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RESEARCH ARTICLE

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Key Points:

- The ARPEGE-Climat 6.3 performance in modeling a stable boundary layer is assessed in a single-column model/large-eddy simulation comparison
- The use of history matching with iterative refocusing indicates turbulence parameterization deficiencies due to poor calibration
- Validation of the statistical framework and guidelines for its use in parameterization development and calibration are discussed

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Modeling the GABLS4 Strongly-Stable Boundary Layer With a GCM Turbulence Parameterization: Parametric Sensitivity or Intrinsic Limits?

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Abstract The representation of stable boundary layers (SBLs) still challenges turbulence parameterizations implemented in current weather or climate models. The present work assesses whether these model deficiencies reflect calibration choices or intrinsic limits in currently-used turbulence parameterization formulations and implementations. This question is addressed for the CNRM atmospheric model ARPEGE-Climat 6.3 in a single-column model/large-eddy simulation (SCM/LES) comparison framework, using the history matching with iterative refocusing statistical approach. The GABLS4 case, which samples a nocturnal strong SBL observed at Dome C, Antarctic Plateau, is used. The standard calibration of the ARPEGE-Climat 6.3 turbulence parameterization leads to a too deep SBL, a too high low-level jet and misses the nocturnal wind rotation. This behavior is found for low and high vertical resolution model configurations. The statistical tool then proves that these model deficiencies reflect a poor parameterization calibration rather than intrinsic limits of the parameterization formulation itself. In particular, the role of two lower bounds that were heuristically introduced during the parameterization implementation to increase mixing in the free troposphere and to avoid runaway cooling in snow- or ice-covered region is emphasized. The statistical tool identifies the space of the parameterization free parameters compatible with the LES reference, accounting for the various sources of uncertainty. This space is non-empty, thus proving that the ARPEGE-Climat 6.3 turbulence parameterization contains the required physics to capture the GABLS4 SBL. The SCM framework is also used to validate the statistical framework and a few guidelines for its use in parameterization development and calibration are discussed.

Plain Language Summary During the night or in snow- or ice-covered region, a stable atmospheric boundary layer (SBL) often develops. Their representation still challenges turbulence parameterizations implemented in numerical weather or climate models. The present work assesses whether the ARPEGE-Climat atmospheric model deficiencies reflect calibration choices or intrinsic limits in its turbulence parameterization using a statistical approach from the Uncertainty Quantification community. A single-column version of the model is evaluated on the GABLS4 case, a nocturnal strong SBL observed at Dome C, Antarctic plateau, and compared to high-resolution simulations. The standard calibration of the ARPEGE-Climat 6.3 turbulence parameterization leads to a too deep SBL and an incorrect wind pattern and so for different vertical resolutions. The statistical tool proves that these model deficiencies are rectified with proper calibration of the turbulence parameterization. In particular, it is shown that two lower bounds, introduced to increase turbulent mixing, are key to capture the GABLS4 SBL. Finally, the potential and relevance of the Uncertainty Quantification approach applied to the single-column model/large-eddy simulation comparison framework for the calibration of climate models are highlighted and few guidelines for its use are proposed.

1. Introduction

In the atmospheric boundary layer, stably stratified conditions generally develop above a surface colder than the overlying air. These Stable Boundary Layers (SBLs) often result from the advection of warm air over a cold surface or from the cooling of the surface. They are frequently observed over ice or snow surfaces (e.g., polar regions, high-latitude continental regions in wintertime) and over land during nighttime. Their development and intensity (e.g., the vertical stratification) are also strongly modulated by the atmospheric synoptic conditions. Nights with weak or no cloud cover at all also favor SBL occurrence, as such conditions

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enhance the radiative cooling of the surface (Mahrt, 1998). In contrast, the occurrence of strong near-surface wind reduces the SBL stratification, or even inhibits their development, through the maintenance of significant mechanical mixing (e.g., Van de Wiel et al., 2012).

SBLs can be classified according to the intensity of their stratification (at first order the vertical gradient of potential temperature), ranging from weak SBLs in which the turbulence remains significant, to strong SBLs, in which the turbulence become intermittent or even disappears (e.g., Acevedo et al., 2016; Mahrt, 1998). In the latter conditions, a mechanical decoupling between the atmosphere and the surface can occur (Derbyshire, 1999): the temperature inversion close to the surface becomes driven by radiation and soil diffusion and the surface turbulent heat flux cannot sustain the surface energy demand enhanced by a strong net surface radiative cooling (e.g., Van de Wiel et al., 2012). Such strong SBLs mostly occur under clear-sky and weak wind conditions, with a strong increase of the near-surface temperature inversion below a critical wind speed (e.g., Vignon, van de Wiel et al., 2017b).

The representation of SBLs in General Circulation Models such as operational climate and weather models is a challenge (e.g., Holtslag et al., 2013): the turbulence is particularly weak and sometimes intermittent (e.g., Mauritsen & Svensson, 2007), and interacts with other small-scale processes (e.g., gravity waves—Steenefeld et al., 2008; Tsiringakis et al., 2017). Under the umbrella of the global energy and water cycle exchanges (GEWEX) project, the GEWEX atmospheric boundary layer study (GABLS) has initiated four model intercomparison projects (Bazile et al., 2015; Bosveld et al., 2014; Cuxart et al., 2006; Svensson et al., 2011) to evaluate and improve the SBL representation in weather and climate models. So far, the GABLS intercomparison exercises revealed:

1. Large-eddy simulations (LES) are able to consistently capture the main properties of stable boundary layers, at least when their resolution is below a few meters for weak to moderate SBL (e.g., Beare et al., 2006) or 1 m for strong SBL (Couvreur, Bazile et al., 2020a). As a result, such LESs provide relevant process-level information to evaluate turbulence parameterizations in an LES/Single-Column Model (SCM) comparison framework (e.g., Randall et al., 1996, 2003)
2. State-of-the-art turbulence parameterizations, such as those with a 1.5-order turbulence closure, are able to reasonably capture the physics of SBLs for a wide range of forcing (e.g., Baas et al., 2018; Cuxart et al., 2006; Vignon, Hourdin et al., 2017a)
3. Nevertheless, current weather and climate models still simulate SBLs that are too deep, with surface drag that is too strong, low-level jets that are too weak and too high and wind veering with height that is too weak (e.g., Cuxart et al., 2006; Holtslag et al., 2013)

The apparent contradiction between the two last conclusions results from the calibration of weather or climate models which, so far, has required an increased turbulent mixing to reduce the activity of synoptic systems, and thereby improve operational scores (e.g., Sandu et al., 2013), or to prevent runaway surface cooling through long-term mechanical decoupling with the atmosphere (Derbyshire, 1999). Such a calibration probably reflects the lack of mixing due to processes that are currently not accounted for in weather and climate models (e.g., surface heterogeneities, internal gravity waves, meso-scale variability, impact of subgrid orography).

Recently Vignon, Hourdin et al. (2017a); Vignon et al. (2018) and Hourdin, Rio et al. (2020a) showed that it is possible to achieve a reasonable representation of SBLs in a climate model (LMDZ), while maintaining reasonable large-scale performance. Starting with an SCM framework built on the very stable boundary layer of GABLS4 (Bazile et al., 2015), Vignon, Hourdin et al. (2017a) underline the importance of (i) the coupling with the surface (snow albedo and thermal inertia) and (ii) the turbulent mixing thresholds usually used in current operational turbulence parameterizations (e.g., for the mixing length or in stability functions). More specifically, the appropriate calibration of surface (snow) properties and the removal of those thresholds in the turbulence parameterization allows the LMDZ SCM to capture well the strong temperature gradient close to the surface (in the first 20 meters), observed at Dome C, Antarctica, during an austral summer night. Vignon et al. (2018) further compare 3D LMDZ simulations with local observations collected at Dome C. On the one hand, the SCM improved results are consistently reported in this 3D configuration, stressing the relevance of the SCM framework for developing and calibrating parameterizations. On the other hand, the new version of the LMDZ model adequately reproduces the annual cycle of Dome

C, the two main SBL regimes discussed above, and preserves satisfying large-scale skills. Finally, the vertical resolution in the lower part of the boundary layer is also shown to be critical for capturing SBLs that cover only a few tens of meters (see also Steeneveld et al., 2006).

Following Vignon, Hourdin et al. (2017a) and Vignon et al. (2018), the general objective of the present work is to document the performance of the CNRM climate atmospheric model, namely ARPEGE-Climat 6.3 (Roehrig et al., 2020), to represent SBLs. The focus is here on its turbulence parameterization, which is based on the work of Cuxart et al. (2000). The parameterization of a given process seeks to represent its effects on the large-scale (or resolved) state of the model. It is based on a set of physical theories or empirical relationships to numerically describe the subgrid-scale processes and their effects. Parameterizations introduce a number of constants, called free parameters in the following, which are often difficult to constrain with observations or other references. A parameterization can thus be seen as a function of the model state variables and of these free parameters. Their calibration, or “tuning”, is a critical step in model development for weather or climate applications (e.g., Hourdin et al., 2017). In the present paper, we therefore propose to address the following specific question: Is it possible to calibrate the ARPEGE-Climat turbulence parameterization to achieve a satisfying representation of SBLs, especially those with a strong thermal stratification? In other words, does the ARPEGE-Climat turbulence parameterization contain the required physics to represent appropriately strong SBLs?

Cuxart et al. (2006) shows that the current turbulence parameterization of ARPEGE-Climat 6.3 is able to capture the main properties of the moderate SBL studied during the first GABLS exercise. We seek to extend this result to the strongly-stratified SBL of the GABLS4 nocturnal phase. We rely on SCM simulations, which have been shown relevant for 3D model configuration (Gettelman et al., 2019; Hourdin et al., 2013; Neggers, 2015; Vignon et al., 2018). We also make use of GABLS4 LESs as references, as they have been shown to capture well the properties of the GABLS4 nocturnal phase (Couvreur, Bazile et al., 2020a). Following Couvreur, Hourdin, et al. (2020b), we use statistical tools developed in the Uncertainty Quantification community, in particular the history matching with iterative refocusing proposed by Williamson et al. (2013) and applied to the SCM/LES comparison. This tool provides the sensitivity analysis of our turbulence parameterization to its free parameters and identifies which part of the full free parameter space provides SCM simulations consistent with the chosen reference, accounting for the various sources of uncertainty (Couvreur, Hourdin et al., 2020b; Williamson et al., 2013, 2017). Thus, we do not seek to optimize the ARPEGE-Climat turbulence parameterization over the GABLS4 SBL in the traditional meaning of searching for the best set of free parameters to reproduce this SBL. In contrast, our objective is to document which sets of free parameters, if such sets exist, provide an appropriate representation of the GABLS4 SBL. This information can then be further used in the whole process of the ARPEGE-Climat calibration and combined with similar information based on other 1D cases or on the full 3D model configuration. Such an approach should ensure that the full model calibration will keep a reasonable process-level behavior for the GABLS4 SBL (see Couvreur, Hourdin et al., 2020b; Hourdin, Williamson et al., 2020b, for a more detailed discussion).

Section 2 introduces the ARPEGE-Climat 6.3 atmospheric model and its turbulence parameterization. The relevant free parameters of the parameterization to be used for calibration are emphasized. Section 3 presents the case study used for the SCM/LES intercomparison and the LES results that serve as a reference. Section 4 describes the statistical framework. Section 5 details the results obtained for two different configurations of the ARPEGE-Climat 6.3 SCM. A third one is rapidly discussed, with more detail provided in Appendix A. Section 6 discusses several aspects of the methodology and section 7 finally concludes the present study.

2. ARPEGE-Climat 6.3

ARPEGE-Climat is a global atmospheric model developed at CNRM for climate studies. Its latest version (6.3, Roehrig et al., 2020) is the atmospheric component of the CNRM ocean-atmosphere climate model CNRM-CM6-1 (Voldoire et al., 2019), and Earth System model CNRM-ESM2-1 (Séférian et al., 2019). The following work uses the single-column model (SCM) version of ARPEGE-Climat (e.g., Abdel-Lathif et al., 2018), in the context of the GABLS4 framework (see Section 3.1). The model physical package is fully

described in Roehrig et al. (2020) and therefore we only insist hereafter on the model features relevant for the present study. ARPEGE-Climat 6.3 standard vertical grid consists of 91 vertical levels, following the progressive hybrid σ -pressure discretization of Simmons and Burridge (1981). The altitude of the first 5 model levels is approximately 8.5, 29, 55, 91, and 132 m. The model timestep is 15 min. A version of ARPEGE-Climat 6.3 with higher vertical resolution (2 m up to 400 m) is also used in Section 5.1. To prevent instabilities, the timestep of this version is reduced to 60 s. Note that the use of this 60-s timestep in the 91-level version of ARPEGE-Climat does not impact much the results of the present work. As described in Section 3, the SCM configuration is run on a idealized case (stable boundary layer, no moisture, no radiation), in which only the turbulence and surface flux parameterizations are activated. These parameterizations are described hereafter, in a dry context. Note that in the following, the names of the free parameters of the turbulence scheme that will be used for model calibration are written in bold.

2.1. Turbulence Parameterization

The turbulence scheme used in ARPEGE-Climat 6.3 follows the work of Redelsperger and Sommeria (1982, 1986); and Cuxart et al. (2000). It relies on the eddy diffusivity approach, coupled to a prognostic equation for the grid-scale-averaged turbulence kinetic energy (TKE) \bar{e} . Given the standard horizontal resolution of ARPEGE-Climat ($\mathcal{O}(100\text{km})$), only the vertical component of turbulent mixing is parameterized. For any variable ψ impacted by turbulent mixing (e.g., wind component u and v , potential temperature θ), the associated second-order turbulent flux $\overline{w'\psi'}$ reads (primes denote fluctuations with respect to the grid-scale average, noted $\bar{\psi}$):

$$\overline{w'\psi'} = -K_\psi \frac{\partial \bar{\psi}}{\partial z} ; K_\psi = \alpha_\psi \mathbf{CM} L_m \sqrt{\bar{e}} \phi_\psi \quad (1)$$

where α_ψ and \mathbf{CM} are free parameters of the parameterization, L_m is the mixing length, and ϕ_ψ is a stability function. ϕ_ψ is taken to 1 for momentum and turbulence kinetic energy ($\psi \in \{u, v, e\}$). For the potential temperature θ , the following formulation is used:

$$\phi_\theta = \frac{1}{1 + C \frac{g}{\theta} \frac{L_m^2}{\bar{e}} \frac{\partial \theta}{\partial z}} \text{ where } C \text{ is a free parameter.} \quad (2)$$

In Equation 1, \mathbf{CM} modulates all turbulent fluxes in the same way. α_u and α_v are taken to 1, and α_θ is the inverse Prandtl number in neutral condition (i.e. when $\phi_\theta = 1$). In the following, α_e and α_θ will be referred to as **AE** and **AT**, respectively. The turbulent fluxes defined by Equation 1 are computed only above the surface layer, namely from the first model half level. The surface flux boundary condition is provided by the surface flux parameterization (see Section 2.2). Eddy diffusivity coefficients K_ψ depend on the intensity of \bar{e} . The time evolution of \bar{e} is given by:

$$\frac{\partial \bar{e}}{\partial t} = -\frac{1}{\rho} \frac{\partial}{\partial z} (\overline{\rho e' w'}) - \left(\overline{u' w'} \frac{\partial \bar{u}}{\partial z} + \overline{v' w'} \frac{\partial \bar{v}}{\partial z} \right) + \frac{g}{\theta} \overline{w' \theta'} - \frac{\bar{e} \sqrt{\bar{e}}}{L_\epsilon} \quad (3)$$

where ρ is the air density, g is the gravity acceleration, and L_ϵ the dissipation length. L_ϵ is assumed to be proportional to the mixing length: $L_\epsilon = \mathbf{CE} L_m$, with \mathbf{CE} a free parameter. The mixing length follows the non-local formulation of Bougeault and Lacarrère (1989) and reads

$$L_m^{\text{BL89}} = \left[\frac{1}{2} \left((L_{\text{up}})^{-2/3} + (L_{\text{down}})^{-2/3} \right) \right]^{-3/2} \quad (4)$$

where L_{up} and L_{down} are, respectively, the maximum upward and downward displacements a parcel can travel within the ambient thermal stratification, given its turbulence kinetic energy, and accounting only for the work of its buoyancy. Close to the surface, the mixing length is supposed to be larger than κz where $\kappa = 0.4$

Table 1
Free Parameters of the Turbulence Parameterization.

