detection spots of CO_2 ice, which are due to interpolation effects on maps generated with a lower data coverage at the beginning of the mission. We can see that our detections of non-polar surface CO_2 frost vary with the L_s : while they remain located in the same two regions mentioned above, the frequency of detections along with the size of the area covered by the frost evolve. Indeed, equatorial and mid-latitudes CO_2 frost is mostly detected for $6^\circ \leq L_s \leq 19^\circ$ and $120^\circ \leq L_s \leq 203^\circ$ (panels a, c & d). Panel b ($49^\circ \leq L_s \leq 100^\circ$) corresponds to the end of Spring and early Summer in the Northern hemisphere, that is, the period of the year with the highest temperatures in the Northern hemisphere. Thus, as the regions where nighttime CO_2 frost is usually detected under midlatitudes are located in the Northern hemisphere, there are almost no detections outside the South polar

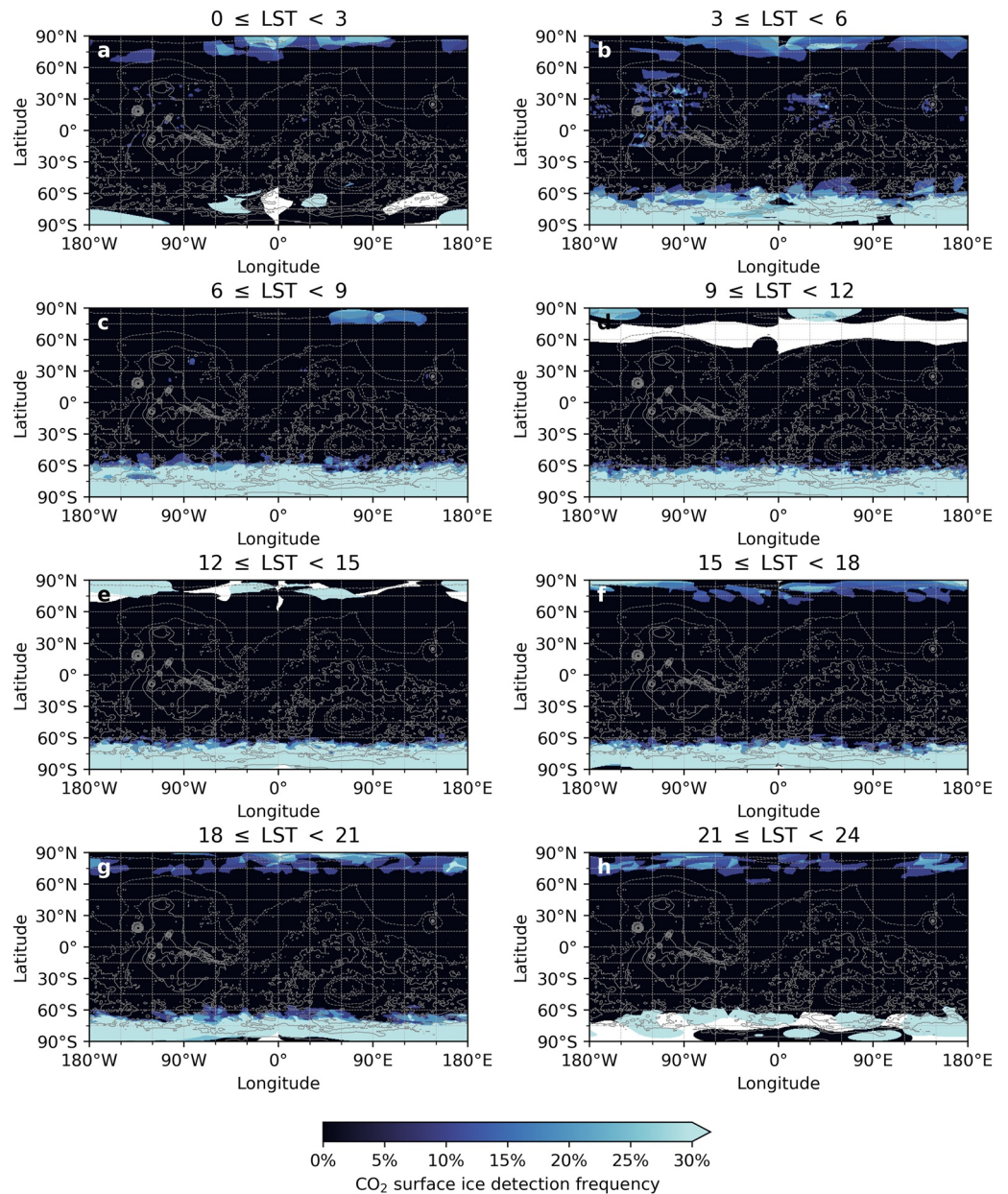


Figure 2. Frequency of detections of surface CO₂ ice from $L_s = 152^\circ$ to $L_s = 203^\circ$ (EMM orbits 156 to 195) for bins of 3 hr of local solar time. Regions without data coverage are in white. Pixels flagged as “low data density” are ignored here. We observe that CO₂ frost is detected here mostly between local solar times of 03:00 and 06:00.