	CM	AE	AT	CE	LMIN	KOZMIN	ZMAX
ARPEGE-Climat 6.3	0.126	2.70	1.13	0.85	10.0	5.10^{-3}	200
Cuxart et al. (2000)	0.0667	6.0	2.5	0.70	10.0	–	–
Lower bound	0.05	0.50	0.20	0.33	0.0	0.0	30
Upper bound	0.30	6.00	3.00	5.00	10.0	5.10^{-3}	400

The values in the standard version of ARPEGE-Climat are those from Cheng et al. (2002). Those from the work of Cuxart et al. (2000) are also included. The bottom two lines provide the range of values that we explore for each parameter.

is the Von Kármán constant. A minimum mixing length, **LMIN**, is introduced to maintain a minimum vertical mixing in stable boundary layers, such that the mixing length L_m finally reads:

$$L_m = \max \left[L_m^{\text{BL89}}, \min(\text{LMIN}, \kappa z) \right] \quad (5)$$

In case of shallow stable boundary layer, another lower bound, which applies mainly close to the surface, is introduced directly on the turbulent fluxes, to avoid runaway cooling of the surface (especially in snow- or ice-covered regions):

$$\overline{w'\psi'} = \max \left[-K_\psi \frac{\Delta \bar{\psi}}{\Delta z}, \alpha_\psi \left(\text{KOZMIN} \left(1 - \frac{z}{\text{ZMAX}} \right) \right) \Delta \bar{\psi} \right] \quad (6)$$

where **KOZMIN** and **ZMAX** are two free parameters, and $\Delta \bar{\psi}$ is the vertical difference of $\bar{\psi}$ between two consecutive model layers (distant of Δz). Above **ZMAX**, no lower bound is used. This formulation, through $\Delta \bar{\psi}$ and Δz depends on the vertical discretization of the model (see also Section 5). This empirical formulation is highly questionable, especially in terms of vertical resolution dependence. Nevertheless, we considered it as it is currently implemented in ARPEGE-Climat 6.3. The present work provides the opportunity to discuss its behavior and possible detrimental impacts with respects to stable boundary layers.

The turbulence parameterization thus includes several free parameters that have to be calibrated. Eight parameters have been identified here. The calibration of the parameterization consists in choosing a value for each of them, accounting for both parameterization performance and physical constraints. In the standard configuration of ARPEGE-Climat 6.3 (see Roehrig et al., 2020), the parameter values follow the work of Cheng et al. (2002) except the parameters **LMIN**, **KOZMIN** and **ZMAX** that were introduced in the course of the parameterization implementation in ARPEGE-Climat and set in a more empirical way. Table 1 provides the values of these parameters as currently used in ARPEGE-Climat 6.3 as well as those initially proposed in Cuxart et al. (2000). Note that the parameter C in Equation 2 is set to 0.143. As the model is not much sensitive to it, the following work does not consider this parameter.

2.2. Surface Flux Parameterization

The SCM configuration of ARPEGE-Climat will be used in two different configurations with respects to the surface boundary conditions, one with prescribed surface sensible heat flux and one with prescribed surface temperature. In both, the roughness lengths for momentum (z_0) and heat (z_{0h}) are prescribed.

2.2.1. Configuration with Prescribed Surface Sensible Heat Flux

The friction velocity u_* is computed following Paulson (1970):

$$u_* = \frac{\kappa \bar{U}_1}{\ln\left(\frac{z_1}{z_0}\right) - \varphi\left(\frac{z_1}{L_{MO}}\right) + \varphi\left(\frac{z_0}{L_{MO}}\right)} \quad (7)$$

where \bar{U}_1 is the wind intensity ($\bar{U}_1 = \sqrt{\bar{u}_1^2 + \bar{v}_1^2}$) at the first model level of altitude z_1 , and L_{MO} is the Monin-Obukhov length. The similarity function φ is given by Paulson (1970). As L_{MO} depends on u_* , the computation is done iteratively, initialized from a neutrally-stable state (i.e. $L_{MO} = \infty$, knowing $\varphi(0) = 0$). The surface momentum flux is finally given by:

$$F_u = \bar{\rho} u_*^2 \quad (8)$$

2.2.2. Configuration with Prescribed Surface Temperature

In this configuration, the standard version of the ARPEGE-Climat surface scheme is used. The surface momentum and heat fluxes are computed based on the formulations of Mascart et al. (1995) involving the bulk Richardson number Ri_b^0 :

$$Ri_b^0 = \frac{gz_1(\bar{\theta}_1 - \theta_s)}{\frac{1}{2}(\bar{\theta}_1 + \theta_s)\bar{U}_1^2} \quad (9)$$

where $\bar{\theta}_1$ and $\bar{\theta}_s$ are the potential temperature at the first model level and at the surface, respectively. A critical Richardson number $Ri_c = 0.1$ is used as an upper bound of the bulk Richardson number: $Ri_b = \min(Ri_b^0, Ri_c)$. Note this critical Richardson number value is most likely not appropriate in the context of strongly stable boundary layers. It may be considered as a free parameter in future work. The exchange coefficients for momentum and heat in the case of stable states ($Ri_b > 0$) following Mascart et al. (1995) and Noilhan and Mahfouf (1996) read:

$$C_d = \frac{\kappa^2}{\left(\ln\left(\frac{z}{z_0}\right)\right)^2} \frac{1}{1 + \frac{B_1 Ri_b}{\sqrt{1 + B_2 Ri_b}}} \quad (10)$$

$$C_h = \frac{\kappa^2}{\left(\ln\left(\frac{z}{z_0}\right)\right)^2} \left(\frac{\ln(z/z_0)}{\ln(z/z_{0h})}\right) \left(\frac{1}{1 + B_3 Ri_b \sqrt{1 + B_2 Ri_b}}\right) \quad (11)$$

and are used to compute the surface fluxes:

$$F_u = \bar{\rho} C_d \bar{U}_1^2 \quad \text{and} \quad F_s = \bar{\rho} C_p C_h \bar{U}_1 (\theta_s - \bar{\theta}_1) \quad (12)$$

where C_p is the heat capacity of air at constant pressure. Note that the present surface flux parameterizations include internal free parameters ($B_1 = 10$, $B_2 = 5$, $B_3 = 15$), which are not considered in the following analysis. They are possibly critical for SBLs (e.g. Vignon, Hourdin et al., 2017a), and will be analyzed in a future work.

3. Experimental Setup and Reference Simulations

3.1. The GABLS4 Framework

The present study is based on the GABLS4 model intercomparison case (Bazile et al., 2014, 2015; Couvreux, Bazile et al., 2020a). It focuses on an austral summer diurnal cycle of the boundary layer at Dome C, Antarctic Plateau (123.3° E, 75.1° S, 3,223 m above sea level, local time (LT) = UTC+8 h) as observed from December 11, 0800 LT to December 12, 0800 LT. During that day, the boundary layer evolved from a 400-m deep convective regime during daytime to a nighttime very stable regime covering a depth shallower than 30 m (Vignon, Hourdin et al., 2017a).

The GABLS4 model intercomparison encompasses three different stages. The first one is dedicated to the intercomparison of SCMs with an interactive snow surface scheme. The second stage prescribes observed surface temperature, thus suppressing several feedbacks between the atmosphere and the surface. The third stage consists in an idealization of GABLS4 stage 2, in which no moisture, no radiation, no large-scale subsidence and no large-scale advection of temperature are considered.

Couvreux, Bazile et al. (2020a) emphasize that the representation of the full GABLS4 diurnal cycle is a challenge for LES models as it requires a large domain for the 400-m deep daytime convective boundary layer and a very high-resolution for the 30-m deep nocturnal very stable boundary layer. Therefore, Couvreux, Bazile et al. (2020a) proposed a complementary setup focused on the GABLS4 nocturnal stable phase, starting at the end of the convective period (1800 LT, i.e. 10 h after the start of the original version of GABLS4 stages) and covering 11 h (until 0500 LT). This new setup, referred to as GABLS4-Stage3-10 h, corresponds to the setup used in the present work.

The initial conditions are obtained from the ensemble mean of three LESs that took part to the GABLS4 Stage 3 LES intercomparison (Couvreux, Bazile et al., 2020a). GABLS4-Stage3-10 h uses the same large-scale forcing as in GABLS4 Stage 3, which thus only includes a large-scale horizontal pressure gradient through a prescribed geostrophic wind. This geostrophic wind is constant in time ($u_g = 1.25 \text{ m s}^{-1}$ and $v_g = 4.5 \text{ m s}^{-1}$) and along height. GABLS4 Stage 3 (and thus GABLS4-Stage3-10 h) assumes a dry atmosphere, with no radiation. The surface pressure is held constant to 651 hPa (Dome C is at 3,223 m above sea level), and the surface temperature is prescribed and evolves with time, following the observations made at Dome C. For the computation of the surface wind stress, the surface roughness length is set to $z_0 = 10^{-3} \text{ m}$ for momentum and $z_{0h} = 10^{-4} \text{ m}$ for heat, following Vignon, van de Wiel et al. (2017b) and Couvreux, Bazile et al. (2020a).

3.2. GABLS4 Large-Eddy Simulations

Couvreux, Bazile et al. (2020a) compare seven LES models over GABLS4-Stage3-10 h. Two of them are not considered here because of a slightly different setup compared to the other five (slightly coarser resolution or different roughness lengths). The five remaining LES models use an isotropic resolution of 1 m over a $500 \times 500 \times 150 \text{ m}^3$ domain. The LES models compute their surface fluxes (momentum and heat) from the prescribed surface temperature and roughness lengths using their own parameterization. In such a setup, especially thanks to the high resolution, the spread among the LES ensemble is rather small, except very close to the surface. In particular, there are substantial differences for the surface sensible heat flux (e.g., -6 to -13 W m^{-2} at 2300 LT) or for the friction velocity (0.07 – 0.11 m s^{-1} — see Figures 1a and 1d). This spread is however much smaller than the spread among the full Stage 3 LES ensemble (Couvreux, Bazile et al., 2020a).

Although the setup is idealized compared to the observed situation, observations have been used qualitatively to evaluate the simulation behavior. The heights of the stable boundary layer and of the low-level jet are overestimated compared to observations but this might be due to the neglect of subsidence in the LESs. Otherwise, the jet intensity and the stratification are consistent with observations. The wind turning at 41 m is also much more realistic in the high-resolution LESs than in the coarse-resolution LESs (Couvreux, Bazile et al., 2020a).

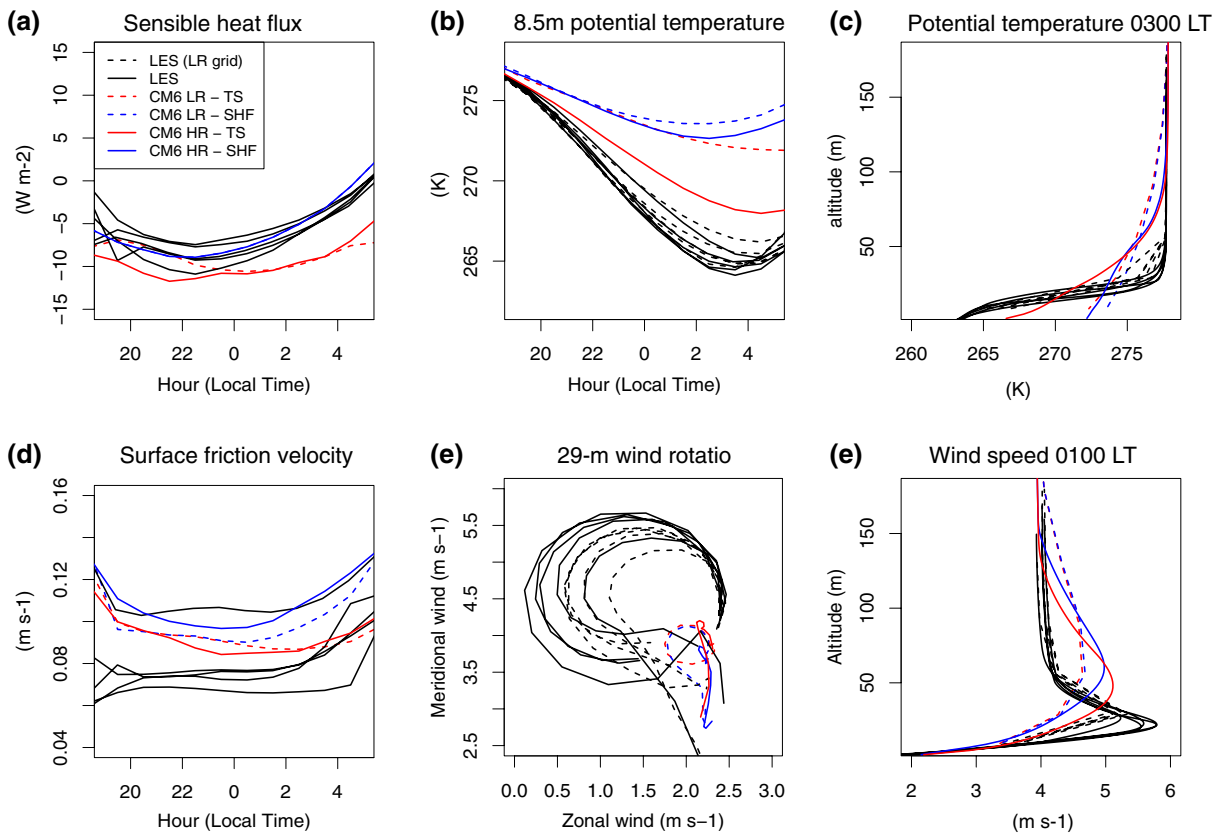


Figure 1. Time evolution of (a) sensible heat flux (W m^{-2}), (b) 8.5-m potential temperature (K), (d) surface friction velocity (m s^{-1}) and (e) 29-m wind rotation, and vertical profile of (c) potential temperature at 0300 LT (K) and (f) wind speed at 0100LT (m s^{-1}), for the LESs (solid black lines), the LESs interpolated on CM6-LR vertical levels (dashed black lines), and CM6-LR (dashed lines) and CM6-HR (solid lines) simulations. CM6 simulations are either forced by the surface temperature (red lines) or by the surface sensible heat flux (blue lines). LES, large-eddy simulation.

3.3. SCM Configurations

The ARPEGE-Climat single column model is run for two different setups. The first one follows exactly the GABLS4-Stage3-10 h setup detailed in Section 3.1. The second one, derived from GABLS4-Stage3-10 h, and referred to as GABLS4-Stage3-10hr-shf, prescribes the surface sensible heat flux, instead of the surface temperature, and thus further removes the coupling between the surface and the atmosphere. In this latter setup, the prescribed surface sensible heat flux corresponds to the ensemble mean of the five LESs.

In both setups, all model parameterizations are deactivated, except those for the atmosphere turbulence and the surface fluxes. Note that, in the GABLS4-Stage3-10 h setup, the surface flux parameterization follows the work of Mascart et al. (1995, see Section 2.2), while in the GABLS4-Stage3-10hr-shf, it is replaced by the simplified version described in Paulson (1970, see also Section 2.2).

We also explore the turbulence parameterization behavior for two different vertical resolutions, namely the standard vertical resolution of ARPEGE-Climat 6.3 (91 vertical levels, with about 15 levels below 1,500 m, indicated by LR when relevant), and a constant high-vertical resolution of 2 m up to an altitude of about 400 m and then decreasing as the standard vertical grid up to the model top (indicated by HR when relevant).

To conclude, the present work addresses three SCM configurations of ARPEGE-Climat 6.3: SCM-HR-SHF in which the SCM has a high vertical resolution and is forced by surface fluxes, SCM-LR-SHF similar to SCM-HR-SHF but with the model standard vertical grid, and SCM-LR-TS similar to SCM-LR-SHF but forced by the surface temperature. The simulations corresponding to the standard calibration of the turbulence parameterization (Table 1) are consistently referred to as CM6-HR-SHF, CM6-LR-SHF, and CM6-LR-TS.

4. Statistical Framework: History Matching with Iterative Refocusing

The calibration of the free parameters of a parameterization is a difficult task due to many degrees of freedom (i.e. the number of free parameters) and the computational cost of operational configuration simulations. For example, the turbulence parameterization used here has eight different free parameters and seven are kept for the calibration experiments. If one wanted to systematically explore this seven-dimensional space, let say with 10 values in each parameter range, this would require 10^7 simulations. This is clearly prohibitive for most model configurations, even in an SCM framework. Therefore, surrogate models, or emulators, are required to appropriately explore the full model parameter space, with a virtually negligible computational cost. In essence, these surrogate models are mathematical functions that approximate, or emulate, SCM outputs, and which are trained on a few tens to hundreds of well-chosen simulations.

In model calibration, one may want to seek for the best or optimal set of model parameters. As discussed in Hourdin et al. (2017) and in Williamson et al. (2015), this may lead to overtuning issues. Therefore, as advocated in Couvreur, Hourdin, et al. (2020), we adopt a different point of view for model calibration. Namely, we use history matching with iterative refocusing as proposed in Williamson et al. (2013). The main objective of the statistical approach is thus not to find the optimal set of parameters, but rather to remove regions of the parameter space in which the model behavior is inappropriate for a given set of performance metrics. The approach thus documents the part of the parameter space which is compatible with the relevant references, accounting for the several sources of uncertainty. The algorithm is iterative, in the sense that several consecutive waves are considered. Each wave starts with a small number of new simulations performed with the full model. These new simulations aim at improving the surrogate model accuracy, but only where needed (i.e., within the acceptable range defined from the previous wave). The search for acceptable parameter values is thus refocused in a reduced space during the following wave. The iterations (waves) stop when convergence is achieved for the space of acceptable parameters.

In the following, history matching with iterative refocusing (or history matching hereafter for the sake of brevity) is applied in a SCM/LES comparison framework as detailed in Couvreur, Hourdin et al. (2020b). Its different steps are briefly presented below.

1. *Selection of the targeted metrics*: These metrics aim at summarizing the model behavior. Reference metric values are computed, here using LES results. In the present framework, a metric is a diagnostic that is used to quantify how a simulation (or its emulated version) is close to the reference. It is simply a scalar function of the model free parameters, such as the temperature or the wind speed at a given altitude and at a given time. It may be more sophisticated and include time or vertical integration of a given model variable, or be a function of several model variables. They generally do not include reference information (e.g., such as scores) as that is included in the implausibility (see hereafter).
2. *Identification of the free parameters*: Free parameters of the model parameterizations are selected and their possible ranges (generally determined from the modeler expertise) defines the input parameter space. In our case, only parameters from the turbulence scheme are selected (see Section 2.1). Their ranges are given in Table 1
3. *Experimental design*: The first wave experimental design (namely the sets of parameters for which a simulation is actually performed) is built by sampling the parameter space previously defined. An optimal sampling of the input parameter space is ensured using a Latin hypercube method (Williamson, 2015).
4. *Building of the surrogate models*: Metrics defined in step 1 are computed from the first wave simulations (performed in step 3) and used as a training sample to build a surrogate model for each metric. Each surrogate model is a Gaussian-process-based statistical model, a tool well-known in the Uncertainty Quantification community, and which has the advantage to predict both the metric and its uncertainty (e.g., Salter & Williamson, 2016). A prior mean function is imposed as a linear combination of simple functions of the input parameters (monomials, Fourier functions and interaction terms chosen according to a predictor selection method—see Williamson et al., 2013). The Gaussian process (GP) assumed a stationary squared exponential kernel, fitted using a Hamiltonian Monte Carlo method (Carpenter et al., 2017). Further details can be found in Couvreur, Hourdin et al. (2020b) and references therein.
5. *History matching*: The parameter space is then systematically explored using the surrogate models built during the previous step, and the emulated metrics are compared to their respective reference. The space of acceptable parameter values is finally determined by ruling out the parts of the input parameter space

which lead to metric values too far from the reference value, accounting for the uncertainty on the references, the SCM and the surrogate models. For a given metric f , the following measure $I_f(\lambda)$, referred to as implausibility, is thus introduced to identify which part of the parameter space has to be ruled out

$$I_f(\lambda) = \frac{|r_f - E[e(\lambda)]|}{\sqrt{\sigma_{r,f}^2 + \sigma_{d,f}^2 + \text{Var}[e(\lambda)]}} \quad (13)$$

where λ is a point of the parameter space and $e(\lambda)$ the metric value predicted by the surrogate model. More precisely, $E[e(\lambda)]$ is the expectation value of the metric Gaussian process, and $\text{Var}[e(\lambda)]$ is its variance. The latter is a measure of the surrogate model uncertainty for the set of parameters λ . r_f is the reference metric value, and, in the present work, the LES ensemble mean of the metric f . $\sigma_{r,f}^2$ is the reference uncertainty and is computed as the LES ensemble variance of the metric f . $\sigma_{d,f}^2$ is the SCM discrepancy or structural error for this metric. This last value is a priori unknown and its estimation could be challenging. Williamson et al. (2017) suggests a “tolerance-to-error” approach for model discrepancy. The implausibility $I_f(\lambda)$ thus measures the distance between the reference and the predicted metric value, normalized by the sum of uncertainties (supposed to be independent). The implausibility can thus be small either because the predicted value is close to the reference, or because the uncertainties are large. From the implausibility, and thus after this first iteration through steps 1–5, the Not-Ruled-Out-Yet (NROY) space for the metric f is defined as:

$$\text{NROY}_f^1 = \{\lambda \mid I_f(\lambda) < T\} \quad (14)$$

where T is a given threshold taken here as 3, following the 3- σ rule of Pukelsheim (1994), which states that at least 95% of any unimodal distribution lies within its 3- σ range centered around its mean. This threshold can be reduced once the surrogate model for the metric f is sufficiently accurate. In the case of multiple metrics, the full NROY¹ space is simply defined as the intersection of the NROY spaces computed for each metric:

$$\text{NROY}^1 = \bigcap_f \text{NROY}_f^1 = \{\lambda \mid I_f(\lambda) < T, \text{ for all } f\} \quad (15)$$

Note that such an option in combining metrics may lead to multiple testing problems (e.g., Vernon et al., 2010; Couvreur, Hourdin et al., 2020b), but the small number of metrics used in this work reduces this risk.

As mentioned above, the implausibility $I_f(\lambda)$ for a set of parameters λ can be small either because the SCM emulated metric is close to its reference, or because the emulator uncertainty at λ is high. Note that $\sigma_{r,f}$ and $\sigma_{d,f}$ are input data to the whole history matching strategy, independent of the surrogate models. The iterative refocusing thus seeks to reduce this emulator uncertainty and thus provide a better estimate of the true space of acceptable sets of parameters. As a result, several iterations (or waves) of the steps 3–5 are performed, with an adapted experimental design in step 3. Namely, we only need to improve the emulator accuracy within the current NROY space, thus we only resample NROY¹ after wave 1, or NROY ^{N} after wave N . Once the surrogate model uncertainty is sufficiently reduced so that the other uncertainties dominate the implausibility, the NROY space is not further shrunk by new waves. The iterative process has thus converged and it is not necessary to perform additional waves. The convergence question is discussed in section 6.2.

5. Ability of ARPEGE-Climat to Simulate the GABLS4 Stable Boundary Layer

The tool introduced in the previous section is used to assess whether the ARPEGE-Climat turbulence scheme is able to capture the main properties of the GABLS4 stable boundary layer. In particular, we determine which part of the space of model free parameter values is compatible with the LES references. The vertical resolution is potentially a critical aspect for the scheme and therefore we start with the SCM-HR high-resolution configuration of ARPEGE-Climat. We also remove a degree of freedom in the surface-atmosphere coupling by prescribing the surface sensible heat flux instead of the surface temperature, to focus

only on the turbulence scheme (SCM-HR-SHF configuration). Then, Section 5.2 discusses results with the standard vertical resolution of ARPEGE-Climat and with two different surface boundary conditions (SCM-LR-SHF and SCM-LR-TS configurations).

5.1. High-Resolution SCM Configuration Forced by Surface Sensible Heat Flux (SCM-HR-SHF)

5.1.1. ARPEGE-Climat Standard Calibration

As a starting point of the present work, we evaluate the standard calibration of the ARPEGE-Climat turbulence scheme (parameter values indicated in Table 1), for a high vertical resolution, and when forced by the LES ensemble mean surface sensible heat flux (CM6-HR-SHF simulation). Figure 1b presents the time evolution of the 8.5-m potential temperature from 1800 to 0700 LT. 8.5 m is approximately the altitude of the first level in the standard vertical resolution model (between the third and fourth level in the present high resolution configuration). Linear interpolation is used in the high-resolution model configuration and in LESs to compute the 8.5-m potential temperature. The LES models (solid gray lines) simulate a significant cooling until 0200 LT, from about 277 K to about 264 K. The minimum potential temperature is reached around 0300 LT. At that time, the potential temperature vertical gradient between 15 and 29 m varies between 0.4 and 0.8 K m⁻¹ among the LESs (Figure 1c). At 0100 LT, a low-level jet is well formed in all the LESs (Figure 1f). The altitude of its peak is similar in all LESs, around 22 m, and its intensity ranges between 5 and 6 m s⁻¹. The inertial rotation of the wind at 29 m is further emphasized in Figure 1e. It is consistent with the theory (e.g., Blackadar, 1957) and representative of the observations collected at Dome C (Gallée et al., 2015).

CM6-HR-SHF severely underestimates the 8.5-m potential temperature cooling during the first half of the night (Figure 1b, solid blue line). The minimum potential temperature reaches about 273 K at 0200 LT, about 8 K warmer than the LES corresponding value. The potential temperature vertical profile at 0300 LT (Figure 1c) emphasizes that CM6-HR-SHF simulates a boundary layer, which is too thick and which stability is underestimated: the potential temperature vertical gradient is at least six times weaker (0.06 K m⁻¹) than in the LESs. Consistently, the CM6-HR-SHF low-level jet is too high, located near 55 m (Figure 1f). Its intensity is significantly weaker than in all LESs but one. The wind rotation is also strongly underestimated at 29 m (Figure 1e).

In order to assess the model sensitivity to its internal turbulent parameters, we choose to synthesize the model behavior with four scalar metrics (step 1 in Section 4). The sensitivity to the choice and number of metrics is discussed in Section 6.1. The nocturnal cooling and boundary layer stability are quantified using the potential temperatures at 2 m and 8 m (referred to as θ_{2m} and θ_{8m} , respectively); these two CM6-HR vertical levels allow to constrain the θ vertical gradient. These two metrics are computed at 0300 LT, when θ is minimum in the LESs. The low-level jet structure is measured using the maximum of the supergeostrophic wind speed and the wind speed at 55 m (referred to as w_{max} and w_{55m} , respectively). The latter altitude corresponds to the level where the wind returns to its geostrophic value in the LESs (it is also the altitude of the CM6-LR third level). These two last metrics are taken at 0100 LT when the low-level jet is well established.

5.1.2. Defining the Acceptable Range of the Turbulence Free Parameters

Based on the seven parameters identified in Section 2.1 and their proposed ranges (step 2 of the statistical framework described in Section 4), the history matching Wave 1 consists of 70 simulations performed with the SCM-HR-SHF configuration. The associated 70 sets of parameters sample the input parameter space following the experimental design proposed in step 3, namely a Latin hypercube. The results for these 70 simulations are shown by the orange lines in Figure 2. Although the majority of these simulations exhibits a too weak cooling, some of them capture the LES behavior, with both a correct θ vertical profile at 0300 LT and a correct overnight evolution of θ at 8.5 m. Concerning the wind, in most simulations, the low-level jet at 0100 LT is too high, so that the return to the geostrophic wind occurs above 100 m. The wind rotation at 29 m is poorly represented, with a too weak meridional component and a too strong zonal component. Reflecting these first conclusions, the metrics computed for each simulation are most of the time fairly far from those computed with the LESs (not shown). Nevertheless, there exist a few simulations, thus a few sets of parameters, for which the chosen metrics have values close to those computed with the LESs. This suggests that, at least for each individual metric, there exists an appropriate calibration of the ARPEGE-Climat

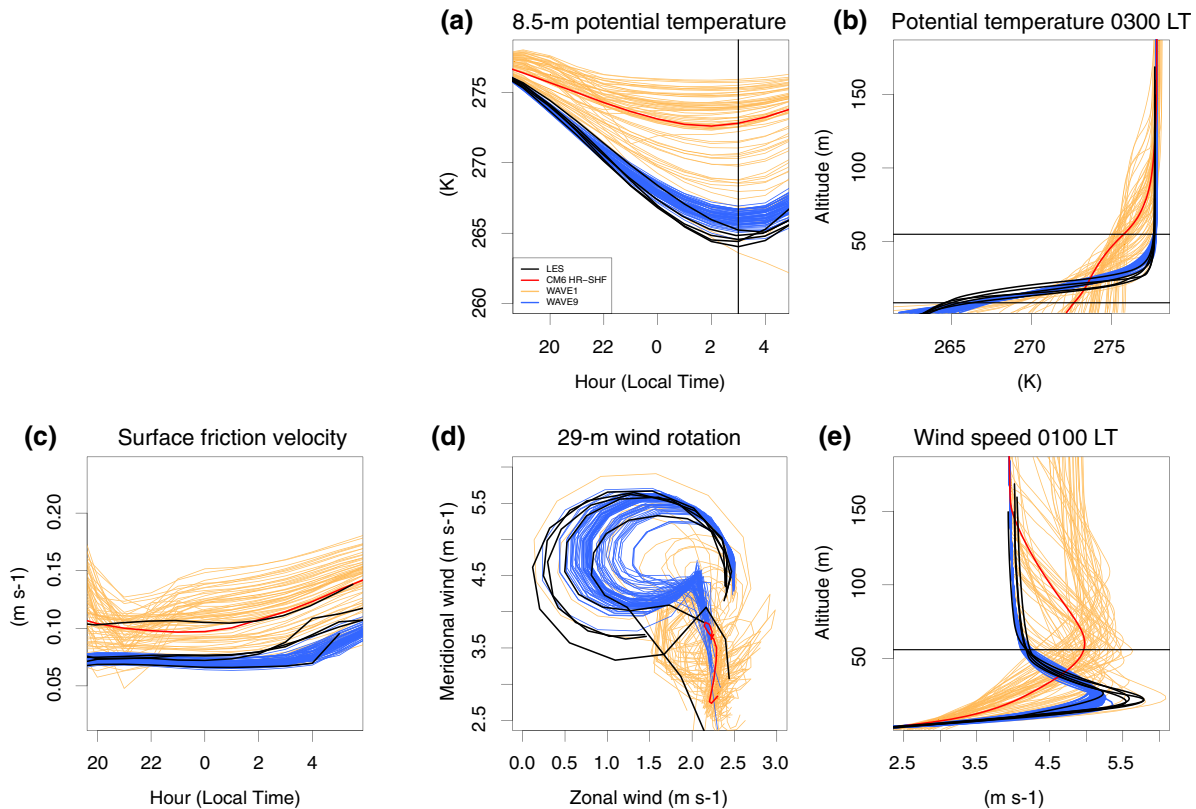


Figure 2. Same as Figure 1 but for the history matching applied to the SCM-HR-SHF configuration. The standard calibration simulation CM6-HR-SHF is indicated with the red line, the Wave 1 70 SCM simulations with the orange lines and the Wave 9 70 SCM simulations with the blue lines. Note that since it is prescribed, the sensible heat flux is not shown. LES, large-eddy simulation; SCM, single-column model.