cap during this period. Even the retrieved nighttime temperatures are not low enough to allow the formation of CO₂ ice at the surface of the planet in these regions. This formation of CO₂ frost late during the night with a quick sublimation at sunrise was expected from the models (e.g., Lange et al., 2022) with implications on the formation of gullies and slope streaks (Khuller et al., 2021; Lange et al., 2022; Pilorget & Forget, 2016). But as previous nighttime observations were conducted only around 03:00, this is the first time we are able to confirm it with direct observations of the diurnal cycle of the surface CO₂ frost at low latitudes.

Similarly, we also report almost no midlatitudes CO₂ frost detections between $L_s = 203^\circ$ and $L_s = 290^\circ$ (panels e and f). This period corresponds to the Southern spring and summer, but also to the part of the Martian Year where the planet is closer to the Sun as the perihelion is reached at $L_s = 251^\circ$. Thus, because of the eccentricity of the Martian orbit, the global temperature at the surface of the planet is significantly higher during the second half

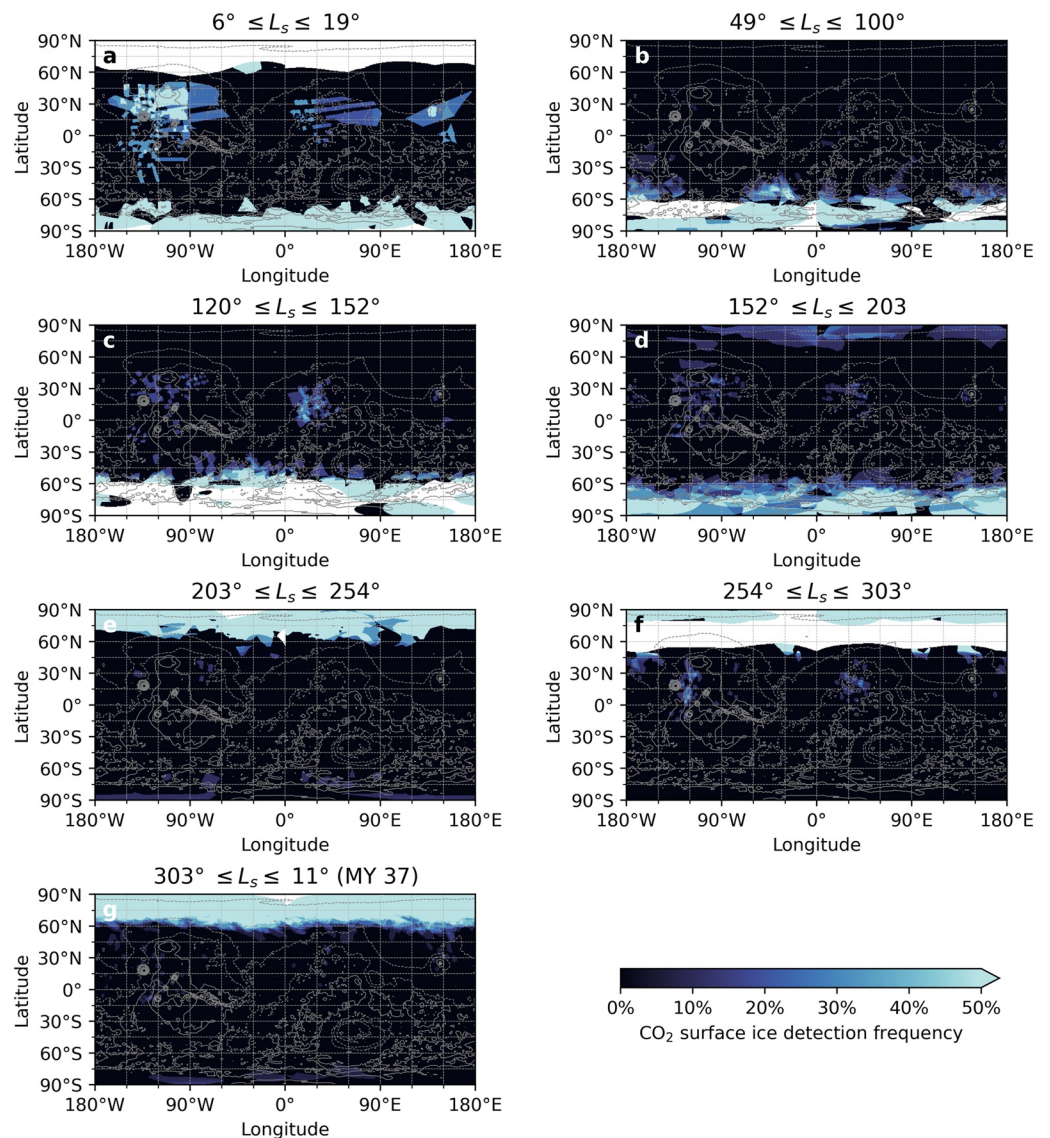


Figure 3. Frequency of detections of surface CO₂ ice between local solar times 03:00 and 06:00 for bins of $\sim 50^\circ$ of L_s over MY 36. Regions without data coverage are in white. Pixels flagged as “low data density” are ignored here.