SCM-HR-SHF configuration. Based on this result, and for the sake of simplicity, in the following, we choose to explore the model performance considering no structural (or tolerance to) error (i.e. $\sigma_{df} = 0$ in Equation 13). Such a choice may lead to overtuning (Williamson et al., 2017) and will need to be reconsidered in the context of the full model calibration. It will also be discussed in Section 6.2.

The metric values, computed for each simulation of this Wave 1, are used as a training sample to build a surrogate model for each of the four metrics (step 4—see Section 4 and references therein for their building). These surrogate models then allow us to explore the parameter space more exhaustively by estimating the value of each metric at as many new points as desired. In practice, the complete parameter space is resampled using a new Latin hypercube of $\mathcal{O}(10^6)$ points and the surrogate models provide estimate (and uncertainty) of the four metrics for all these new points. As explained in Section 4, the following step consists of history matching (step 5): for each point sampled in the parameter space, the implausibility with respect to each metric is calculated, and the maximum implausibility over the four metrics (i.e., the most discriminating metric) is used to characterize NROY¹. Sampled points with an implausibility greater than a threshold set to three are ruled out. After this first wave, the remaining space is 3.0% of the initial space, so a large part of the input parameter space is rejected. The NROY space obtained at this stage (not shown) shows that **AE**, **KOZMIN** and **ZMAX** have a negligible influence on the results after this first iteration. The model behavior as a function of the other four parameters indicates a preference for values that significantly reduce the turbulent mixing. Table 2 details for each metric the performance of the 70 simulations performed for this wave 1 and for each of the following waves. For the first wave, all the metrics are very discriminating (more than two thirds of the simulations are incompatible with the LES reference values) except the one representing the jet intensity.

We then iterate several times over steps 3–5 described in Section 4. 70 points are randomly sampled in NROY¹ and the corresponding simulations are carried out. These new simulations feed the wave 2. Steps

Table 2
Evolution of the Different Metrics (Columns) Over the Successive Waves (Rows) for the History Matching Applied to the SCM-HR-SHF Configuration.

Wave	θ_{2m}		θ_{8m}		W_{max}		W_{55m}	
	Spread	$I > 3$	Spread	$I > 3$	Spread	$I > 3$	Spread	$I > 3$
1	260.3–277.9	58	266.3–276.5	64	4.8–5.7	2	4.1–5.3	46
2	252–269.4	10	263.2–270.7	3	4.7–5.8	2	4.1–5.6	31
3	255.2–265.2	10	262.3–268.7	5	5.0–5.9	0	4.1–4.5	6
4	259.1–266.1	7	264.4–267.7	4	5.0–5.5	0	4.1–4.3	0
5	258.9–266.5	10	264.4–267.3	0	5.0–5.4	0	4.1–4.5	1
6	261–265.1	4	264.7–267.4	0	5.0–5.4	0	4.1–4.3	2
7	261.5–264.9	3	264.9–267.4	2	5.0–5.5	0	4.1–4.3	2
8	261.7–264.8	4	265.4–266.9	0	5.0–5.4	0	4.1–4.3	0
9	261.7–265	2	265.7–267	0	5.1–5.3	0	4.1–4.3	0
LESs	263.4 ± 0.43	–	265.0 ± 1.91	–	5.6 ± 0.68	–	4.2 ± 0.06	–

For each metric, the first column indicates the spread (minimum–maximum) of the metric value computed for the 70 simulations performed at the beginning of Wave $N = 1, \dots, 9$. For the LESs in the last row, the LES ensemble mean $\pm 3\sigma_{rf}$ is indicated. The second column (Implausibility $I > 3$) gives the number of these simulations that do not belong to the NROY space estimated at the end of Wave N (namely $NROY^N$). LES, large-eddy simulations; NROY, not-ruled-out-yet; SCM, single-column model.

4 and 5 then provide updated surrogate models in $NROY^1$ and a new estimate of the NROY space, namely $NROY^2$. A third wave can then start. The process is repeated until it is no longer possible to reduce the NROY space (i.e., $NROY^N \sim NROY^{N+1}$ if N is the number of a wave). Table 2 indicates no clear further reduction of the NROY space from Wave 6, that is the 70 simulations performed at the beginning of Wave 6 are for most of them still in the NROY space estimated at the end of Wave 6 ($NROY^6$). This thus suggests convergence of the results and that we can stop the iterations. From Wave 4, w_{max} , w_{55m} and θ_{8m} metrics are no longer discriminating, that is, for these metrics, almost no more part of the parameter space is ruled out. For θ_{2m} , the process evolves more slowly. At Wave 5, 10 simulations are still ruled out because of this metric. From Wave 6 onwards, the number of rejected simulations varies between 2 and 4. The difference between two consecutive waves is then only due to sampling effects of the remaining space. Note that at a given wave N , the current tool version has the limitation to not use the simulations performed during the previous waves $k < N$ to build the wave N emulators. We stopped after 9 waves.

Figure 2 compares the vertical profiles and time series of temperature and wind between the simulations of Wave 1 (orange lines), those of Wave 9 (blue lines), and the reference LESs (black lines). The 70 simulations performed in Wave 9 belongs to $NROY^8$. Given the chosen metrics and the LES uncertainty, it is not possible to discriminate them (except for 2 of them as indicated in Table 2). Figure 2b shows that the θ vertical profile at 0300 LT (time used to compute θ -related metrics) is very close to that of the LESs. However, it seems that there is still room for improvement close to the surface but performing an additional iteration does not reduce the near-surface spread of the SCM simulations. Pushing the calibration exercise further would require more simulations but the results obtained are already qualitatively satisfying. Looking at the time series of θ at 8.5 m (Figure 2a), there is also a very clear improvement in the results all along the simulation although we only use the metric at one given time (0300 LT). The results for the wind are good also. The Wave 9 simulations capture well the wind vertical profile at 0100 LT and the 29-m wind veering, falling within the LES spread (Figures 2d and 2e). Note that the four instantaneous constraints put on the SCM simulations are sufficient to achieve model calibrations that perform well over the entire GABLS4 stable boundary layer, thus illustrating the consistency of the ARPEGE-Climat turbulence scheme physics.

Figure 3 provides a visualization of the seven-dimensional NROY space, obtained after 9 waves, namely $NROY^9$. This half matrix tentatively synthesizes the space containing the optimal sets of parameters. Each panel of the half-matrix is a “projection” of $NROY^9$ in the two-dimensional space defined by solely two

parameters, the names of which are given in the diagonal. For instance, the X -axis of the upper-right panel corresponds to **ZMAX**, varying between its minimum and maximum values provided in Table 1 and indicated in the upper part of its diagonal box (30 and 200 m, respectively) and its Y -axis corresponds to **CM**, varying between 0.05 and 0.3 (see the right part of its diagonal box). To numerically characterize $NROY^9$ in this two-dimensional parameter space, each panel consists of 15×15 pixels, which values quantify the fraction of the full input space in the other 5 directions that belongs to $NROY^9$. This value is estimated thanks to the implausibilities computed over 10^6 sets of parameters which sample the input parameter space (Latin hypercube). The $NROY^9$ fraction of this input space (also equal to the sum of the 15×15 pixel values of each panel) is indicated as text in the upper-right part of the figure (0.0959% in the present case). Gray color of a pixel indicates that the fraction is zero given the input space sampling, meaning that there is mostly no way to fit the reference by varying the other 5 parameters when the 2 panel parameters are fixed at the pixel corresponding values. Dark blue to yellow colors indicate that the combination of the 2 panel parameters is the most likely to provide satisfying results.

As shown in Figure 3, the history matching method selects parameter values that reduce turbulent mixing. The exchange coefficients are significantly reduced (mainly through the high probability of low **CM** and **LMIN** values, see the panels on the first and third rows and on the third column of the half matrix). The calibration also leads to small values of **CE** (panels on the second row and the second column of the half matrix), which controls the dissipation length scale: the TKE dissipation is increased and thus the TKE and the turbulent mixing are reduced. These results are consistent with previous works that generally indicate a too strong turbulent mixing in numerical weather and climate models (e.g., Beljaars & Viterbo, 1998; Sandu et al., 2013). The case of **LMIN** is also consistent with the work of Vignon, Hourdin et al. (2017) which showed the importance of removing most of the bounds in the turbulence parameterization used in global models to better represent stable boundary layers. The whole framework highlights the high sensitivity of the model results to this parameter (Figure 3), and illustrates the difficulty to calibrate the model for stable boundary layers with **LMIN** values greater than about 4–5 m. The influence of **AT**, which controls the turbulent flux of θ , is significant but not as decisive as that of the three previous parameters. The value currently used in ARPEGE-Climat (black squares on Figure 3) is appropriate. The results are not sensitive to **AE**, so that there is a priori no need to change its current value. This parameter influences the vertical diffusion of TKE. It therefore appears that this term is negligible compared to the other terms in the TKE equation (Equation 3) in the case of the GABLS4 very stable boundary layer (not shown). Finally, **KOZMIN** and **ZMAX** which directly limit the turbulent fluxes, are not relevant to adjust this high-resolution SCM configuration. This is expected given the bound formulation (Equation 6) which tends to zero as the vertical grid spacing goes to zero. These parameters are more critical for the model standard resolution (see Section 5.2).

As a preliminary conclusion, the history matching approach demonstrates that the ARPEGE-Climat turbulence scheme contains the required physics to represent the GABLS4 strongly-stable boundary layer, at least for vertical resolution of $\mathcal{O}(1\text{m})$. Besides, if the ARPEGE-Climat standard calibration is not appropriate (i.e. the free parameter standard values are not retained in the $NROY$ space), the standard values of **CM** and **CE** are reasonable (i.e. very close or within the $NROY$ space). It is rather the lower bound **LMIN** on the mixing length that is the most discriminating, a conclusion similar to the one obtained by Vignon, Hourdin et al. (2017a).

5.2. Standard-Resolution SCM Configuration Forced by Surface Sensible Heat Flux (SCM-LR-SHF)

The behavior of the standard-resolution standard-calibration SCM configuration, forced by the LES ensemble-mean surface sensible heat flux (CM6-LR-SHF), is summarized in Figure 1 (dashed blue line). In order to compare the low-resolution SCM results with those of the LESSs, the latter are regridded onto the vertical grid of CM6-LR. The CM6-LR-SHF overnight cooling is weak, and slightly weaker than in CM6-HR-SHF, thereby indicating sensitivity of the model results to vertical resolution. The minimum potential temperature is reached earlier than in the LESSs (as for CM6-HR-SHF) and is 8 K warmer. The low-level jet is weakly marked. The altitude of the maximum wind speed is too high (55 m) and the wind speed remains too strong above 55 m, where it should be geostrophic. The wind rotation is slightly better represented than in CM6-HR-SHF but still too weak. The CM6-LR-SHF behavior is thus broadly similar to its high-resolution

SCM-HR-SHF NROY - Wave 9

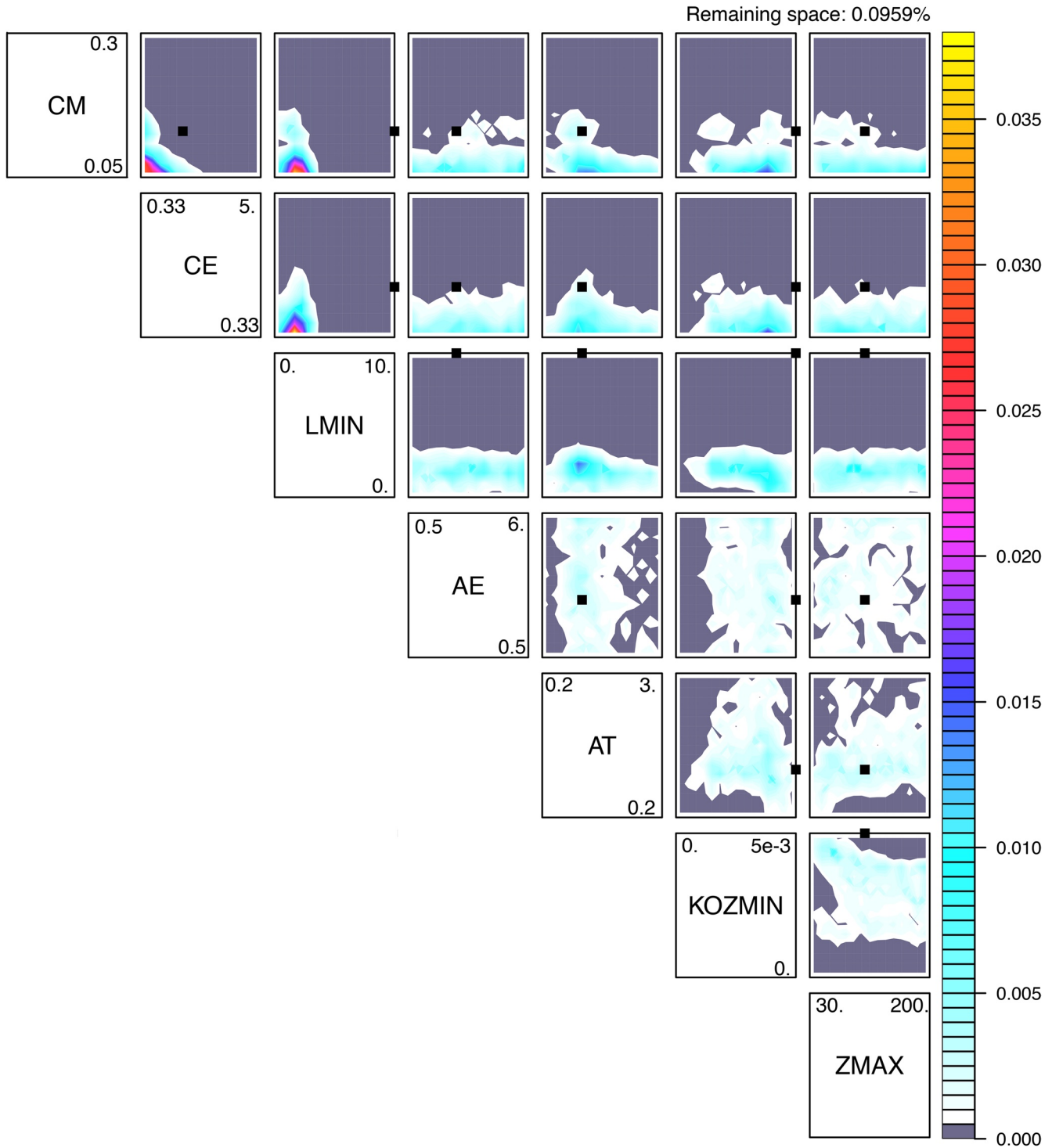


Figure 3. NROY half matrix for Wave 9 of the SCM-HR-SHF history matching (NROY⁹). Each panel displays the fraction of the input space with implausibility lower than the chosen threshold (i.e. NROY⁹, see Section 4) as “projected” in a two-dimensional parameter space, the name of the 2 parameters being given in the diagonal of the half matrix. In the diagonal, each parameter box provides the minimum and maximum values to be used for the X-axis of the panels above the box (numbers in the upper part of the box) and the Y-axis of the panels to the right of the box (numbers to the right in the box). The black square on each panel locates the parameter values of the ARPEGE-Climat CM6-HR standard calibration. See text for further details. NROY, not-ruled-out-yet; SCM, single-column model.

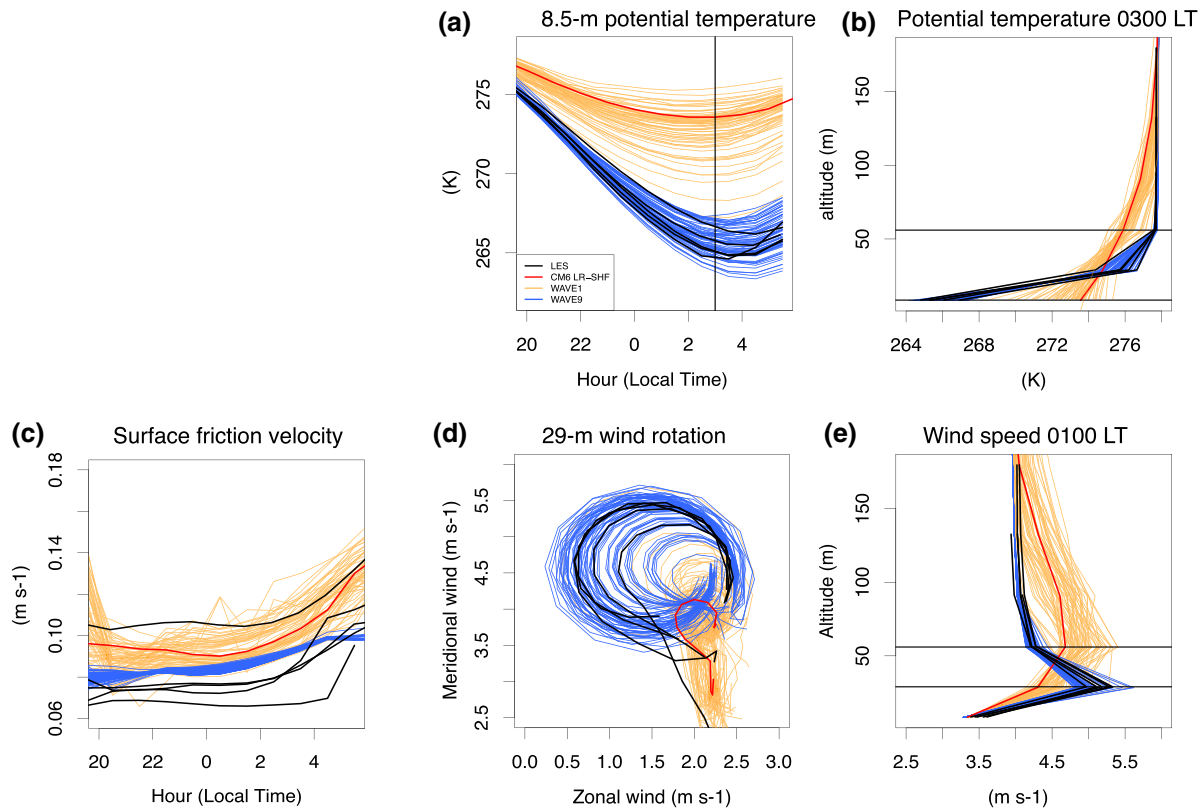


Figure 4. Same as Figure 2 but for the history matching applied to the SCM-LR-SHF configuration. The standard calibration simulation is CM6-LR-SHF (red line). LES, large-eddy simulation; SCM, single-column model.

counterpart, and consistent with an overestimated turbulent mixing. We now investigate whether this behavior is intrinsic to the parameterization (for this standard resolution), or results from a poor calibration for stable boundary layers. Similar scalar metrics to those used in Section 5.1 for SCM-HR-SHF are chosen to constrain the θ vertical gradient and the low-level jet structure. The associated altitudes are slightly adapted to be consistent with the SCM standard vertical resolution. The θ gradient is characterized by θ at the first and third model levels at 0300 LT, namely 8.5 m ($\theta_{8.5m}$) and 55 m (θ_{55m}). Because of the rather large uncertainty of θ at the second SCM level (29 m) in the LESs (about 5 K), choosing this metric would have been less efficient in reducing the free parameter space. As in the high-resolution experiment, the jet structure is summarized by the wind speed at 29 m (second SCM-LR level and altitude of the wind maximum— w_{29m}) and at 55 m (third SCM-LR level— w_{55m}).

The Wave 1 simulations present a large spread, with only a few simulations getting close to the LES results (Figure 4, orange lines). Note here that before introducing the parameters **KOZMIN** and **ZMAX** in the history matching process, this variety of behaviors was not observed and none of the simulations performed well (not shown). Indeed, with the default setting, this minimum bound for the mixing coefficients was systematically reached and any modifications of the other parameters had very little effect as hidden by the minimum bound. The consideration of **KOZMIN** and **ZMAX** in the history matching was crucial in the standard resolution configuration. Table 3 shows the evolution of the metrics for nine waves. From Wave 3, the results are already satisfying for most of the metrics. Only θ_{55m} needs a few more waves to achieve reasonable values. The Wave 9 NROY space, namely $NROY^9$, is finally slightly less than 0.1% of the input parameter space (Figure 5).

The θ vertical profiles at 0300 LT of all the Wave 9 simulations are very close to the LESs (Figure 4b, blue lines). The time evolution of the 8.5-m potential temperature still presents some dispersion, which is consistent with the LES uncertainty. Wave 9 simulations also capture well the sharpened wind speed vertical

Table 3
Same as Table 2, but for the History Matching Applied to the SCM-LR-SHF configuration.

WAVE	$\theta_{8.5m}$		θ_{55m}		W_{29m}		W_{55m}	
	Spread	$I > 3$	Spread	$I > 3$	Spread	$I > 3$	Spread	$I > 3$
1	267.9–275.7	67	275.3–277.2	65	4.6–5.5	4	4.4–5.4	59
2	263.4–271.5	8	276.5–277.9	24	4.7–6.0	1	4.1–4.7	9
3	262.9–269.0	1	277.4–277.8	2	4.8–5.3	0	4.1–4.6	3
4	263.2–267.8	0	277.4–277.8	4	4.8–5.4	0	4.1–4.6	1
5	263.2–268.8	1	277.4–277.8	7	4.8–5.5	0	4.1–4.3	0
6	263.9–267.5	0	277.5–277.8	2	4.7–5.8	1	4.1–4.3	0
7	264.4–268.4	0	277.5–277.8	0	4.8–5.7	1	4.1–4.4	1
8	263.6–268.0	0	277.5–277.8	4	4.8–5.5	0	4.1–4.3	0
9	264.2–267.3	0	277.6–277.7	0	4.8–5.6	0	4.1–4.4	3
LESs	265.6 ± 2.3	–	277.6 ± 0.32	–	5.2 ± 0.39	–	4.3 ± 0.19	–

Note that the LES metric values might be slightly different from those in Table 2, as they are computed here using LES data interpolated onto the SCM-LR vertical grid. LES, large-eddy simulations; SCM, single-column model.

structure at 0100 LT, and much better than CM6-LR (red line). There is still some significant spread in the jet intensity, but again, it reflects the LES discrepancies (Figure 4e). All the simulations reproduce well the 29-m wind rotation (Figure 4d).

The half-matrix of Figure 5 presents the remaining space after nine waves, that is, NROY⁹. In contrast with the SCM-HR version, the **KOZMIN** parameter appears critical. The history matching thus shows that the SCM can only behave well over the GABLS4 case for very low values of **KOZMIN**. Given that this parameter is used to maintain significant turbulent fluxes within the first levels of the model, it was rather expected. For the other parameters, the results are similar to those obtained with the SCM-HR-SHF configuration: parameter values that allow weak turbulent mixing are found the most suitable for adjusting the turbulence parameterization. Note that the CM6 values (black square) of **KOZMIN** and **LMIN** are well beyond their acceptable range (already eliminated at Wave 1). As for the high-resolution configuration, the SCM behavior is also significantly sensitive to **CE** and **CM** with acceptable values leading to decrease of the turbulent mixing but the CM6 default values are not ruled out. **AE** and **AT** have no significant impact. To summarize, the ARPEGE-Climat turbulence parameterization is able to capture the GABLS4 stable boundary layer as described by the reference LESs, even with the standard-resolution SCM-LR configuration. Such an acceptable behavior requires to almost remove the two bounds on the turbulent mixing induced by **LMIN** and **KOZMIN**, while the other free parameters of the parameterization can be kept close to their original values, similar to those proposed in Cuxart et al. (2000) and Cheng et al. (2002). The existence of the **LMIN** and **KOZMIN** in the model is thus strongly questioned by the present results. At least only small values of these parameters are compatible with the representation of strongly-stable boundary layers, which suggests that the associated formulations need to be revised.

The history matching approach was also applied to a configuration in which the surface boundary condition is provided by the surface temperature (SCM-LR-TS configuration), similarly to the LES setup, but also similarly to a more operational use of the model. In this case, the surface fluxes for momentum and heat follow the formulation described in Section 2.2.2. The results are mostly similar to those obtained with the SCM-LR-SHF configuration used here above and therefore are only reported in Appendix A.

6. Discussion

In this section, we discuss in more detail two features of the calibration statistical approach that are rather subjective. The selection of metrics is first emphasized as a key step that deserves some caution. We discuss the selection we made in Section 2 and analyze the respective role of each selected metrics. Second,

SCM-LR-SHF NROY - Wave 9

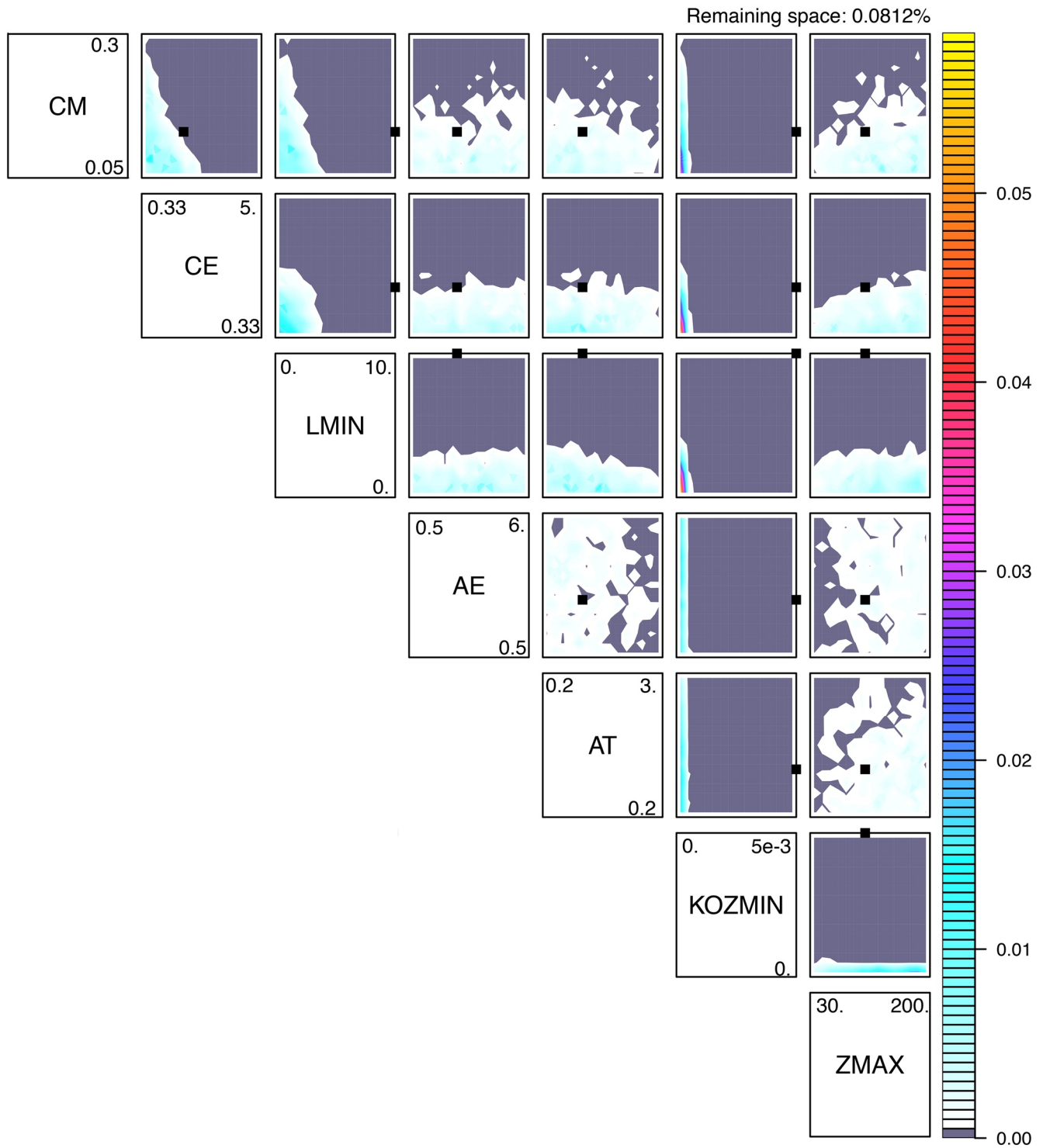


Figure 5. Same as Figure 3, but for the history matching applied to the SCM-LR-SHF configuration (NROY space obtained at Wave 9). NROY, not-ruled-out-yet; SCM, single column model.

Table 4
Results of the History Matching Applied to the SCM-LR-SHF Configuration, Considering Only One Metric at a Time.