of the year (e.g., Bell, 2008; Smith, 2006), which explains the fewer nighttime CO₂ frost detections at equatorial and midlatitudes.

This seasonal trend was previously observed by MCS and THEMIS (Khuller et al., 2021; Piqueux et al., 2016). However, one may note the fewer frost detections in our EMIRS results, which may be related to the difference in terms of spatial resolution between the instruments. Indeed, the spatial resolution of the EMIRS pixels is comprised between 100 and 300 km (Edwards et al., 2021), which is much larger than the few km of the MCS footprints (McCleese et al., 2007) or the 100 m of THEMIS (Christensen et al., 2004). Plus, the detections made by both MCS and THEMIS reveal that the CO₂ frost detections may be localized to areas that are sub-pixel at the EMIRS resolution (Khuller et al., 2021; Piqueux et al., 2016). Thus, if an EMIRS pixel is only partially covered by surface CO₂ frost, the retrieved temperature, averaged over the entire footprint, will be higher than the CO₂ freezing temperature as a portion of the footprint will be unfrozen. This results in a lower sensibility in the surface CO₂ frost retrieving process, but also strengthen our detections as they testify of the presence of CO₂ ice over all our pixel footprints.

4. Conclusion

In this paper, we present the results of our study on the monitoring of the Martian surface ices with the EMIRS instrument onboard EMM. From the variation of the surface temperature over the day, we developed a method to automatically map the presence of diurnally stable ice at the surface, which allows us to monitor the seasonal variations of the two polar caps. Then, based on the method previously developed for MCS data (Piqueux et al., 2016), we use the surface temperature to detect and map the presence of CO₂ frost at the surface of the planet. We confirm the diurnal apparition and disappearance of surface CO₂ frost under equatorial latitudes previously reported by Piqueux et al. (2016) and observationally monitor for the first time its evolution as a function of the local time thanks to the unique orbit of the EMM probe.

We monitor the evolution of the seasonal polar caps from $L_s = 57^\circ$ (MY 36) to $L_s = 11^\circ$ (MY 37), with a temporal resolution of 5° of L_s (10 Earth days). Plus, the large-scale view of EMIRS and the automatization of our method allow us to continuously and simultaneously monitor the annual variations of both polar caps.

Also, we are able to observe for the first time the apparition and disappearance of low-latitude nighttime CO₂ frost at the surface of the planet, thanks to the unique ability of EMM instruments to provide full coverage in terms of local time. CO₂ ice is detected at the surface down to the equator around spring and fall equinoxes in the second half of the night (essentially between 3 and 6 a.m.) with a quick sublimation at sunrise, which confirms previous model expectations.

Data Availability Statement

The SPiP module is freely available on GitHub at <https://github.com/NAU-PIXEL/spip> (Stcherbinine, 2023). Data from the Emirates Mars Mission (EMM) are freely and publicly available on the EMM Science Data Center (SDC, <http://sdc.emiratesmarsmission.ae>). This location is designated as the primary repository for all data products produced by the EMM team and is designated as long-term repository as required by the UAE Space Agency. The data available (<http://sdc.emiratesmarsmission.ae/data>) include ancillary spacecraft data, instrument telemetry, Level 1 (raw instrument data) to Level 3 (derived science products), quicklook products, and data users guides (<https://sdc.emiratesmarsmission.ae/documentation>) to assist in the analysis of the data. Following the creation of a free login, all EMM data are searchable via parameters such as product file name, solar longitude, acquisition time, sub-spacecraft latitude & longitude, instrument, data product level, etc.

Data products can be browsed within the SDC via a standardized file system structure that follows the convention: /emm/data/< Instrument>/< DataLevel>/< Mode>/< Year>/< Month>.

Data product filenames follow a standard convention: emm_< Instrument>_< DataLevel>_< Start TimeUTC>_< OrbitNumber>_< Mode>_< Description>_< Kernel-Level>_< Version>.< FileType>.

EMIRS data and users guides are available at: <https://sdc.emiratesmarsmission.ae/data/emirs>.

The Mars PCM v6 and the MCD are available from <http://www-mars.lmd.jussieu.fr>.

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