Experiment	$\theta_{8.5m}$	θ_{55m}	W_{29m}	W_{55m}
CM6-LR-SHF	273.5	275.5	4.7	4.7
LESS	265.7	277.7	5.2	4.2
LES std	0.83	0.027	0.14	0.035
NROY ¹	2.6 %	2.8 %	25.4 %	6.0 %
NROY ²	1.9 %	0.7 %	22.4 %	0.5 %
NROY ³	1.8 %	0.6 %	21.2 %	0.4 %
NROY ⁴	1.6 %	0.5 %	20.8 %	0.3 %
NROY ⁵	1.4 %	0.5 %	19.2 %	0.3 %
NROY ⁶	1.4 %	0.4 %	19.1 %	0.3 %
NROY ⁷	1.4 %	0.4 %	18.7 %	0.3 %
NROY ⁸	1.4 %	0.4 %	18.2 %	0.3 %
NROY ⁹	1.3 %	0.4 %	18.0 %	0.3 %

The first line gives the value of each metric computed with CM6-LR-SHF, the second line gives the value of each reference metric computed as the LES ensemble mean and the third line gives the LES uncertainty computed as the LES ensemble standard deviation. The following lines give the evolution of the NROY space fraction (with respects to the input parameter space) over 9 waves. LES, large-eddy simulation; NROY, not-ruled-out-yet.

we investigate more in depth the convergence criteria, and how we have tackled it. Finally, the SCM computationally cheap approach provides the opportunity to perform a large simulation ensemble, which can serve as a basis for the evaluation of the full statistical framework.

6.1. Choice of Metrics

The choice of metrics in this type of framework is crucial, yet it is difficult to draw up a definitive list in advance. In this section, we illustrate the metric selection procedure in the case of the SCM-LR-SHF history matching and give some guidelines. It might be tempting to choose a large number of metrics to precisely control the model. But for efficiency, especially for the analysis of the results of each wave, the number of metrics should be limited. In particular, redundant metrics should be avoided. Note also that if, in the course of the framework, biases that were not accounted for emerge, new metrics can be added “on the fly” at the beginning of a wave.

The first step consists in conducting an evaluation of the model behavior in order to identify the main model biases (and especially those we care about) and then to design the appropriate metrics to quantify them. We have already seen in Section 5.2 that the two main identified biases are a weak nocturnal cooling and an insufficiently marked low-level jet structure. Analysis of the potential temperature profile at 0300 LT (Figure 4b), the time at which the minimum of θ at 8.5 m, the first CM6-LR level, is reached in the LES (Figure 4a), shows that the θ vertical gradient in the nocturnal boundary layer is too weak, with a warm bias (+8 K) at the first model level and a cold bias (about -2 K) at the third model level. The

values of θ at these two levels (respectively $\theta_{8.5m}$ and θ_{55m}) are therefore used as metrics to control the model thermodynamics. Concerning the low-level jet structure at 0100 LT (the time at which it is the sharpest in the LESS, Figure 4e), the main features to be considered are the altitude and the intensity of the wind maximum (which are respectively too high and too weak in CM6-LR), and the thickness of the jet, which can be represented by the altitude of the wind return to its geostrophic value. The first two features can be summarized by a single metric, namely the wind intensity at the second model level (w_{29m}), which corresponds to the wind maximum in the regridded LESS. The third feature can be captured by the wind intensity at the third model level (w_{55m}). Wind intensity at the fourth level could also be chosen and this choice leads to similar results (not shown).

As explained before, the framework used here takes into account the different sources of uncertainty through the notion of implausibility. Note that the lower the uncertainty on the reference metrics the more constraining the metric: particular attention was devoted to the associated reference uncertainty, and metrics with weak reference uncertainty were preferred. When it is significant, at least compared to the Wave 1 simulation errors, it may rapidly dominate the implausibility and the framework will only weakly constrain the model behavior. For the four metrics defined on the basis of model biases, the uncertainty of the LESS regarding these metrics is low in front of the biases of Wave 1 simulations (for $\theta_{8.5m}$ and w_{29m}) or almost negligible (for θ_{55m} and w_{55m}).

To understand how each of the selected metrics constrains the model, four history matching experiments are carried out, considering only one of the four metrics at a time. For each metric, the evolution of the remaining space throughout the different iterations is presented in Table 4. The w_{29m} metric is much less discriminating than the other three. 18% of the input parameter space is compatible with the references after 9 waves, as opposed to around 0.5%–1% for the other metrics. The rather large LES uncertainty for this metric explains part of this result. The importance of the reference uncertainty relative to the bias may be quantified as the ratio between the two quantities (Table 4), but consideration of this ratio is not sufficient to presuppose the importance of each metric. The results of these experiments are presented in Figure 6. The θ -related and 55-m wind speed metrics rule out around 99% of the input parameter space.

History matching with only one θ -related metrics (either $\theta_{8.5m}$ or θ_{55m}) lead to simulations with a proper nocturnal cooling (a little more pronounced when considering only θ_{55m} —compare the red shading with the gray shading in Figures 6d and 6h). They also show a low-level jet significantly better than CM6-LR with a wind maximum at the right altitude (Figures 6a and 6e). However, the return to geostrophic wind is weakly constrained and for many simulations it occurs at a too high altitude, especially for those resulting from the history matching based on $\theta_{8.5m}$ only (Figure 6a). The history matching based on w_{55m} only leads to a jet with a structure at 0100 LT very close to the LESs (Figure 6m), but with a too weak nocturnal cooling for many Wave 9 simulations (Figure 6p), even if the improvement compared to CM6-LR is still significant. These different metrics reduce the space of the parameters differently (Figure 7). If high values for **LMIN** and **KOZMIN** are systematically eliminated for each of these three metrics (Figures 7a, 7b, and 7d), they differ for the history matching results for **CM**, **CE**, and **AT**. w_{55m} (Figure 7d) strongly constrains the sum of **CM** and **CE** but has no influence on **AT**, whereas thermodynamic metrics (Figures 7a and 7b) have a less marked sensitivity to **CM** and **CE** but significantly constrain **AT** by eliminating the highest values. The weaker sensitivity of $\theta_{8.5m}$ and θ_{55m} to **CM** is explained as the turbulent heat flux depends on the product of **CM** by **AT** (see Section 2.1). The w_{29m} metric poorly constrains the nocturnal cooling and the altitude of the return to geostrophic wind (Figures 6i and 6l). The usefulness of this particular metric may be questioned. An experiment conducted without w_{29m} however slightly degrades the low-level jet structure (not shown) and leads to a slower reduction of the parameter space.

In this study, we use simple scalar metrics at given times. The time at which the metrics are calculated is chosen according to the maximum intensity of the phenomena they capture. Thus the wind-related metrics are computed at 0100 LT which is the time when the jet is the sharpest. The θ -related metrics are computed at the time of the strongest cooling (0300 LT). More complex (vector) metrics could be used, such as vertical profiles or time series (e.g., Salter et al., 2019; Williamson et al., 2017), but the results presented in Section 5 emphasize that these very simple metrics are sufficient to constrain significantly the model behavior over the entire duration of the simulation.

We would like to finally stress here that the choices made in Section 5 result from a trial-and-error empirical approach, which we advocate for. To help to analyze the results, a small number of metrics were selected from a much larger list of potential metrics. It is certainly possible to use a greater number of metrics, in practice all those that the modeler finds relevant. Nevertheless, we emphasize hereafter a few guidelines on this process of metric selection:

1. The analysis of the default configuration but also of the first wave simulations is an important step, which helps to identify model biases or appropriately simulated features and build metrics to quantify them
2. As the method takes into account the different sources of uncertainty, particular attention should be paid to the uncertainty in the LESs and to ensure that it does not dominate the biases identified above. Besides, the first wave spread might also give indications about the discriminatory ability of a given metric
3. Metrics do not need to be taken all at once starting at Wave 1 and sometimes it is preferable to start with few discriminant metrics (with priority metrics being defined by the modeler expertise) in order to ease the construction of the emulators during the following waves as over a smaller parameter space. This is ensured by the tool flexibility, as the modeler can add new metrics from a given wave
4. A preference is given to the use of simple metrics, easier to compute and interpret

6.2. Iterative Refocusing Convergence and Its Link with the Sources of Uncertainty

Deciding when to stop the iterations is not trivial. For the different experiments carried out in this study, we decided to stop the iterations when the size of the NROY space no longer decreases as we perform more waves (e.g., Table 5). Using the SCM-LR-SHF experiment, we illustrate here the NROY convergence and analyze the evolution of the different uncertainty sources that we need to take into account to monitor this convergence.

In the SCM-LR-SHF experiment, the NROY space size has converged from about the fourth wave (Table 5). The NROY space estimated at the end of Wave N is defined using the implausibility which depends on the

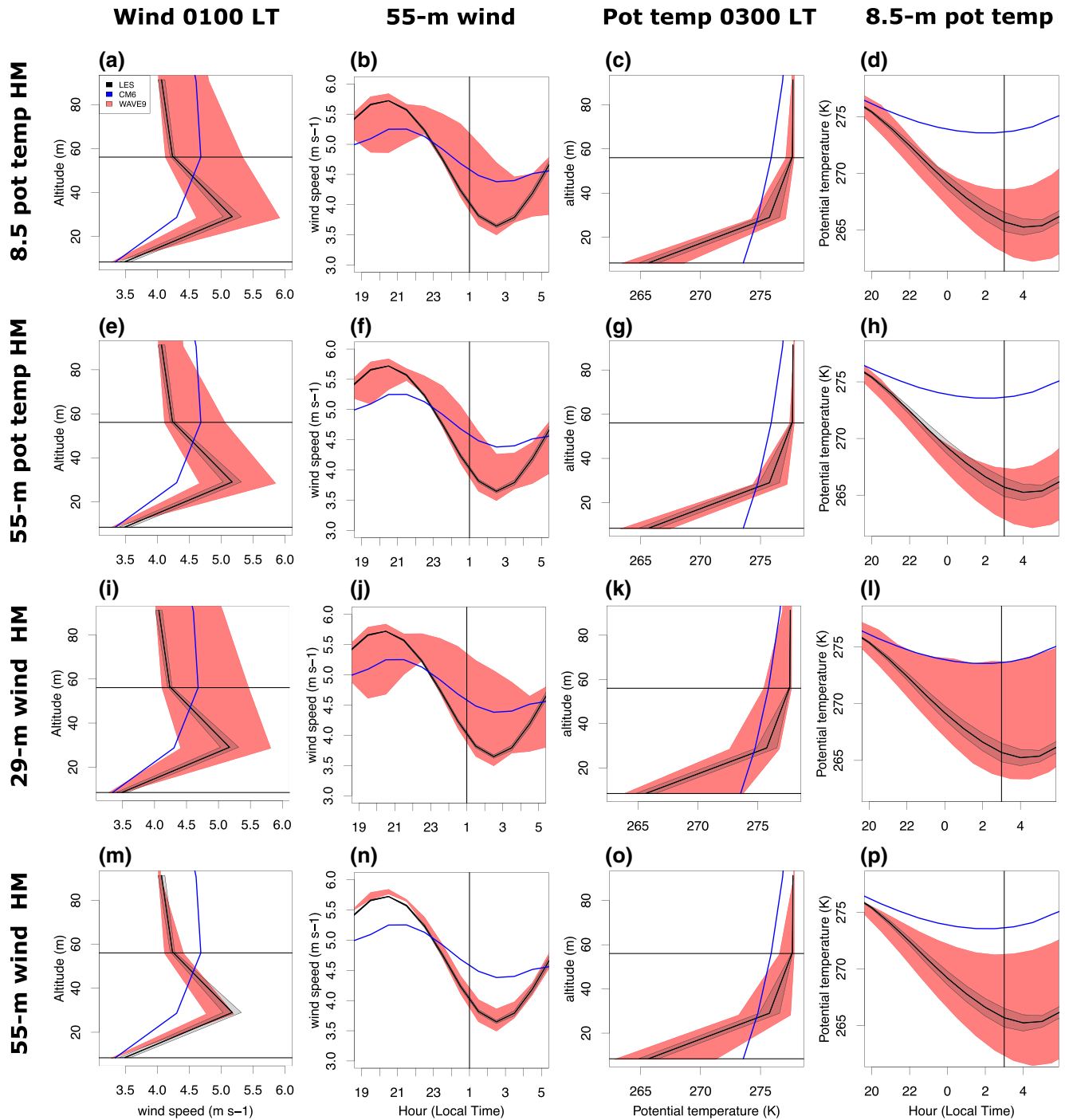


Figure 6. Results of the SCM-LR-SHF history matching experiment, but considering only (a to d) the $\theta_{8.5m}$ metric, (e to h) the θ_{55m} metric, (i to l) the w_{29m} metric and (m to p) the w_{55m} metric. The first column corresponds to the wind speed profile at 0100 LT ($m s^{-1}$), the second column to the wind speed time evolution at 55 m ($m s^{-1}$), the third column to the potential temperature profile at 0300 LT (K) and the fourth column to the 8.5-m potential temperature time evolution (K). On each panel, the blue line indicates the CM6-LR-SHF simulation, the black line the LES ensemble mean with plus or minus one standard deviation as the gray shading and the red shading the Wave 9 simulation envelope. LES, large-eddy simulation; SCM, single-column model.

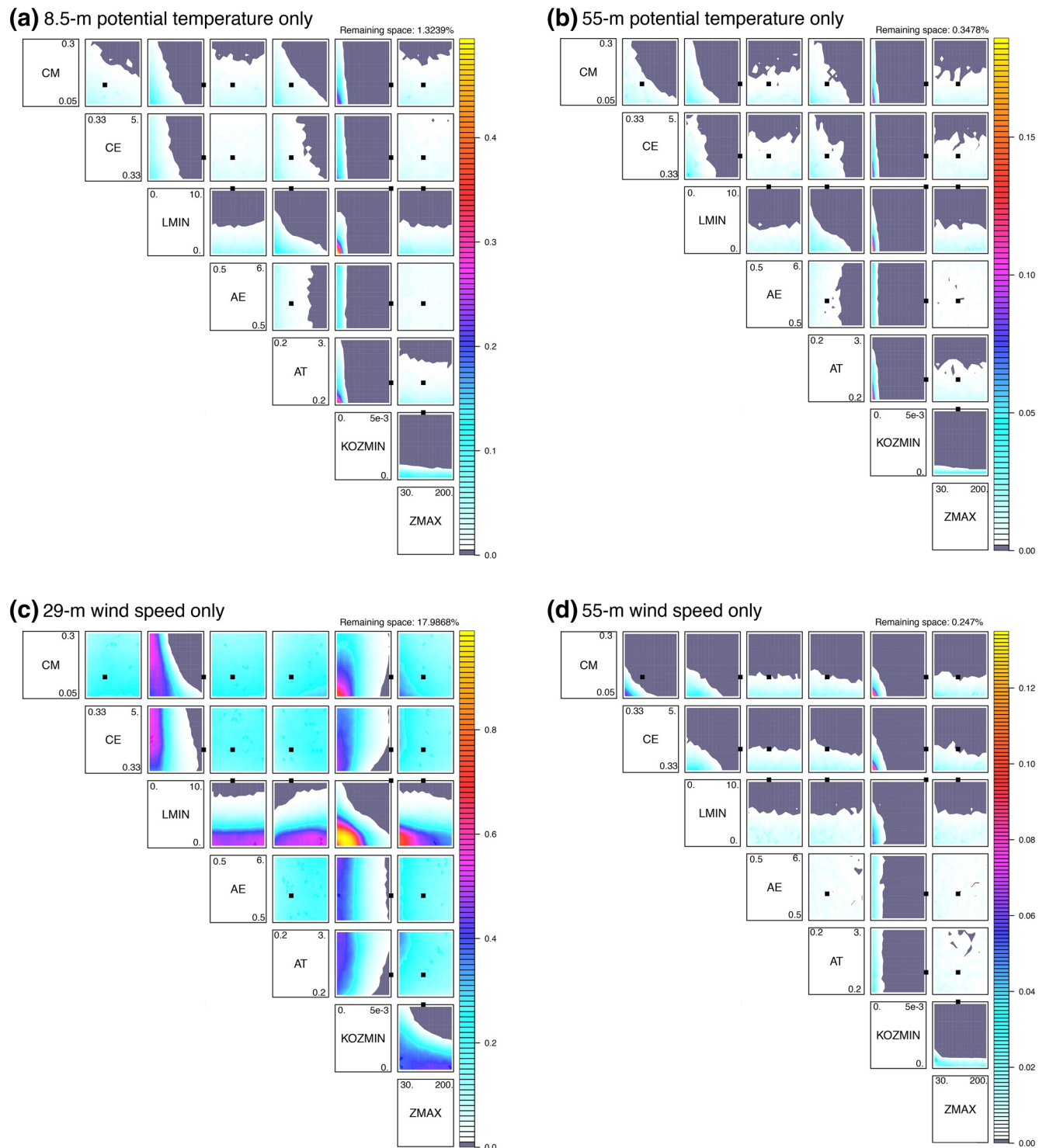


Figure 7. Same as Figure 5, but for the history matching applied to the SCM-LR-SHF configuration considering only (a) the $\theta_{8.5m}$ metric, (b) the θ_{55m} metric, (c) the w_{29m} metric, and (d) the w_{55m} metric. On each panel, the NROY space is displayed after 9 waves. NROY, not-ruled-out-yet; SCM, single-column model.

Table 5
Evolution of the NROY Space Fraction (With Respects to the Input parameter Space) for the History Matching Experiments Applied to the SCM-HR-SHF and SCM-LR-SHF Configurations

Experiment	SCM-HR-SHF	SCM-LR-SHF
NROY ¹	3.0%	0.59%
NROY ²	0.97%	0.17%
NROY ³	0.52%	0.14%
NROY ⁴	0.39%	0.12%
NROY ⁵	0.27%	0.11%
NROY ⁶	0.21%	0.10%
NROY ⁷	0.16%	0.09%
NROY ⁸	0.14%	0.08%
NROY ⁹	0.12%	0.08%

NROY, not-ruled-out-yet; SCM, single-column model.

metric value estimated using the surrogate model, the different sources of uncertainty and the chosen threshold. Figure 8 presents the evolution through the different waves of the implausibility distribution (third row), computed for each of the four metrics considered here, as well as the different quantities involved in its computation. For Wave N , the implausibility is computed using the points in the Wave $N-1$ NROY space. From Wave 3, and for each metric, almost all points in the NROY space have an implausibility lower than 3 (the chosen cutoff, cf. Section 4) despite the reduced emulator uncertainty within the NROY space compared to the previous wave.

As the NROY space fraction reduces during the successive waves, the 70 SCM runs sample a much smaller space, thus leading to improved surrogate models within the NROY space, with reduced uncertainty. This is particularly the case for the θ_{55m} and w_{55m} metrics, for which from Waves 2–3, the surrogate model uncertainty (or variance) falls mostly below or is of the same order of magnitude as the reference uncertainty (Figures 8f and 8h). For $\theta_{8.5m}$, the surrogate model uncertainty is also reduced after Wave 1 (Figure 8e), but there is not a strong decrease of it. It is already significantly lower than the reference uncertainty and thus does not play

much in the implausibility for this metrics. Besides, as the NROY space does not anymore reduce dramatically from Wave 2 onwards, the new SCM runs mostly sample the same space and thus no surrogate model improvement is expected.

The different metrics estimated by the emulators also converge rapidly and from Wave 3 (Wave 5 for w_{55m}) onwards (Figures 8a–8d), only a few outliers fall outside the range of 3 standard deviations around the reference. In other words, from Wave 5 onwards, the simulations are consistent with the LESs given our four metrics, and taking into account the LES uncertainty. To go further in the calibration process, the threshold initially taken to three could be changed to a lower value, especially because the emulator uncertainty is not anymore the dominating uncertainty.

The analysis of the metric convergence also emphasizes the question of discrepancy. If for $\theta_{8.5m}$, θ_{55m} , and w_{55m} , the SCM simulations converge to the mean reference, this is not the case for w_{29m} . For this metric, the LES ensemble mean is about 5.2 m s^{-1} , while the SCMs struggle to exceed 5.0 m s^{-1} . If a non-null discrepancy is not necessary to avoid an empty space, the latter point advocates for taking a structural error of the model at least equal to 0.2 m s^{-1} . Our choice for no tolerance to error seems reasonable for exploring a single case study and only a few metrics. It is however not recommended in a more comprehensive calibration of the full model as it is likely to end up with overtuning (Williamson et al., 2017). When no information can help in suggesting or quantifying this discrepancy, tests during the first wave (when computing the implausibilities and identifying the NROY space) are likely to provide some upper bound on it. The following waves will possibly help to reduce it as the modeler gains a better quantification of his model behavior and can further confront what he wants and what the model can really do.

Finally, we want to stress here the importance of analyzing the comparative importance of the different sources of uncertainty when the procedure seems to converge (i.e. when the space of the parameters no longer reduces significantly from one wave to the next) before reducing the threshold for example.

6.3. Evaluation of the Statistical Framework: Comparison with a 100% SCM Approach

As SCM low-resolution simulations are computationally cheap, it is possible to perform an independent large ensemble of simulations to assess the overall framework used in Section 4, in particular, the quality of the GP-based surrogate models. In the present section, we focus on the SCM-SHF-LR setup for which 10^4 simulations are performed. The parameters are sampled according to a conventional Latin hypercube of the input space. The implausibility of each of these 10^4 simulations is evaluated. As there is no emulator used in this experience, Equation 13 reduces to

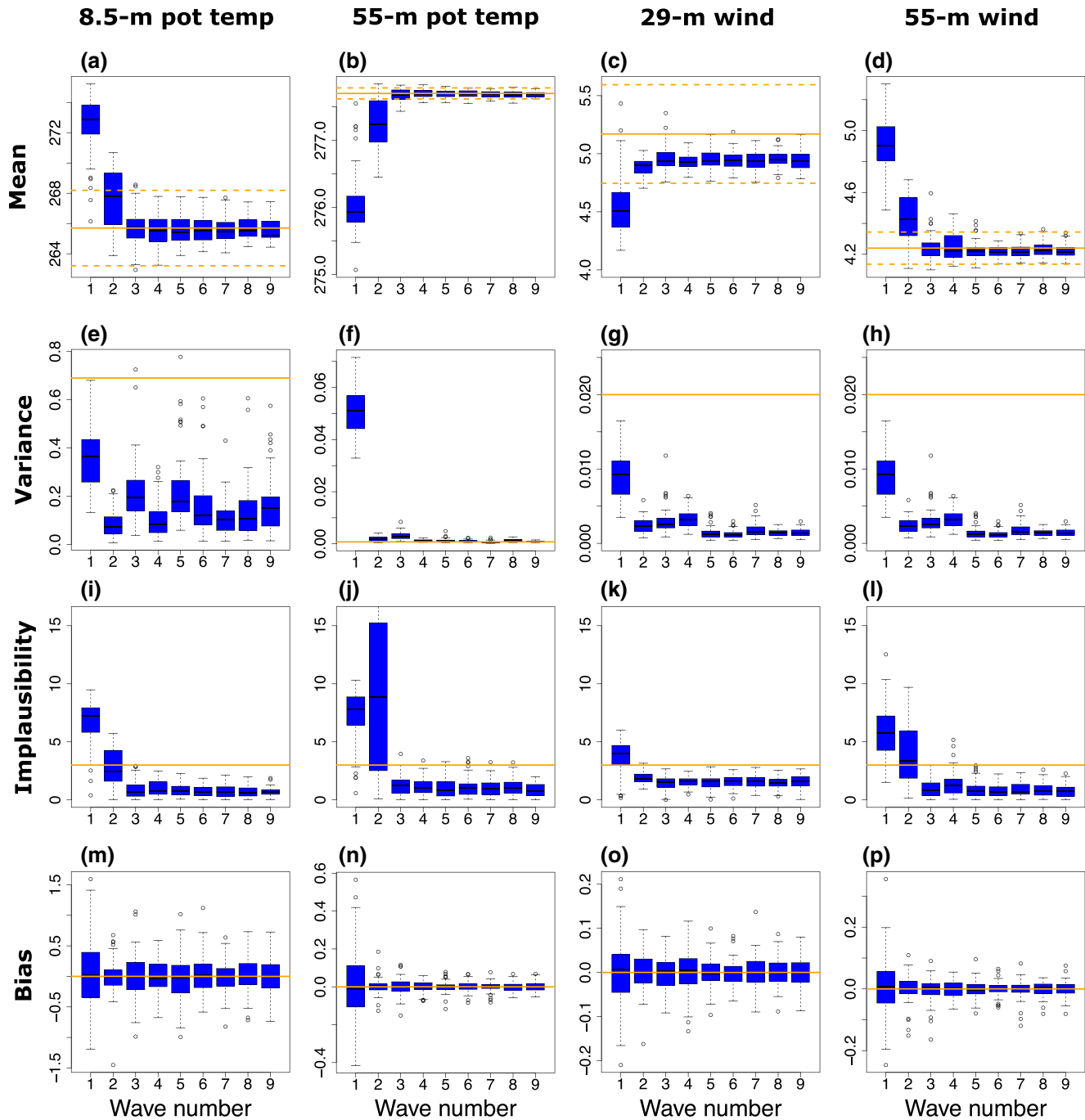


Figure 8. Evolution across the SCM-LR-SHF 9 waves of the distribution of various variables involved in the implausibility computation: (a to d) metrics values with the LES ensemble mean (orange solid line) and plus or minus three standard deviation (dashed orange lines), (e to h) surrogate model variance with the LES ensemble variance (orange solid line), (i to l) implausibility with the cutoff of 3 (orange solid line), and (m to p) the GP-based emulator bias with respect to the SCM ensemble mean of the corresponding wave. The first column corresponds to the $\theta_{8.5m}$ metric, the second column to the θ_{55m} metric, the third column to the w_{29m} metric and the fourth column to the w_{55m} metric. The distributions are computed using points sampling the NROY space obtained at the end of the previous wave (input space for Wave 1). LES, large-eddy simulation; NROY, not-ruled-out-yet; SCM, single-column model.

$$I_f(\lambda) = \frac{|r_f - f(\lambda)|}{\sqrt{\sigma_{r,f}^2 + \sigma_{d,f}^2}} \quad (16)$$

where $f(\lambda)$ is the value of the metric f for the set of parameters λ , directly computed from the SCM simulations. As in the previous history matching experiment and in order to compare results, the structural error $\sigma_{d,f}^2$ is taken to zero. This formulation is used to estimate the NROY space, directly from the 10^4 SCM ensemble. Note that to appropriately characterize the NROY space, it is necessary that it contains enough points, at least a few hundreds. The NROY space obtained using the four metrics in Section 5.2 is about 0.1% of the input space. It has been numerically evaluated using the implausibility computed over 10^6 points, so that it contains in the end around 10^3 points. With only 10^4 points in the input space, only an order of 10 points is expected to remain, which is clearly not sufficient to fully characterize the NROY space and makes its estimate much sensitive to the input space sampling. This further emphasizes the relevance of using surrogate models, even in this SCM framework. Therefore, the statistical framework is evaluated using only the w_{29m} metric, as it a priori keeps a sufficient fraction of the initial space, namely about 18% using emulators (thus about 1,800 points). In parallel, a NROY space is estimated using the emulators built in Section 5.2 on the same 10^4 points in the parameter space. The use of the same points for both estimates avoids sampling effects in the comparison. In experiments using emulators, the threshold for implausibility is set at 3. This choice is deliberately conservative and reduces the risk of ruling out a set of parameters that would in fact leads to results consistent with the reference. The threshold of 3 follows the 3- σ rule of Pukelsheim (1994) which states that 95% of any unimodal distribution lies in the range of $\pm 3\sigma$, σ being the standard deviation of this distribution. To compare the two methods, keeping this threshold of 3 for the direct approach is not relevant, as in our full SCM approach, there is only one source of uncertainty. Assuming a Gaussian distribution for the reference, 95% of the distribution is in the range of $\pm 1.96\sigma$ around the mean. We thus reduce the threshold to 1.96 in the full SCM approach when ruling out parameter values as this provides a more consistent comparison with the GP-based framework.

Figure 9 compares the NROY spaces for the w_{29m} metric, obtained either directly from the 10^4 SCM simulations (NROY_{SCM}) or with Wave 9 emulator ($\text{NROY}_{\text{GP}}^9$). NROY_{SCM} is significantly smaller than $\text{NROY}_{\text{GP}}^9$ (9.8% against 18.0%), but the two NROY spaces are highly consistent. This comparison clearly validates the statistical framework used in Section 5 and the quality of the surrogate models in predicting the metric and its uncertainty.

7. Conclusion

Stable boundary layers are still critical features for weather and climate models. In the present work, we seek to assess whether these model deficiencies reflect calibration choices, or whether they are more deeply rooted in the formulations and implementations of the turbulence parameterization themselves. In the latter case, this would clearly point to intrinsic limits of current parameterizations and possibly to missing processes key for stable boundary layers. To address this question, we took the example of the CNRM atmospheric model, namely ARPEGE-Climat 6.3, which implements the Cuxart et al. (2000) 1.5-order turbulence parameterization, with a few bounds that were historically and empirically added to prevent undesirable model behavior under certain circumstances (e.g., runaway cooling over Antarctica). At this stage, our example solely makes use of a single-column model framework, based on the very stable boundary layer regime of GABLS4 (Bazile et al., 2015). The Large-Eddy Simulation (LES) ensemble analyzed by Couvreur, Bazile et al. (2020a) serves as reference for evaluating and constraining the model behavior. A statistical approach, based on history matching and Gaussian-Process-based surrogate models, is then used to identify whether there exists or not some calibration of the free parameters of the turbulence parameterization, which can provide satisfying results on this GABLS4 case. More precisely, our framework follows the process-based tuning advocated by Couvreur, Hourdin et al. (2020b) to characterize the part of the free parameter space, which leads to SCM simulations compatible with the LES references, given the various sources of uncertainty.

We have addressed this experience using two vertical resolutions, namely the standard ARPEGE-Climat 6.3 vertical resolution and a LES-type vertical resolution (2 m), and using two configurations for the interaction

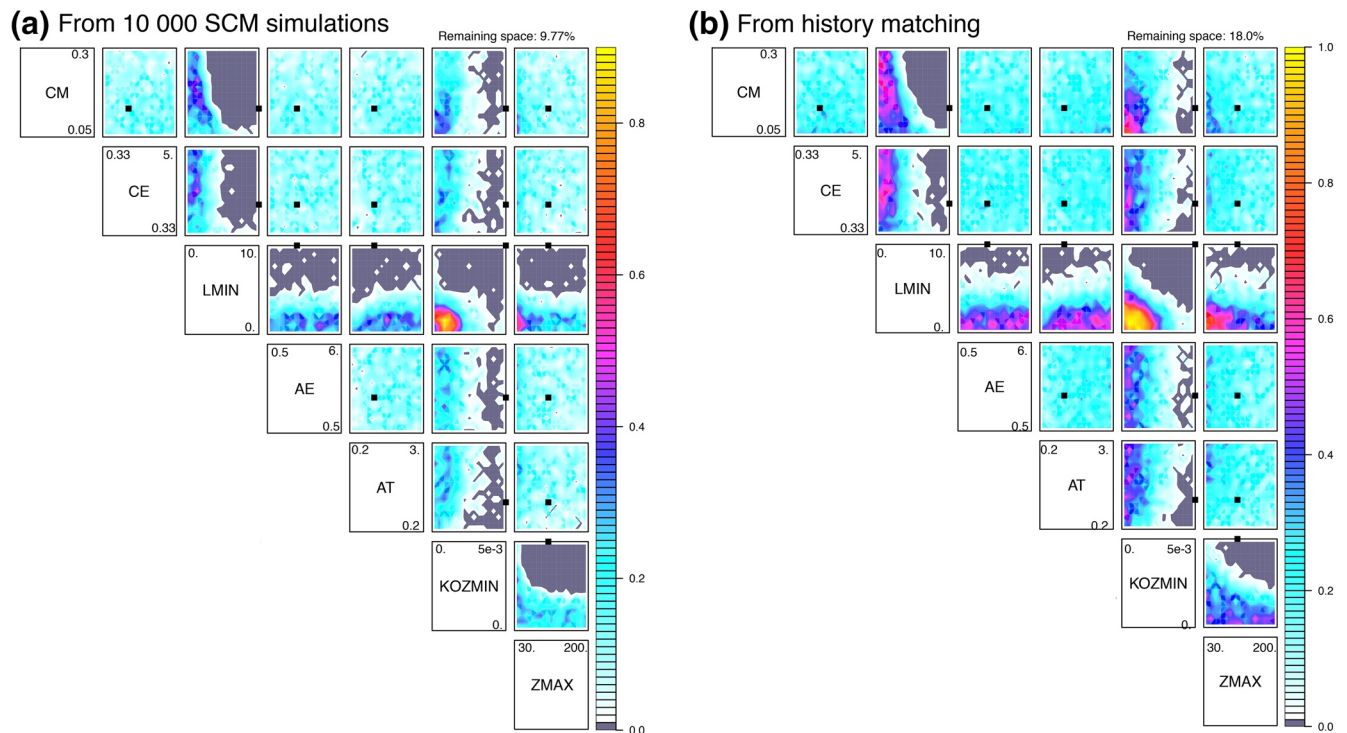


Figure 9. NROY spaces (a) computed from the 10^4 SCM-LR-SHF simulations and (b) computed from the Wave 9 history matching framework applied to SCM-LR-SHF, considering only the w_{29m} metric. See text for details. NROY, not-ruled-out-yet; SCM, single-column model.

with the surface (prescribed surface fluxes and prescribed surface temperatures). Using four metrics (two characterizing the temperature profile, and two characterizing the wind profile), sampled at a given time of the GABLS4 nighttime stable boundary-layer regime, we proved that for each SCM configurations, there exist calibrations of the Cuxart et al. (2000) turbulence parameterization, as implemented in ARPEGE-Climat 6.3, which provide results consistent with the LES references. This indicates in an unambiguous manner that this turbulence parameterization contains sufficient physics to capture strongly-stable boundary layers. As expected, such acceptable model behavior requires calibration that allows to weaken turbulent mixing. This is mostly achieved when strongly reducing the impact of lower bounds historically introduced for maintaining a minimum turbulent mixing (mixing length and minimum flux close to the surface). In contrast, and even though it can be revisited, the calibration of other turbulent parameters can broadly remain consistent with the previous proposals of Cuxart et al. (2000) and Cheng et al. (2002). This importance of lower bounds in the turbulence parameterization clearly echoes similar results obtained by Vignon, Hourdin et al. (2017a) for the climate model LMDZ.

The present work is also the opportunity to gather and formalize our experience with the statistical tools used here and borrowed from the Uncertainty Quantification community (history matching, GP-based surrogate models). As such, we attempt to provide guidance for their use in the context of parameterization and atmospheric model calibration. The importance of the different sources of uncertainty is emphasized. The choice of metrics is an important step that is case-dependent: if the analysis of the standard-calibration model simulation is important by highlighting its main biases, the analysis of the model behavior over the first wave simulations is equally important. It allows to explore, to some extent, what the model can do and where the uncertainties lie. We also illustrate how to understand and tackle the question of the framework convergence by comparing the emulator uncertainties to the other sources of uncertainty.

The present work is a first step in the full calibration of the ARPEGE-Climat atmospheric model. It will be first complemented with the addition of new 1D cases that explore a wide variety of atmospheric regimes. For instance, preliminary new waves that follow the 9 waves performed with the SCM-LR-TS configuration, and which include additional metrics from the first GABLS case (GABLS1, Cuxart et al., 2006), seem to

indicate that the moderately-stable boundary layer of GABLS1 requires a minimum mixing. This clearly contrasts with what we obtained here with GABLS4, thus emphasizing the benefit of using a large diversity of SCM cases (see also Couvreur, Hourdin et al., 2020b). Then, when the calibration performed in the SCM/LES framework has shrunk the space of acceptable parameters, it can continue within this Not-Ruled-Out-Yet space with 3D model configurations. Working in this NROY space constrained with 1D cases ensures that the model behavior at the process-level remains acceptable, while it dramatically reduces the parameter space to be explored with the more expensive 3D configurations. Hourdin, Williamson et al. (2020b) recently provided a first proof of concept of this whole framework, in the context of boundary-layer clouds.

Appendix A: Standard-resolution SCM configuration forced by surface temperature (SCM-LR-TS)

Similar to Section 5.2, this appendix briefly documents the history matching results in the case where the ARPEGE-Climat SCM is forced by the surface temperature (SCM-LR-TS configuration) rather than by the surface sensible heat flux.

The standard calibration SCM version (CM6-LR-TS) behaves similar to CM6-LR-SHF with an underestimated cooling of the atmospheric low levels and a too weak low-level jet (Figures 1 and A1). A notable difference is that the model continues to cool for a longer period (until around 0600 LT). The wind rotation at 29 m remains weakly captured, as in CM6-LR-SHF.

History matching is applied as for SCM-LR-SHF, using the same metrics as in Section 5.2. Nine waves are performed with a similar convergence as for the SCM-LR-SHF configuration (not shown). The Wave 9 SCM simulations are shown with blue lines in Figure A1. Overall, the associated sets of free parameters provide improved and satisfying results over the GABLS4 stable boundary layer, with only a systematic warm bias

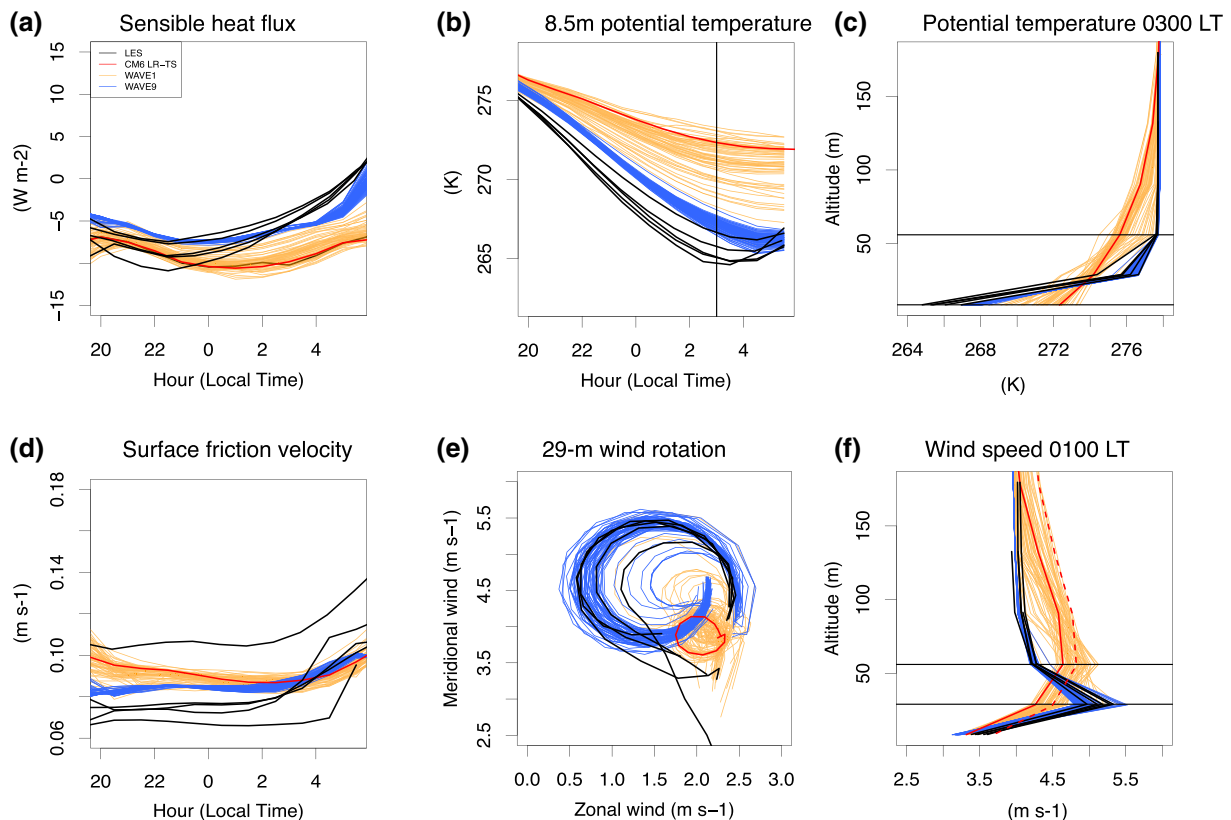


Figure A1. Same as Figure 2 but for history matching applied to the SCM-LR-TS configuration. The standard-calibration simulation CM6-LR-TS is shown with the red line. LES, large-eddy simulation; SCM, single-column model.

SCM-LR-TS NROY - Wave 9

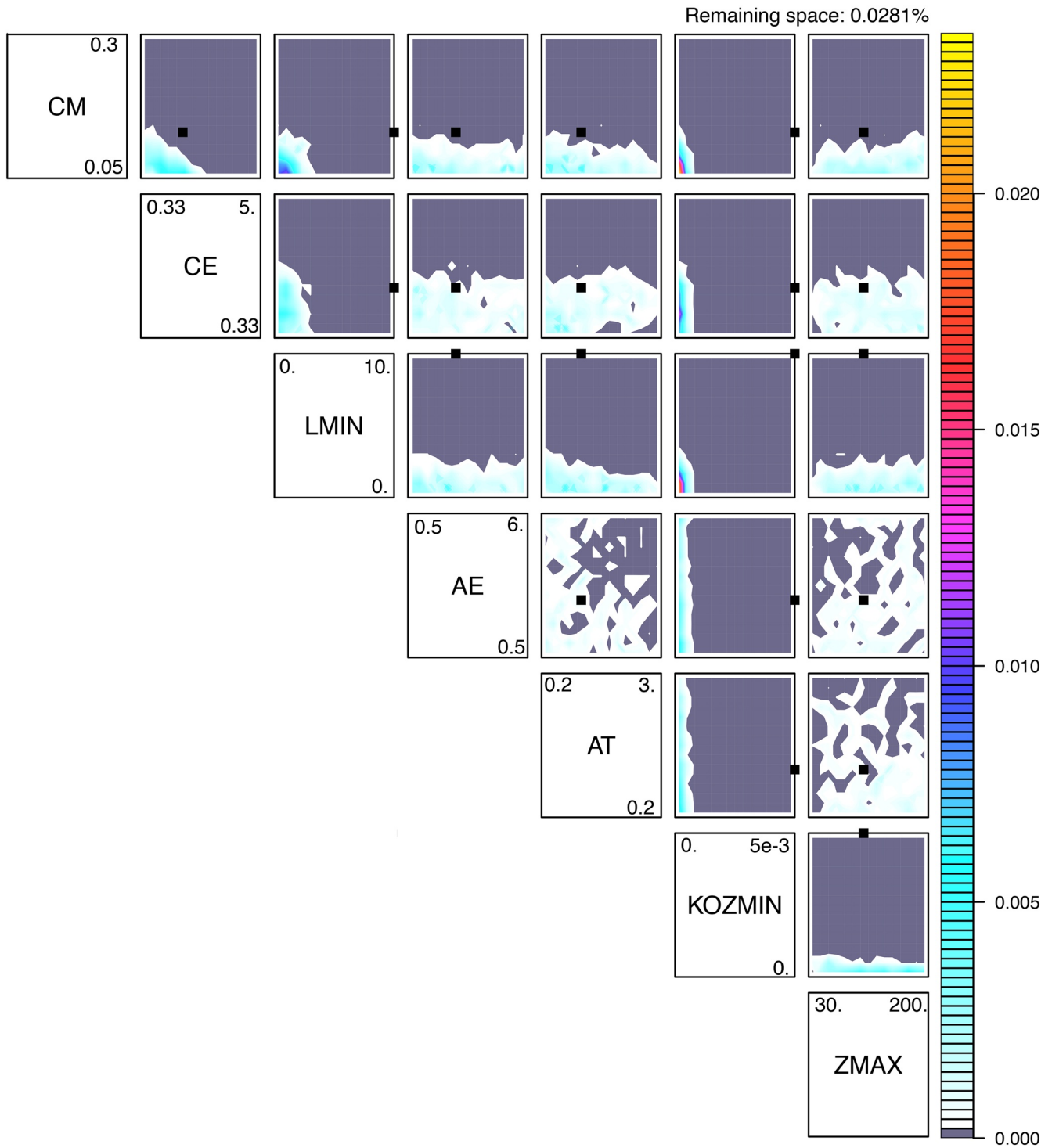


Figure A2. Same as Figure 5, but for the history matching applied to the SCM-LR-TS configuration (NROY space obtained at Wave 9). NROY, not-ruled-out-yet; SCM, single-column model.

at 8.5 m at 0300 LT (1.1–1.3 K), mainly because of a 2-h delay in the SBL cooling. Regarding the wind, the Wave 9 simulations capture well its vertical structure and rotation, within the LES uncertainty (Figures A1e and A1).

Finally, Figure A2 presents the NROY space after the 9 waves for this SCM-LR-TS configuration. The conclusions are similar to those made with the SCM-LR-SHF configuration, with a critical role of **KOZMIN**, **LMIN**, **CM**, and **CE**. For these parameters, values leading to low turbulent mixing are required.

Data Availability Statement

The High-Tune Explorer (htexplo) tool used here for history matching with iterative refocusing is available through the open source version control system “subversion” (svn) at <https://svn.lmd.jussieu.fr/HighTune.org/10.1002/2017MS001077>. A snapshot version of the tool used here is also available at <http://doi.org/10.14768/70efa07b-afe3-43a4-8334-050354f9deac>. LES and SCM data is available at <https://doi.org/10.5281/zenodo.4350522>.

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